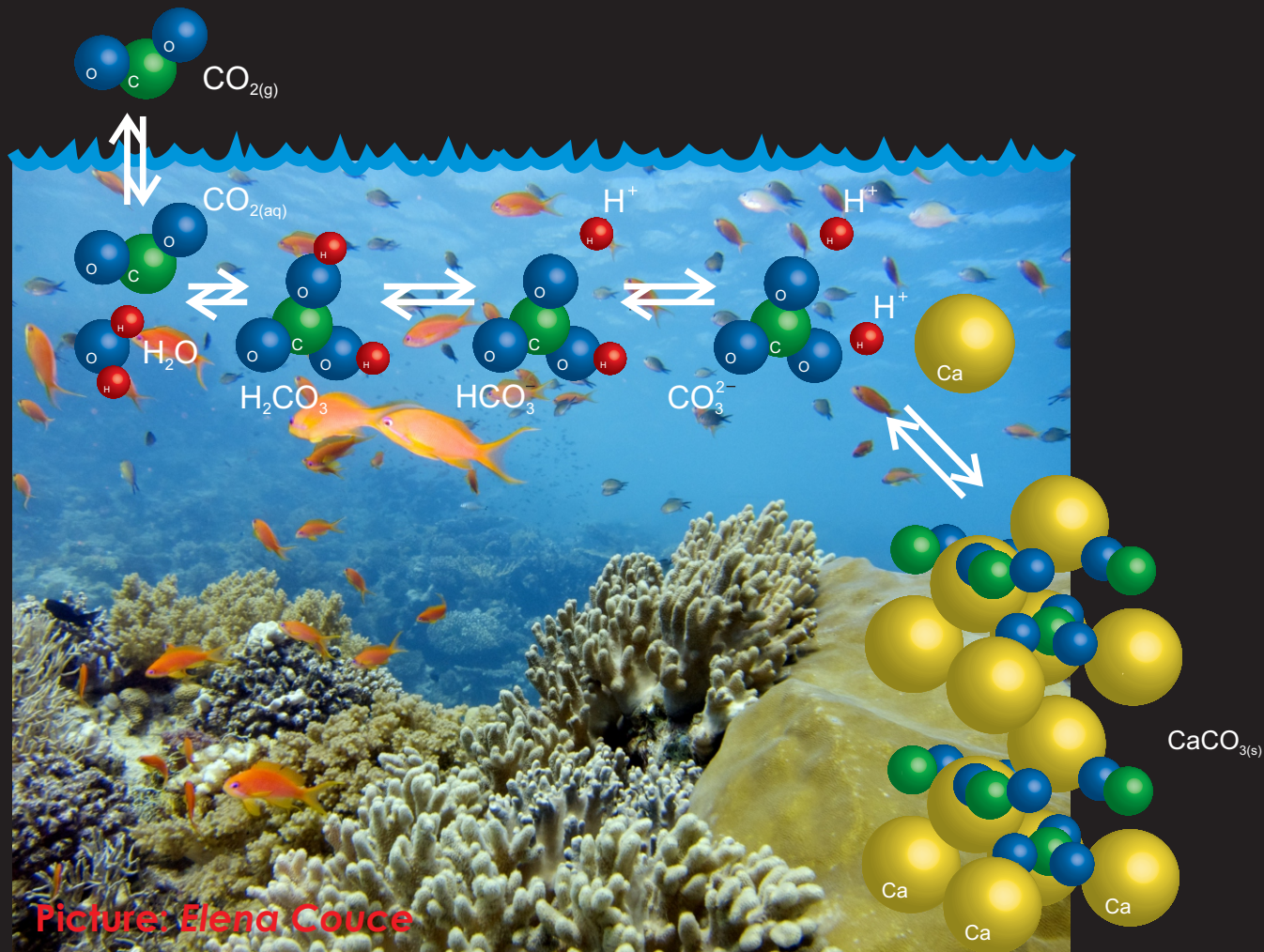
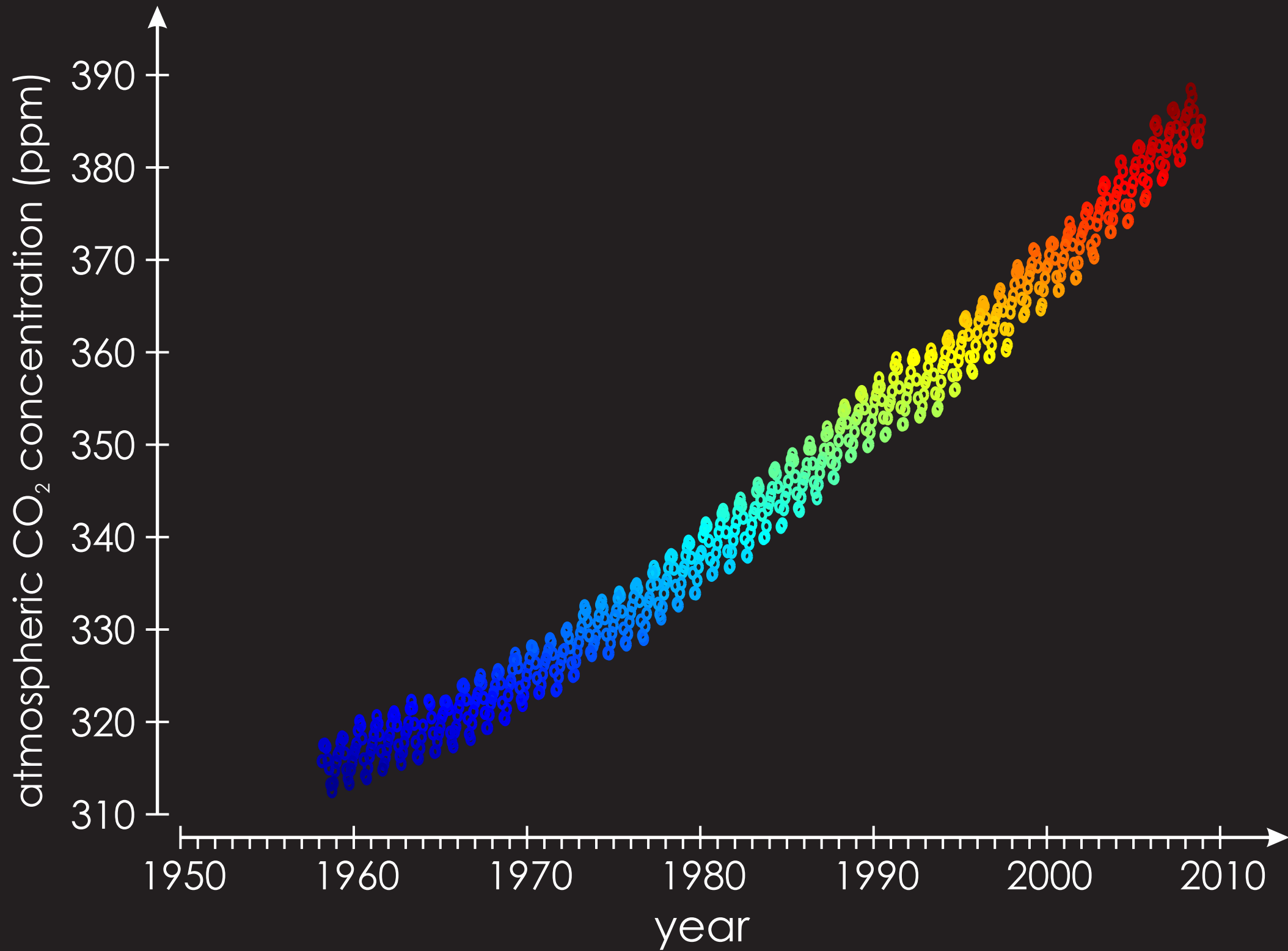


# Ocean Acidification

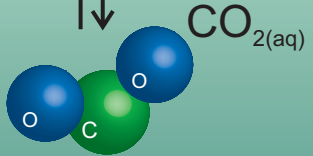
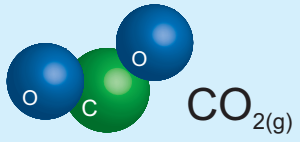


Picture: Elena Couce





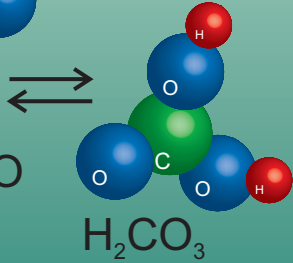
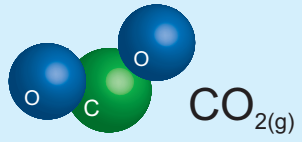
atmosphere



ocean

$\text{CO}_2$  chemistry  
in seawater

atmosphere

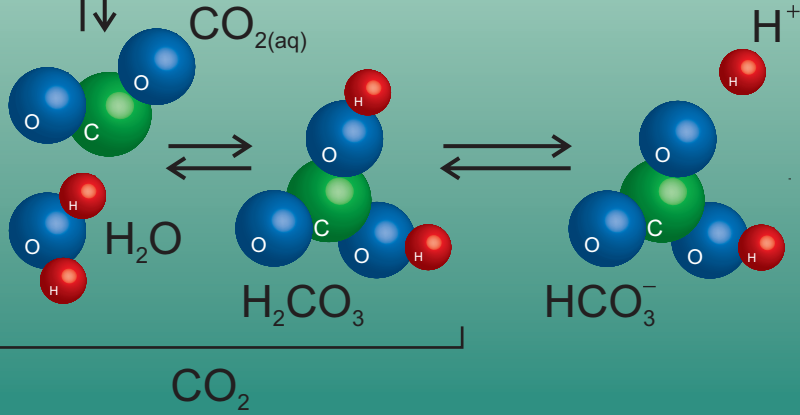
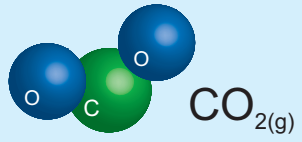


