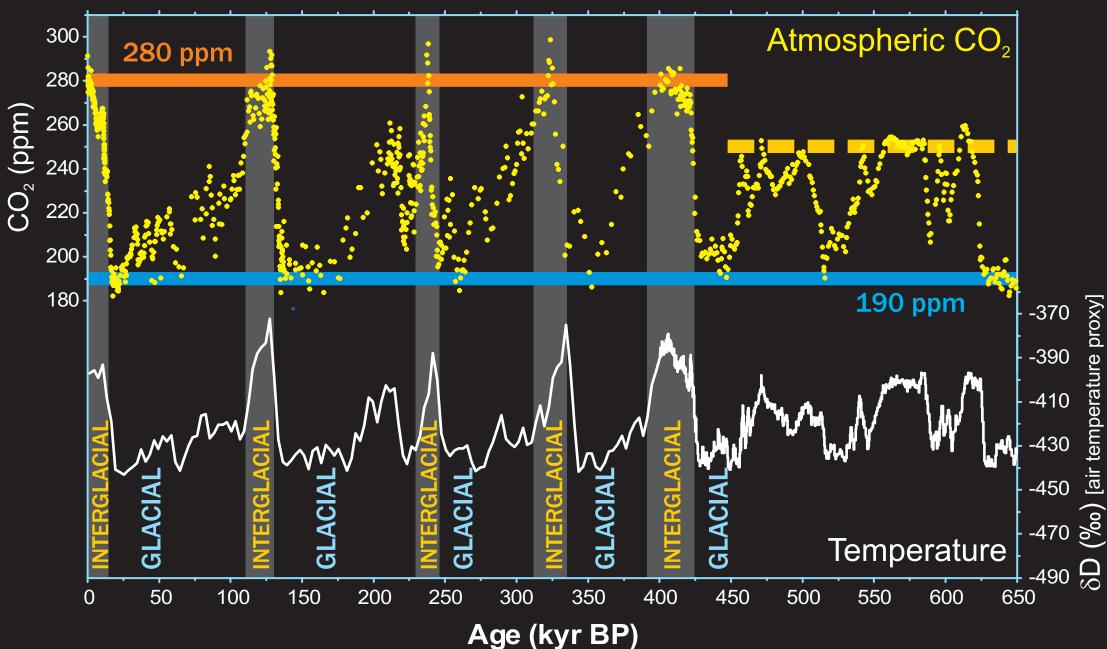
### Classic $CO_2$ (and climate) problems in 'shallow time': Why was glacial $CO_2$ 'low'?

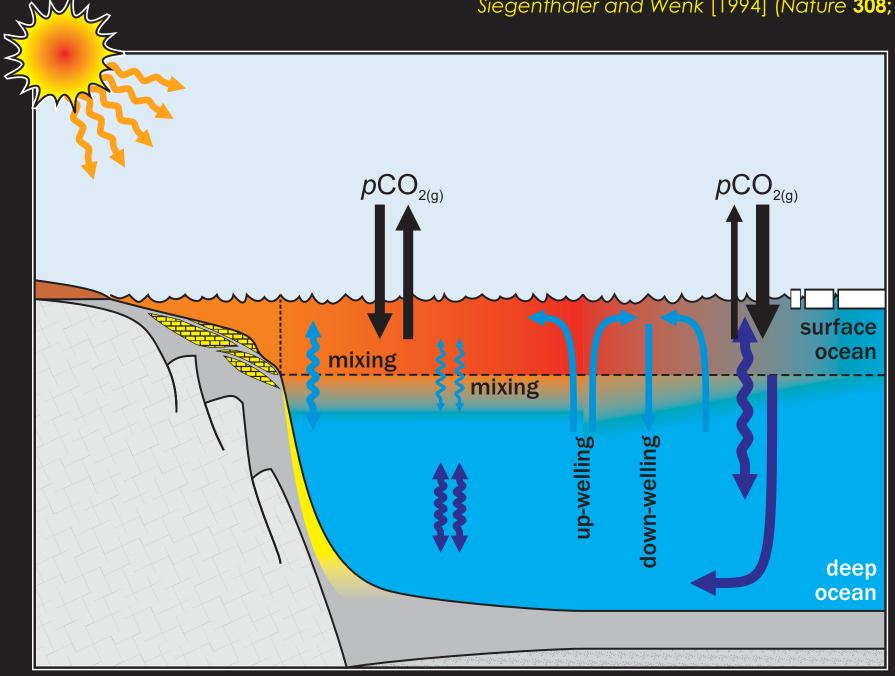
Dome C and Vostok ice core data; Siegenthaler et al. [2005] (Science 310)