$\text{CO}_2$

ocean

$\text{CO}_2$  chemistry  
in seawater

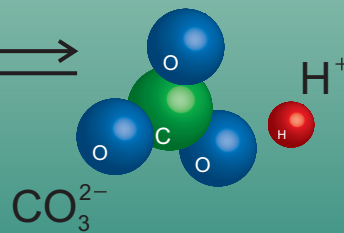
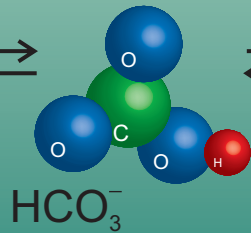
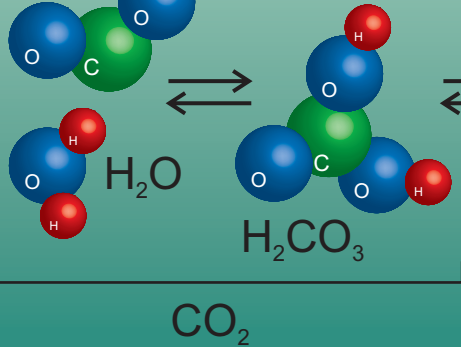
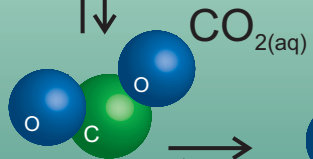
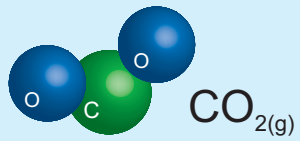
atmosphere



$\text{CO}_2$  chemistry  
in seawater

ocean

atmosphere



ocean

## $\text{CO}_2$ chemistry in seawater

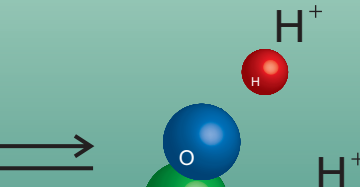
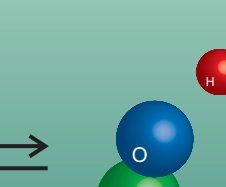
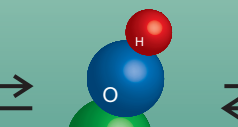
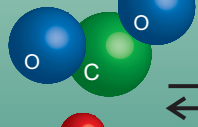
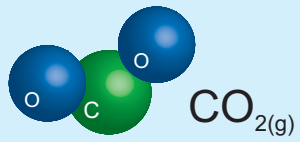
Under typical marine conditions, dissolved carbon dioxide ( $\text{CO}_{2(aq)}$ ) will largely hydrate to form a proton ( $\text{H}^+$ ) and a bicarbonate ion ( $\text{HCO}_3^-$ ), which can dissociate to form another hydrogen plus carbonate ion ( $\text{CO}_3^{2-}$ ).

The sum total;

$\text{CO}_{2(aq)} (+ \text{H}_2\text{CO}_3) + \text{HCO}_3^- + \text{CO}_3^{2-}$   
is collectively termed dissolved inorganic carbon ('DIC').



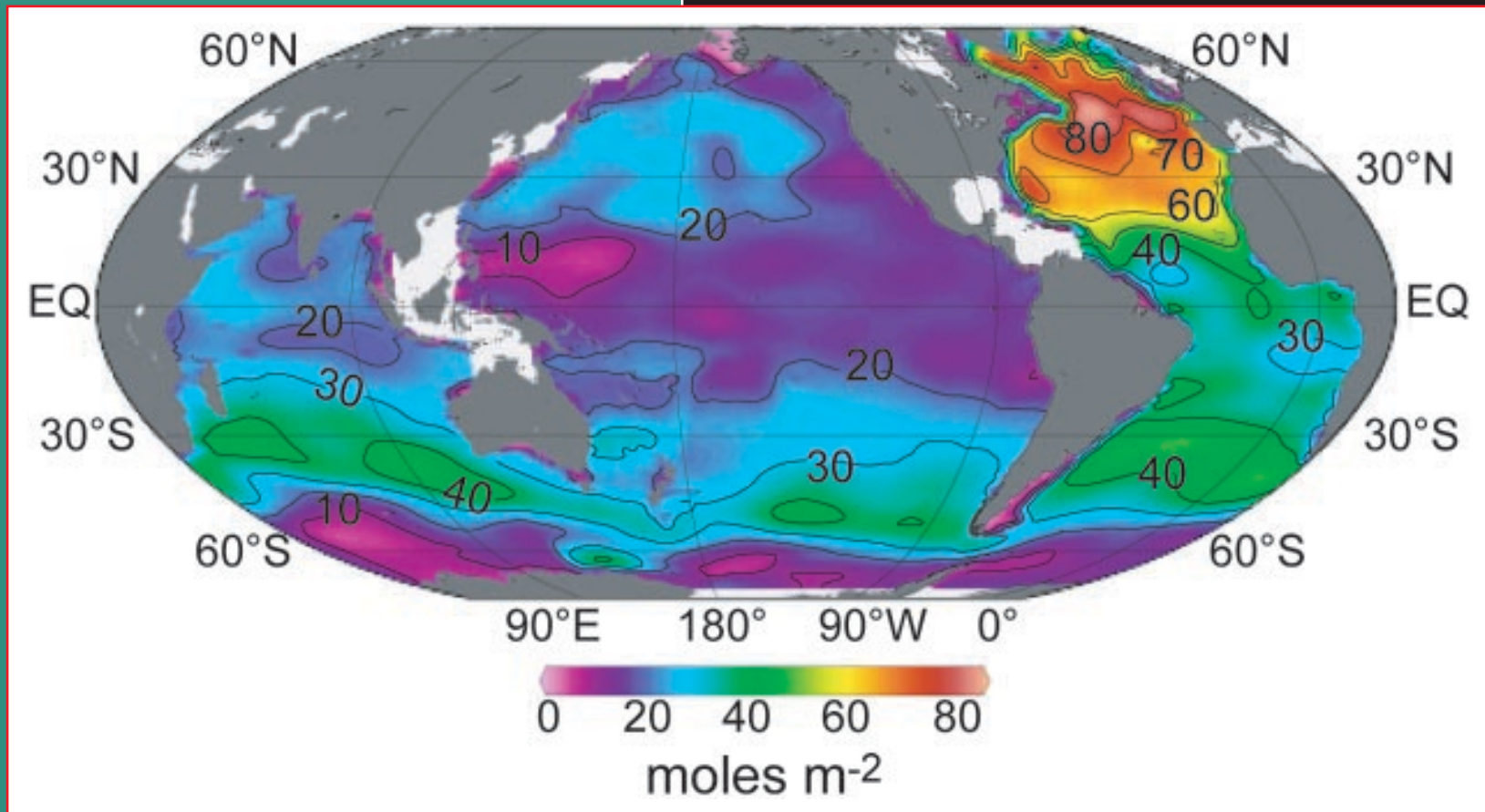
# atmosphere



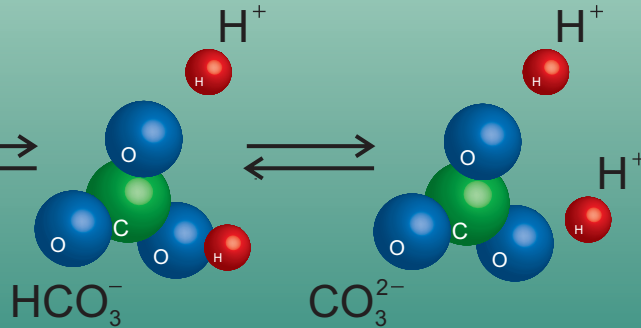
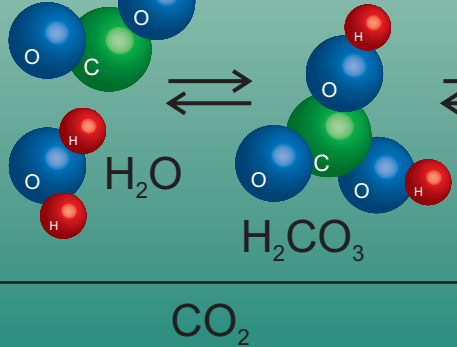
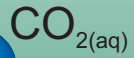
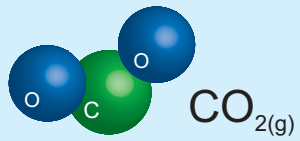
$\text{CO}_2$

# $\text{CO}_2$ chemistry in seawater

From: Sabine et al. [2004] (Science 305)



atmosphere



ocean

## $\text{CO}_2$ chemistry in seawater

When  $\text{CO}_2$  dissolves in seawater, the  $\text{CO}_{2(aq)}$  concentration changes only slightly because the system is buffer by carbonate ions:  $\text{CO}_3^{2-}$

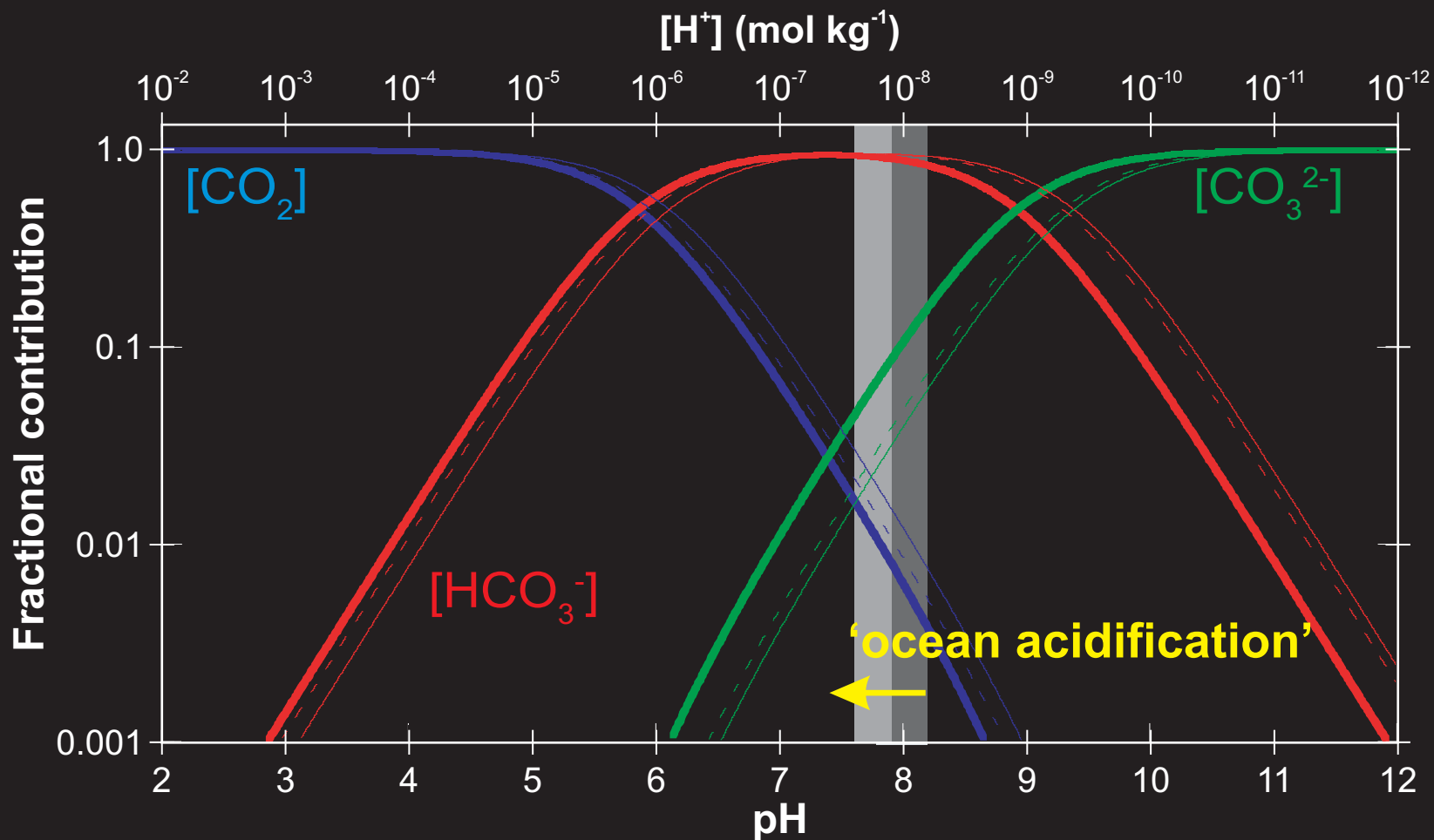
$\text{CO}_2$  is scavenged according to the reaction:



However, a small part of the resulting  $\text{HCO}_3^-$  dissociates into  $\text{CO}_3^{2-}$  and  $\text{H}^+$ , which is where the 'acidification' in ocean acidification comes from.

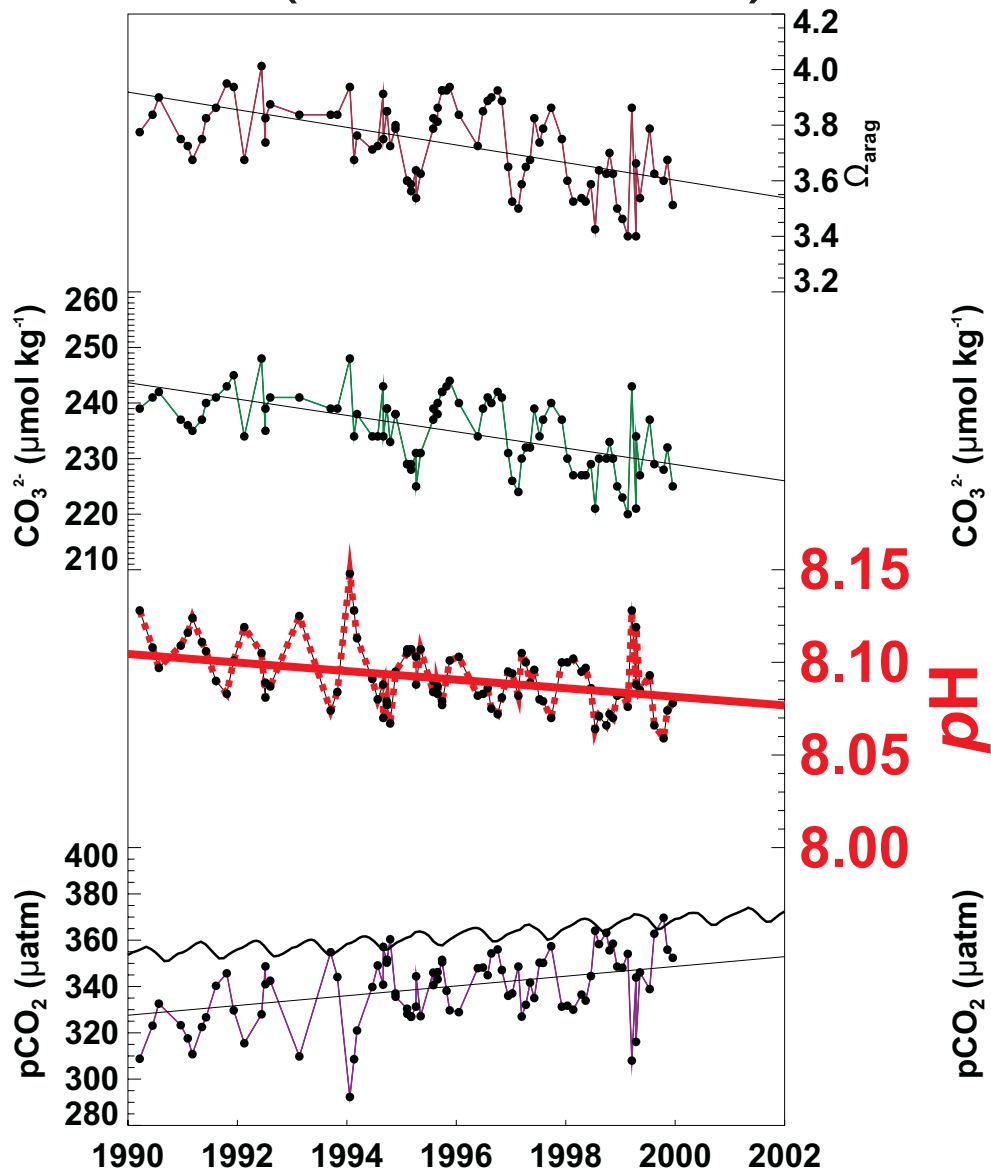
Don't forget that pH is a log scale, so that a 1.0 unit decrease in pH corresponds to a factor 10 increase in hydrogen ion concentration ( $[\text{H}^+]$ ):  $\text{pH} = -\log_{10}([\text{H}^+])$