CO <sub>2</sub> terms	Glacial-interglacial CO <sub>2</sub> audit	CO <sub>2</sub> (ppm)	LOSU
SST	splinity	+26 (21/30)	High
sea-level	non-salinity effects	-6 (-7/-5) -7 (-9/-6)	High
terrestrial biosphere		-22 (-35/-12)	High
partial linear sum: High LOSU terms		-9 (-30/6)	High
AMOC strength SO stratification whole ocean		12 (7/15) 33 (31/35) 27 (3/57)	Med- Low
iron fertilization		15 (5/28)	Med- Low
metabolic rate		15 (0/30)	Low
coral-reef re-growth		12 (6/20)	
silica fertilization	aeolian	4 (1/8)	Low
wind-speed		0 (0/5)	Med
weathering		0 (-5/5)	Med
sea-ice cover		-5 (-5/0)	Med- Low
partial linear sum: Med/Low LOSU		68 (-4 / 153)	Med/ Low
TOTAL LINEAR SUM OF CO2 TERMS BY TIME-SCALE	'fast' + 'slow' terms 'fast' terms	59 43	
$D_2$ change	glacial→interglacial amplitude (20– deglacial transition (17→11 ka)		
(ppm) -4	0-30-20-10 0 10 20 30 40 50 60 7	0 80 90 1	00

### Lower LGM ocean surface temperatures enhance the solubility of $CO_2$ , increasing the sequestration of $CO_2$ in the ocean interior.

Siegenthaler and Wenk [1994] (Nature 308; Keir [1993] (JGR 98)



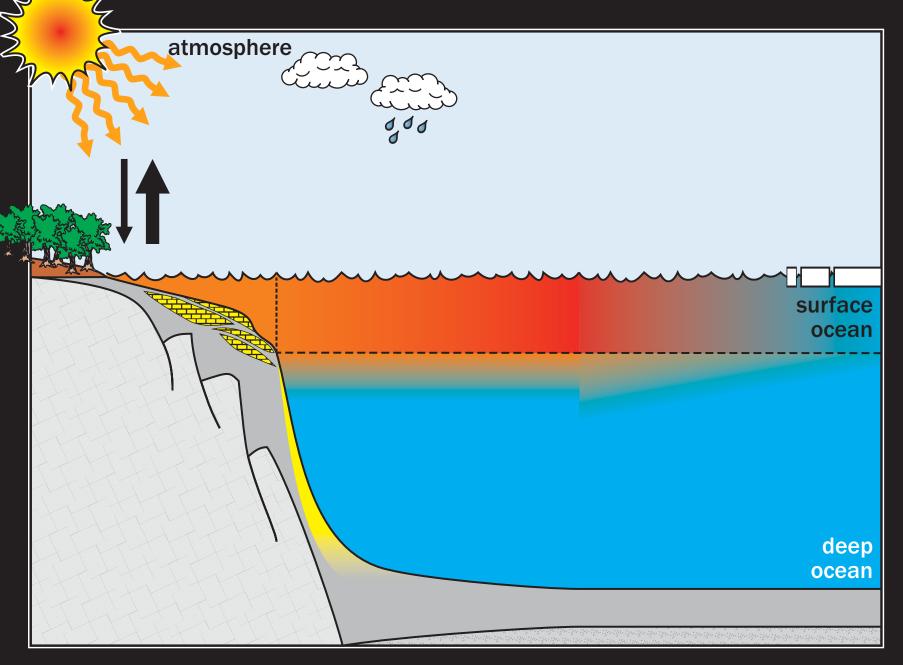
SST



> Change the ocean temperature in the **C-Model** and explore how sensitive atmospheric  $pCO_2$  is to an e.g. ~4°C glacial cooling (relative to the Holocene).

## Terrestrial carbon

A colder, drier glacial climate, and lower pCO<sub>2</sub> and hence reduced CO<sub>2</sub> 'fertilization' of productivity, is likely to have resulted in a reduced inventory of carbon in the terrestrial biosphere.





> Fossil fuel  $CO_2$  emission experiments have been envisaged as the main application of the model but this can be made use of.

- > Add Fossil Fuels
- > Fossil Fuels \*On\*
- > \*Sinewave\*

(avoids complications and additional emissions associate with SRES emissions scenarios)