# pH (and acidity vs. alkalinity)

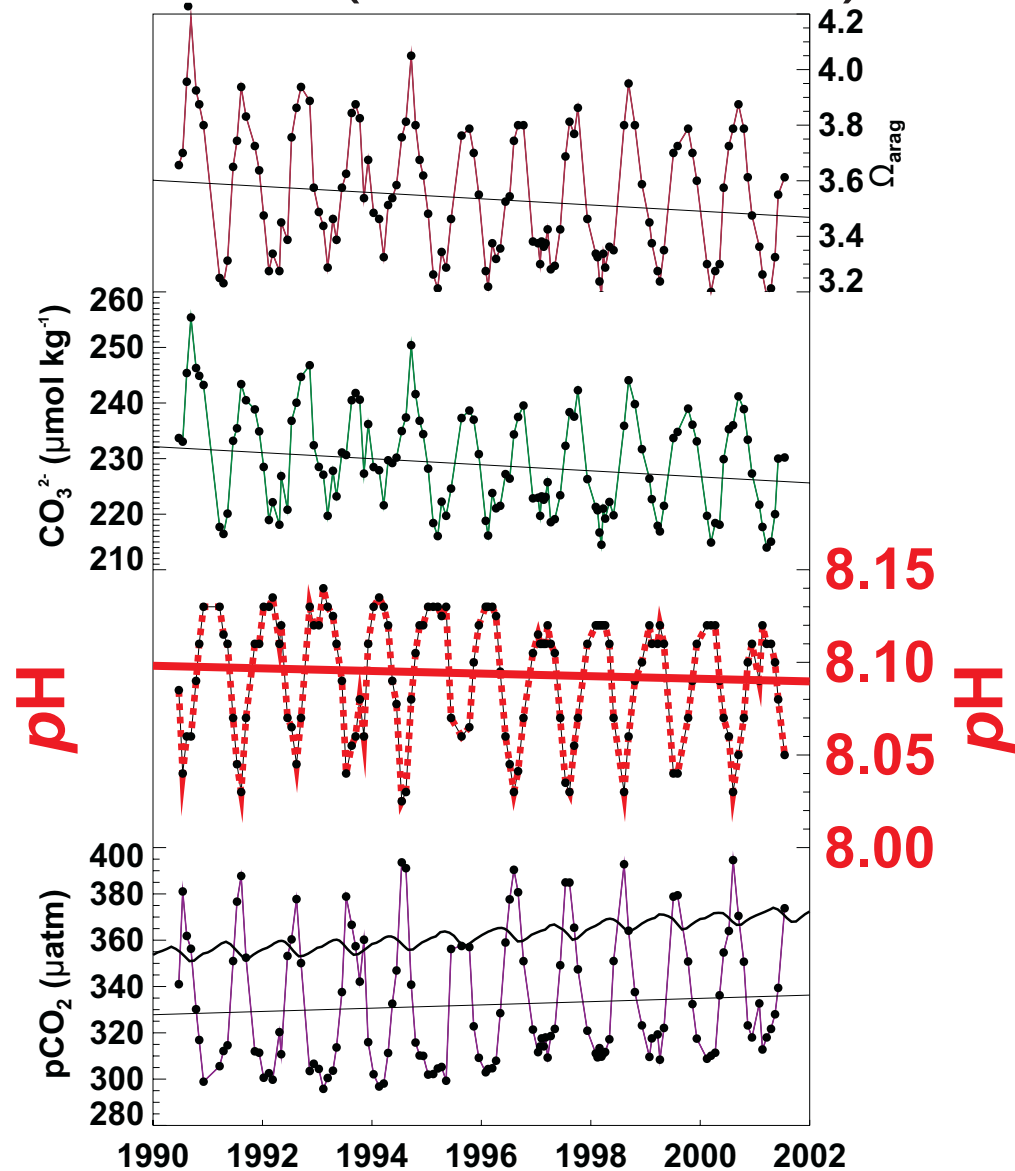


# Observations of declining ocean surface pH

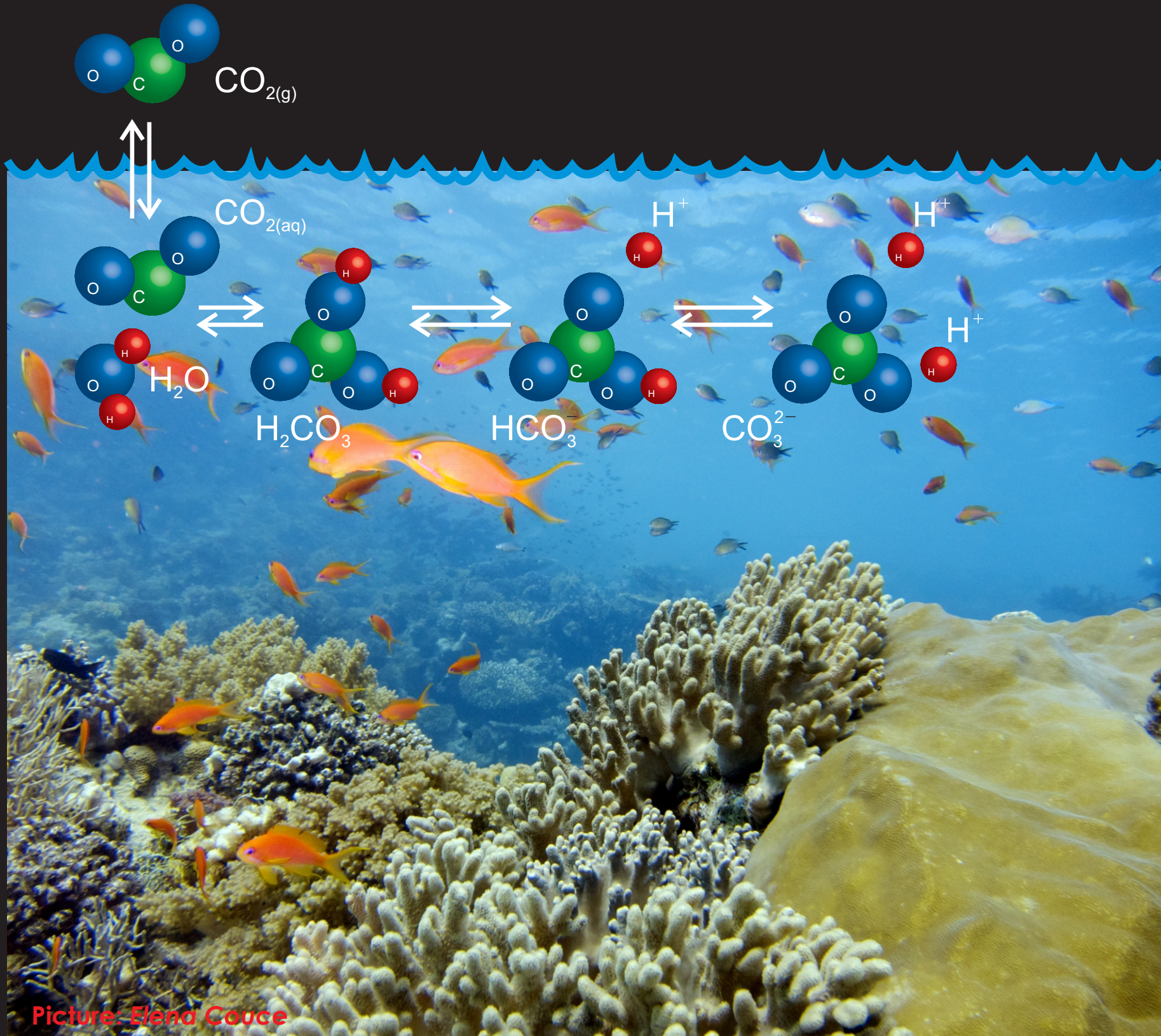
## HOT (Pacific Ocean)



## BATS (Atlantic Ocean)

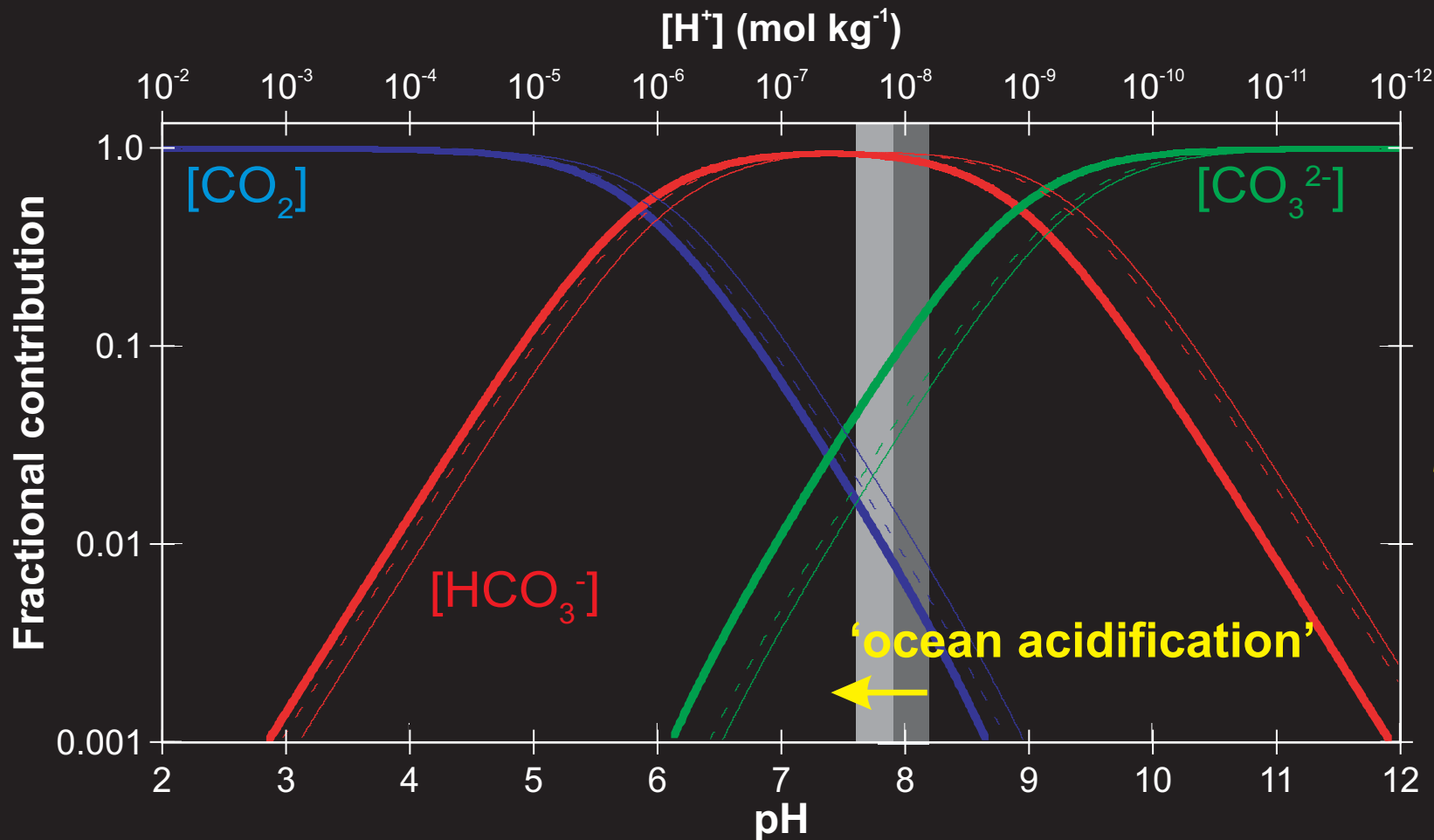






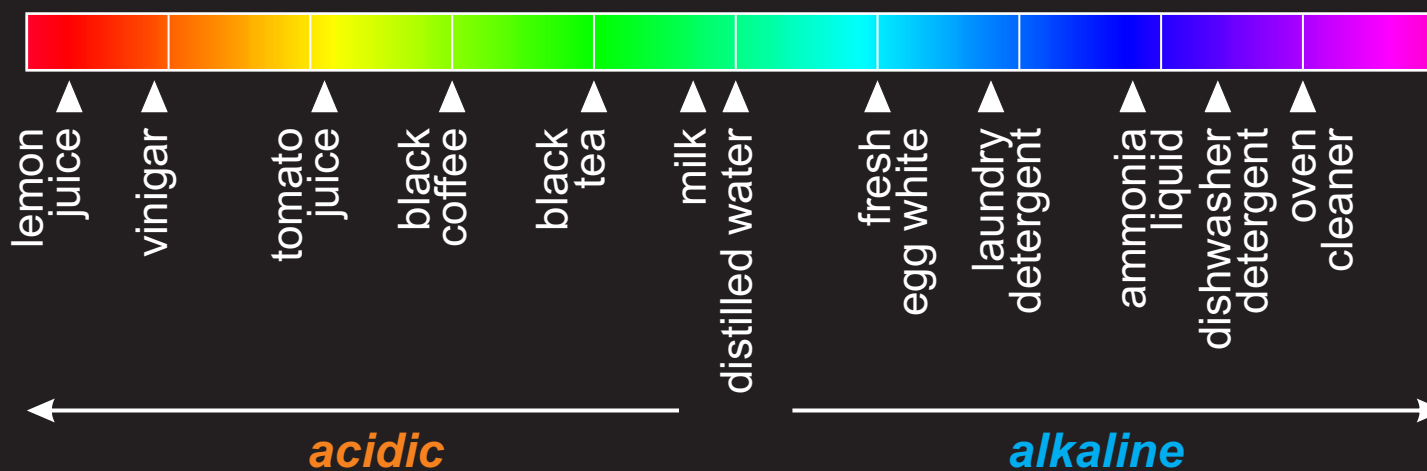
Picture: Elena Couce

# pH (and acidity vs. alkalinity)



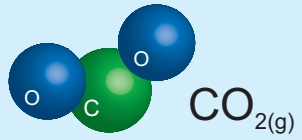
Direct impacts on micro-organisms in the ocean are currently thought likely to be small, although changes in pH will affect nutrient concentrations and trace metal speciation.

Larger (multicellular) animals may be affected as increased  $CO_2$  will acidify the body tissues and fluids and affect the ability of blood to carry oxygen ('hypercapnia').

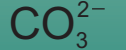
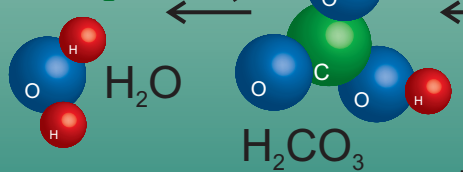
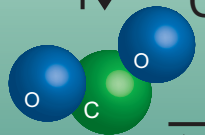




atmosphere



$\text{CO}_{2(aq)}$



Ca

$\text{CaCO}_{3(s)}$

Ca

Ca

Ca

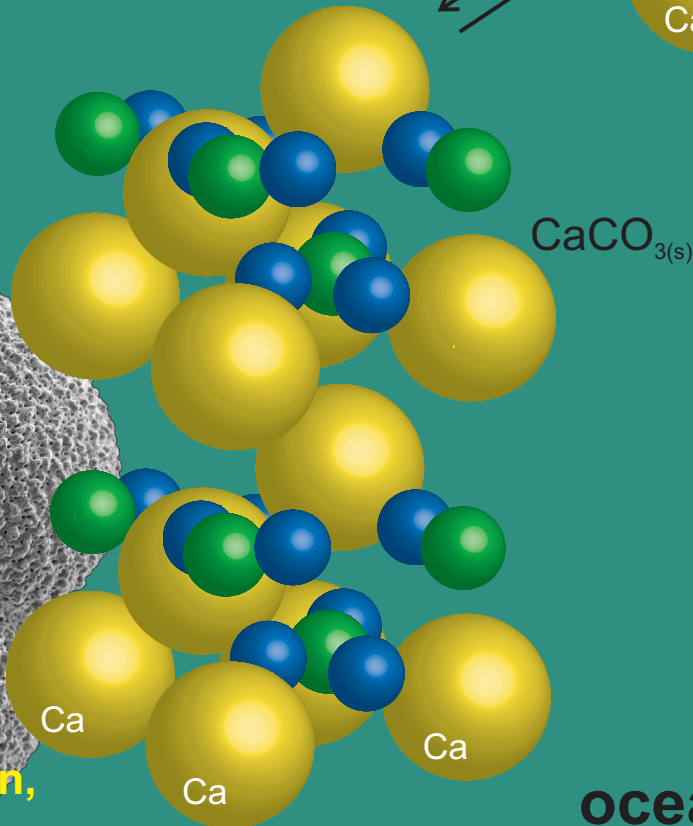
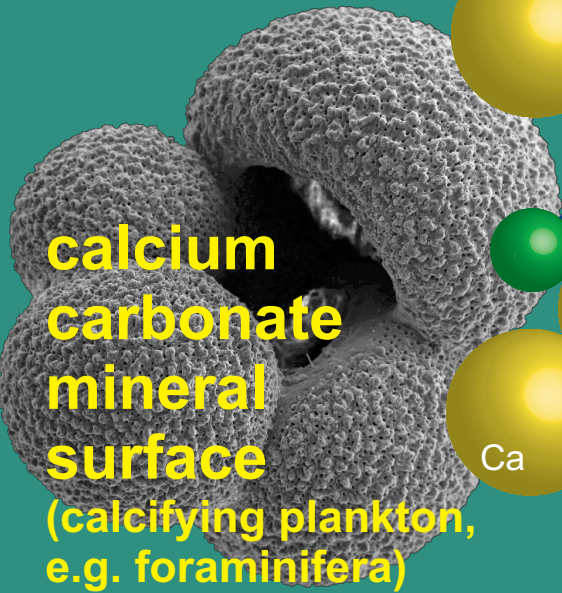
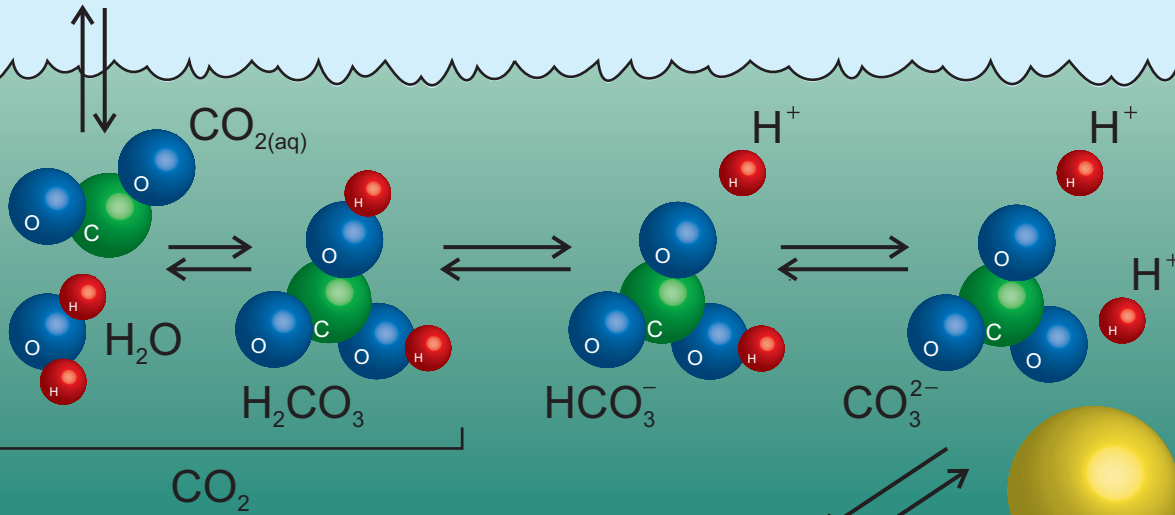
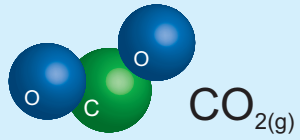
ocean

calcium  
carbonate  
mineral  
surface

(calcifying plankton,  
e.g. foraminifera)

$\text{CO}_2$  chemistry  
& mineral phases

atmosphere



ocean

# $\text{CO}_2$ chemistry & mineral phases



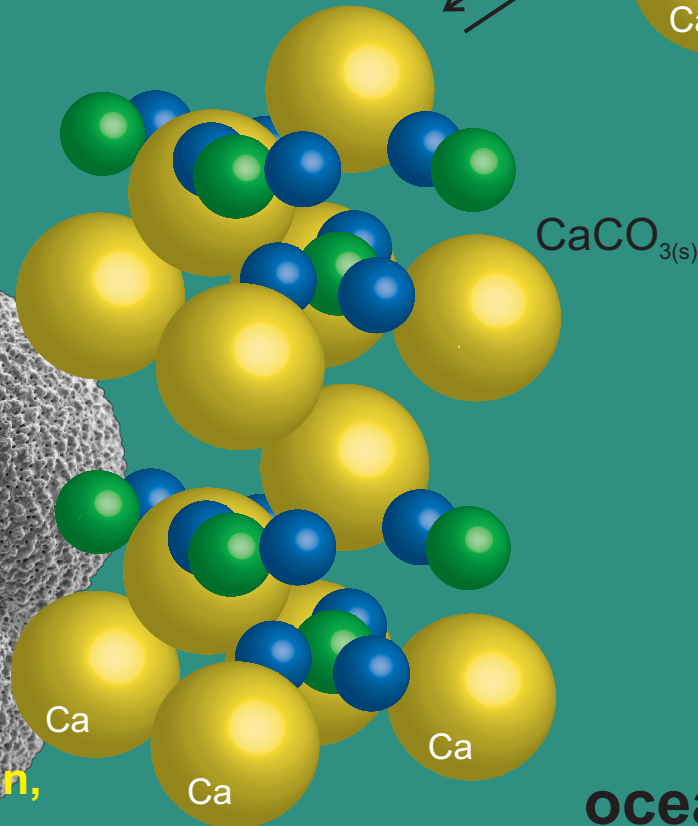
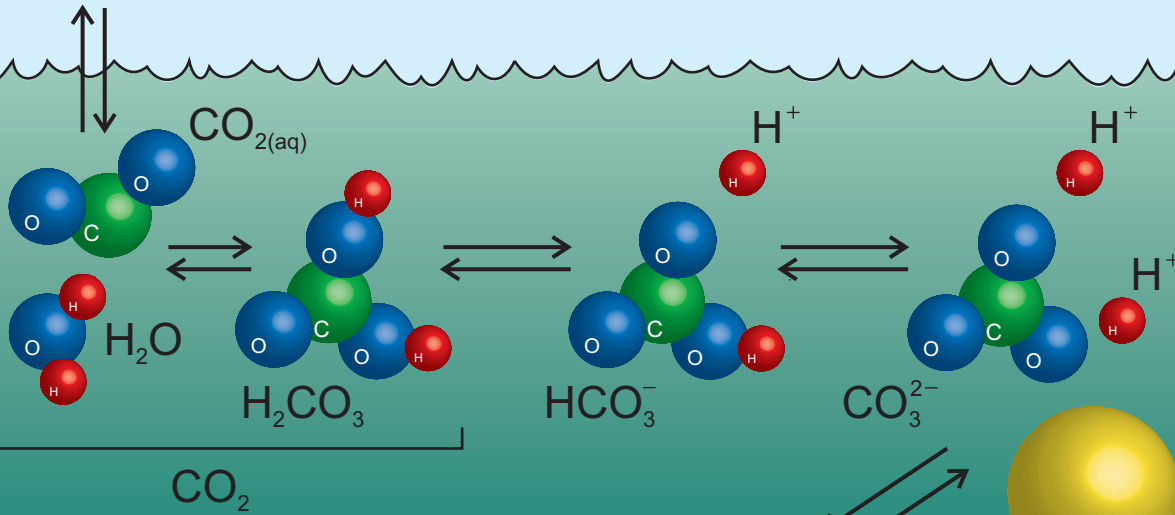
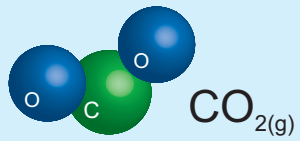
**Aragonite: less stable**  
orthorhombic polymorph (e.g., many corals, pteropods)