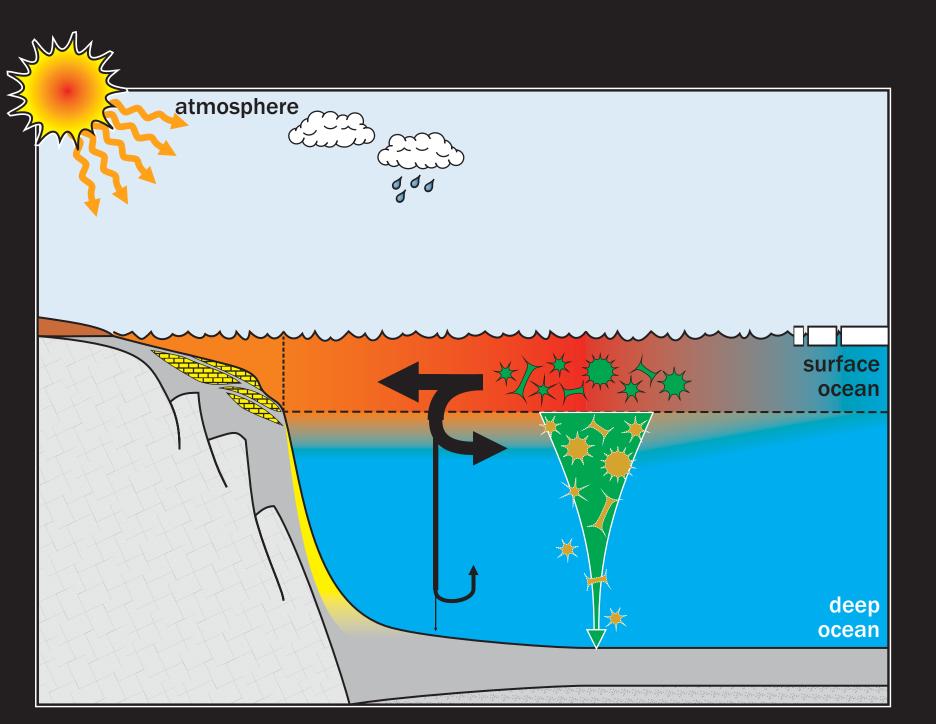
#### > Enter a total amount

> Hint: estimates of  $CO_2$  addition to the ocean/atmosphere at last glacial range from ca. 300-700 PgC.

Q. What constraints might there be on the amount of terrestrial carbon transferred to the ocean+atmosphere?

Q. How much added  $CO_2$  stays in the atmosphere compared with in the ocean? (Is this independent of the amount of  $CO_2$  released?)

# Marine biology and the 'biological pump'





> Other ways of changing ocean
productivity (and the biological
pump) that are plausible and
testable?

- > Model parameters
- > Biological
- > Lower the 'half saturation' value. Be careful not to increase the algal maximum growth rate too high ... you can crash the model ...

> Or:

- > Set Initial Conditions
- > Phosphorous

> \*Surface\*, \*Middle\*, and/or \*Deep\*
Phosphorous

Q. What sort of paleoceanographic constraints can be brought to bear on the model?

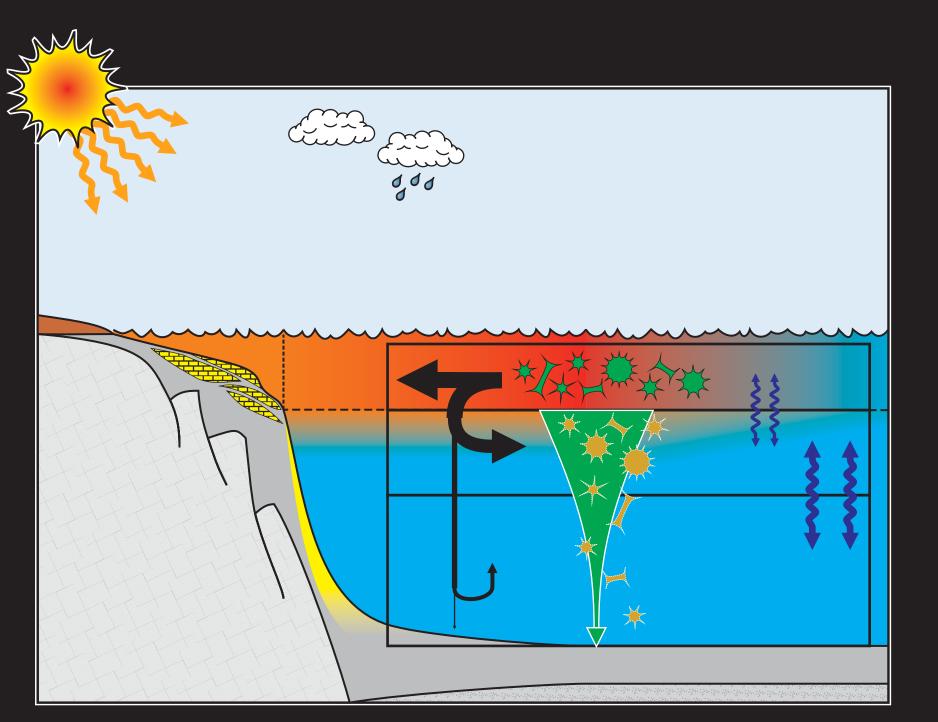


Q. Other further ways of changing the biological pump (efficiency)?

- > Model Parameters
- > Organic Carbon
- > \*Fraction Remineralized\*

> Q. What sort of paleoceanographic constraints can be brought to bear on the model? (And/or ice-core constraints?)

## Ocean circulation





> AMOC, AABW/CPDW production changes
... ?

> What about testing deep stratification? (i.e. the idea that e.g. a deep highly saline layer sat at the bottom of the ocean).

- > Model Parameters
- > Physical
- > \*mixing\* (intermediate to deep)