**Calcite: more stable**  
(and more abundant)  
trigonal polymorph (e.g., coccolithophorides, foraminifera)



# atmosphere



ocean

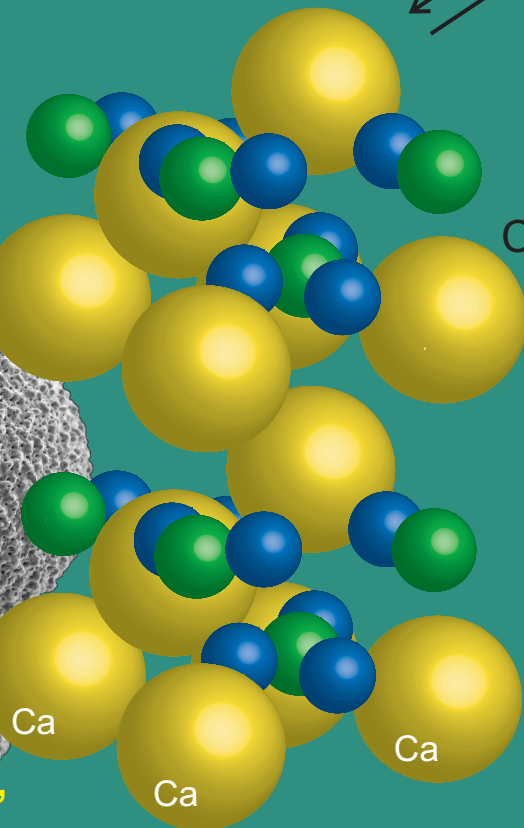
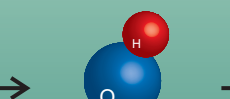
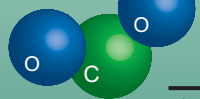
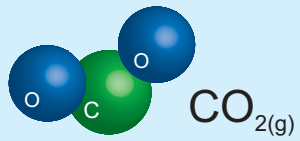
# $\text{CO}_2$ chemistry & mineral phases

The addition of (fossil fuel)  $\text{CO}_2$  to seawater results in a decrease in carbonate ion ( $\text{CO}_3^{2-}$ ) concentration and 'ocean acidification'. A decrease in  $\text{CO}_3^{2-}$ , in turn, suppresses the stability of  $\text{CaCO}_3$ , defined by its 'saturation state':

$$\Omega = [\text{Ca}^{2+}] \times [\text{CO}_3^{2-}] / k$$

$\Rightarrow$  The thermodynamic efficiency of precipitating  $\text{CaCO}_3$  is a function of  $[\text{CO}_3^{2-}]$  (and carbonate 'saturation').

atmosphere



Ca

Ca

Ca

ocean

calcium  
carbonate  
mineral  
surface

(calcifying plankton,  
e.g. foraminifera)

# $\text{CO}_2$ chemistry & mineral phases

The bottom-line:

more (fossil fuel)  $\text{CO}_2$



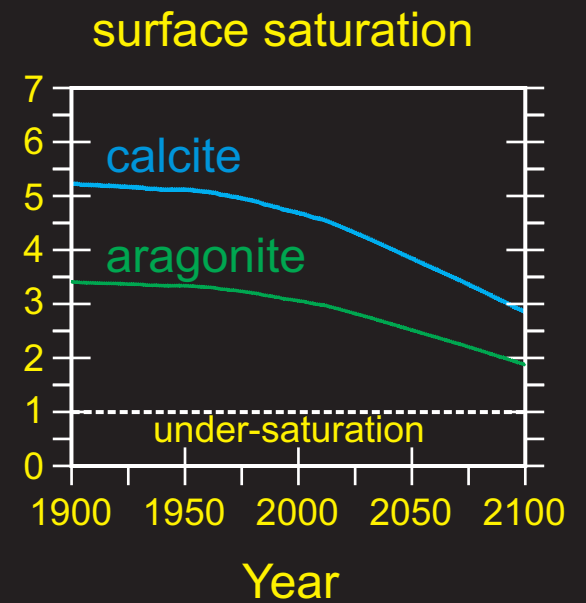
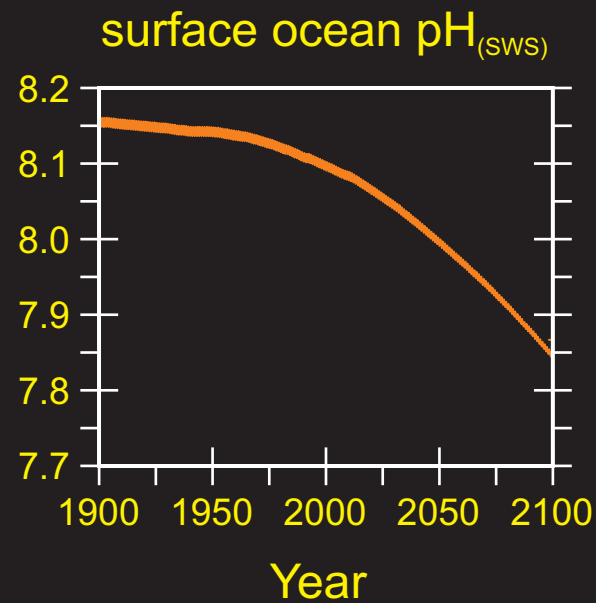
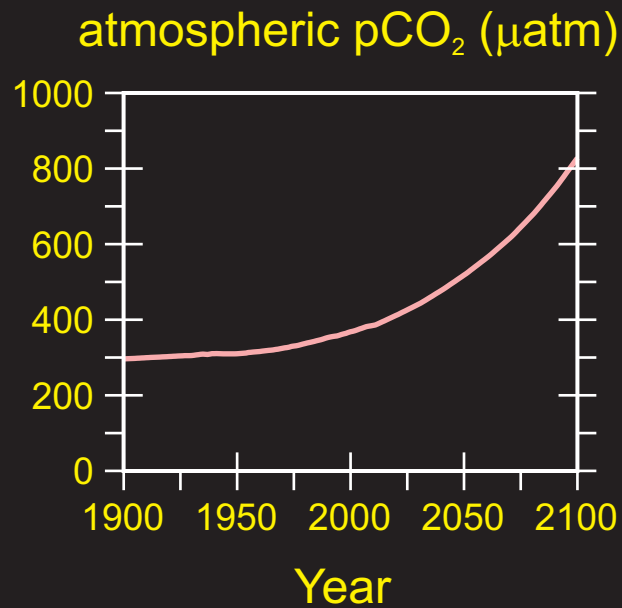
less  $\text{CO}_3^{2-}$  (& lower pH)

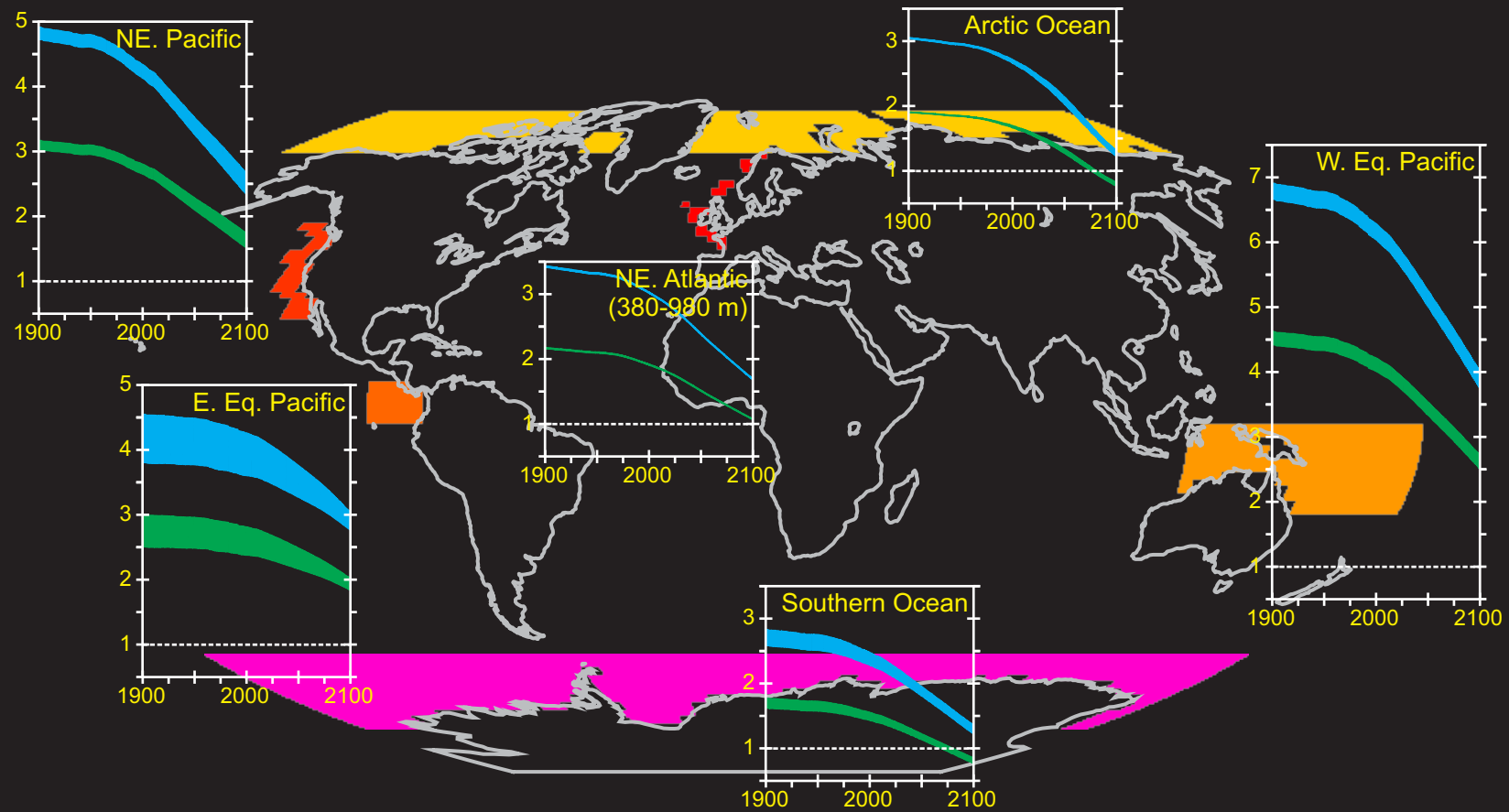


lower saturation ( $\Omega$ )  
& less stable  $\text{CaCO}_3$

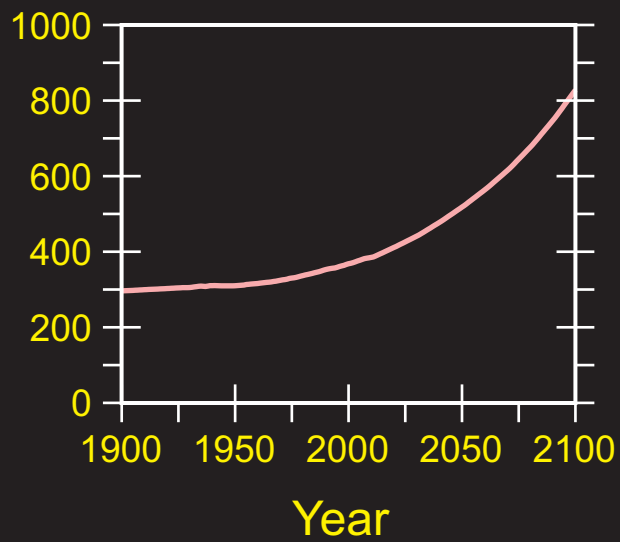
(i.e., calcite and aragonite will  
dissolve more readily or be less  
easily precipitated by  
organisms)

# CHANGING OCEAN CHEMISTRY

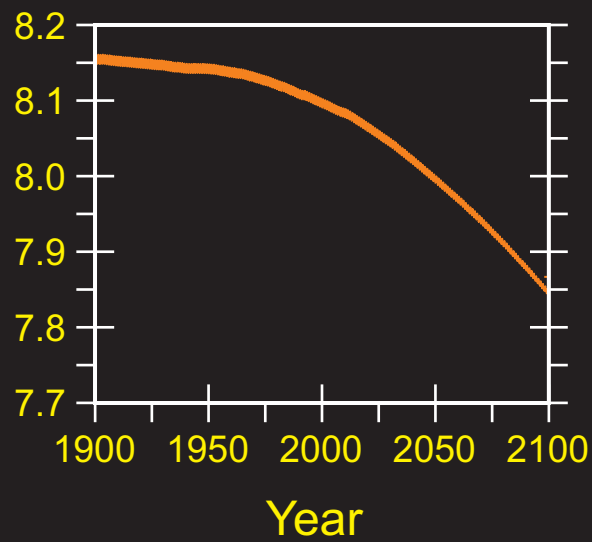




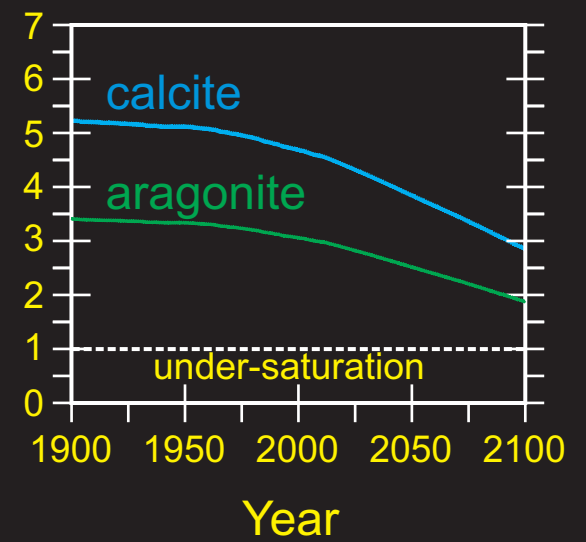
atmospheric pCO<sub>2</sub> (μatm)



surface ocean pH<sub>(SWS)</sub>

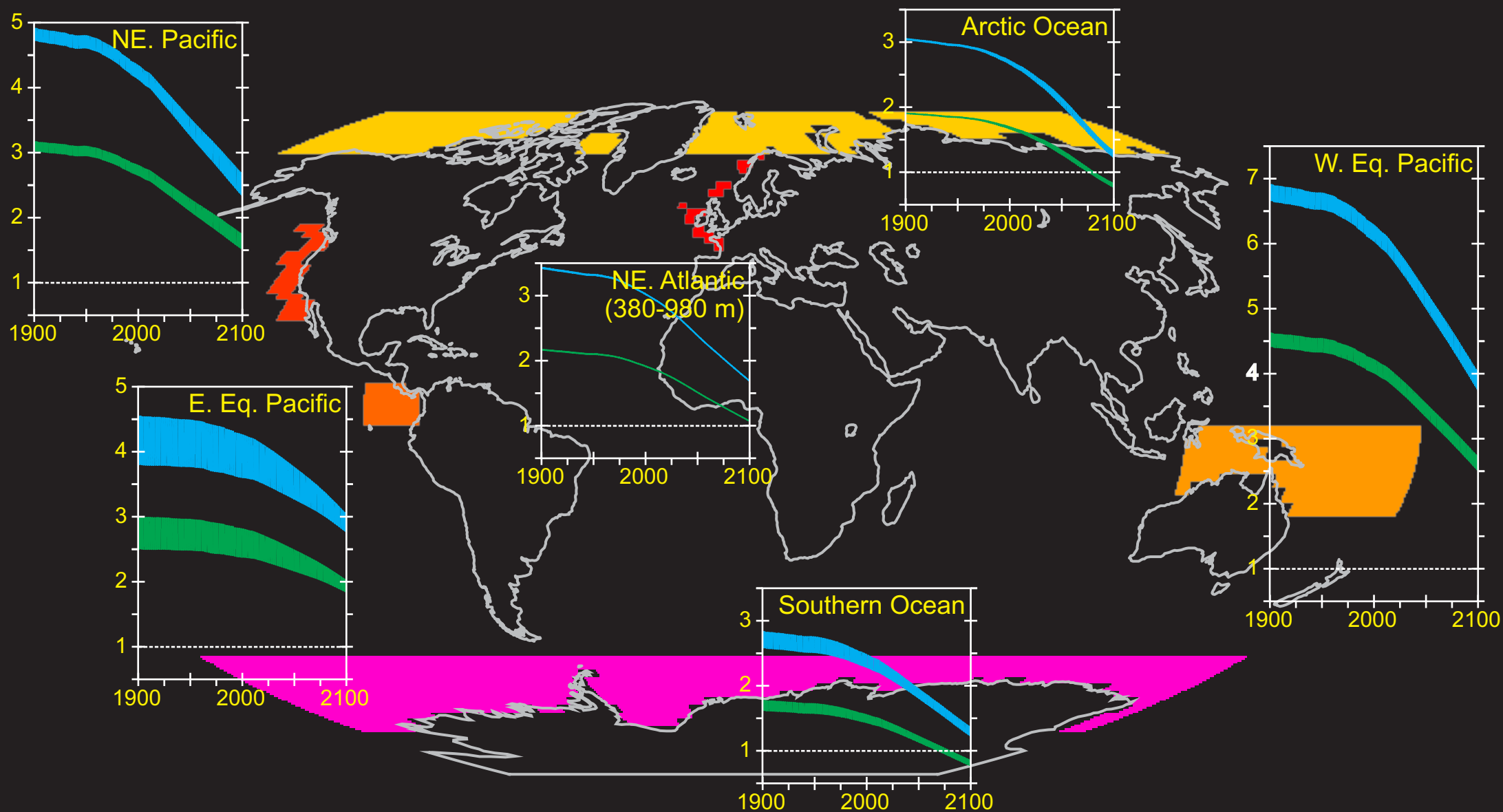


surface saturation

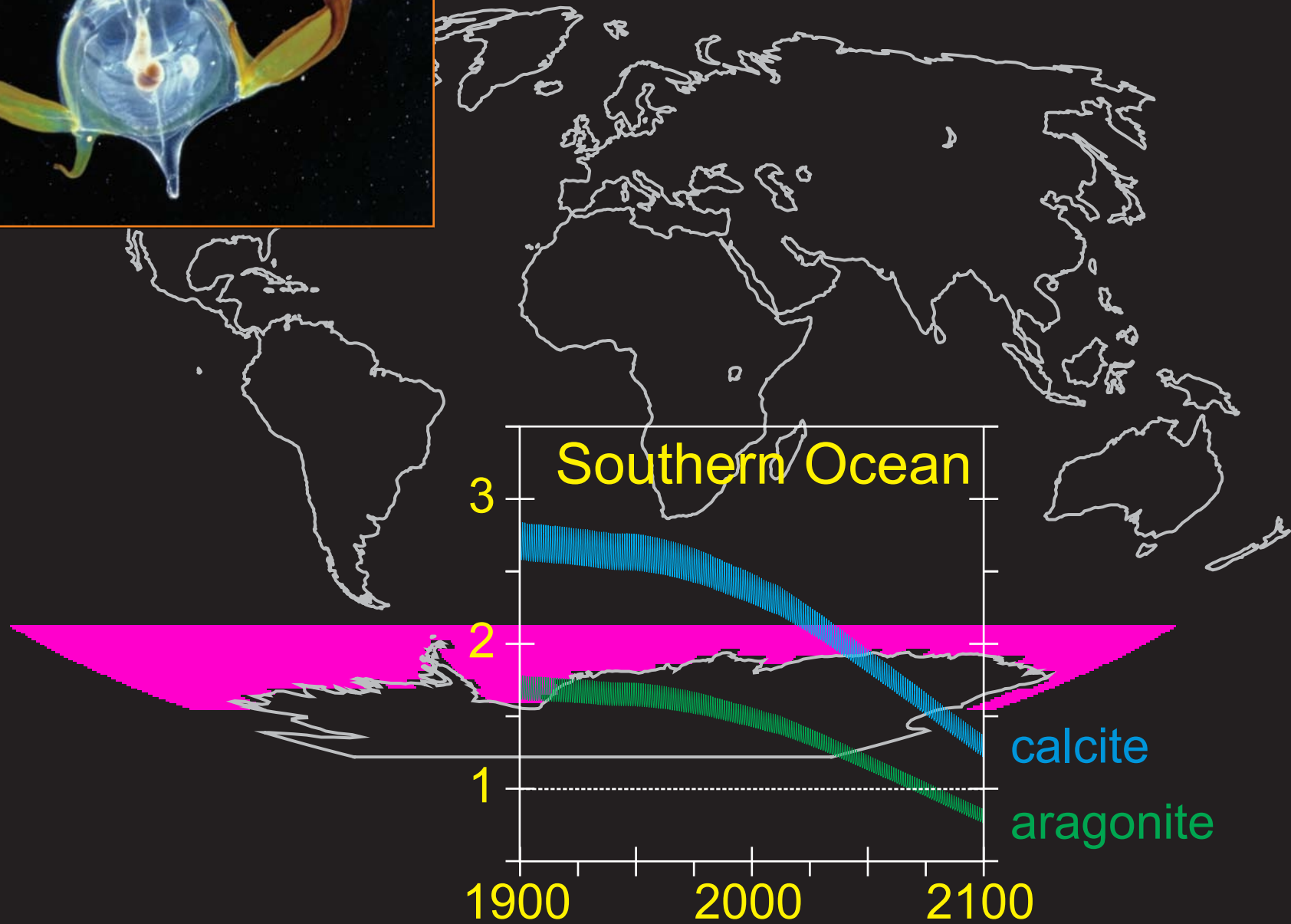




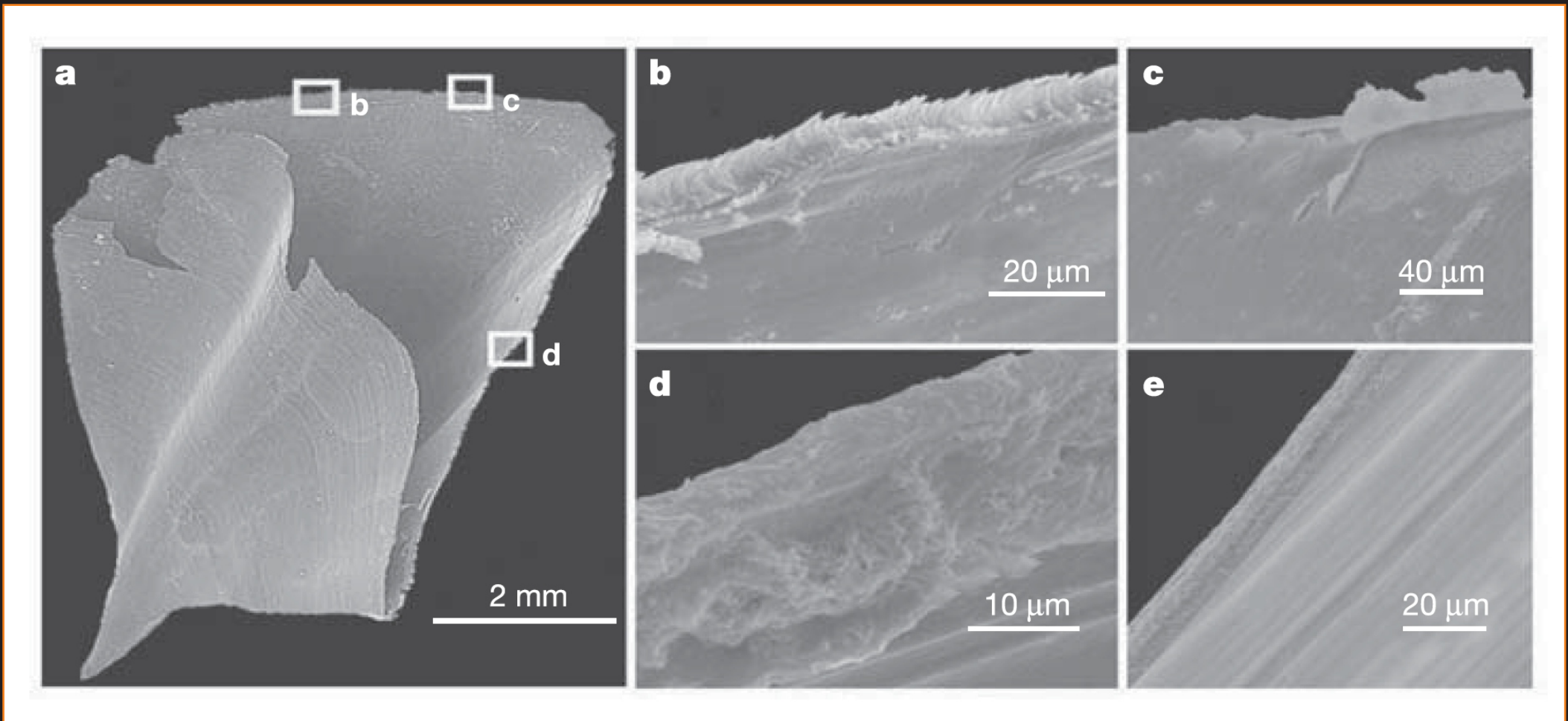
# SEASONAL/REGIONAL IMPACTS OF OCEAN ACIDIFICATION



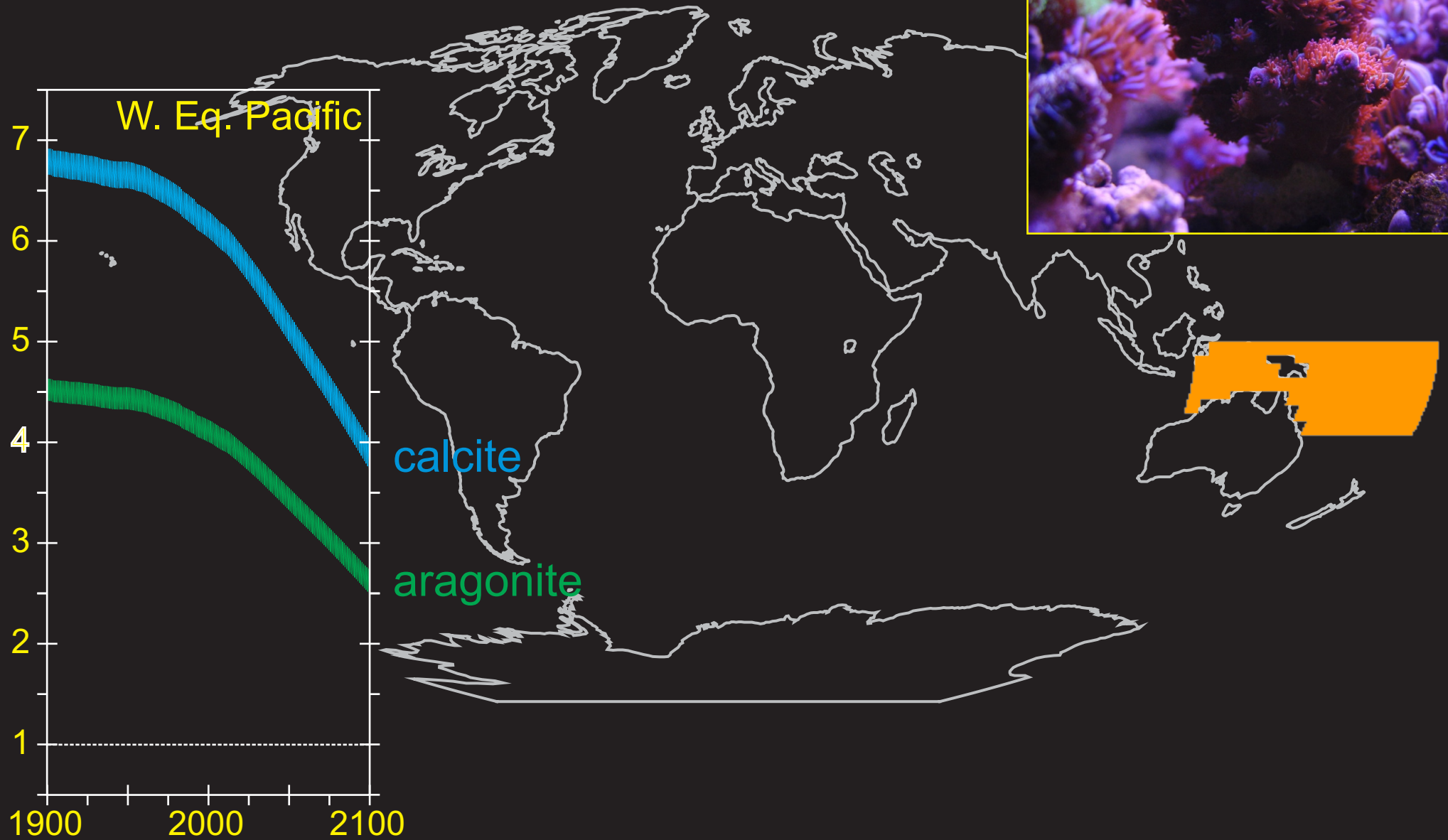
# Ocean biological consequences(?)



# Ocean biological consequences(?)

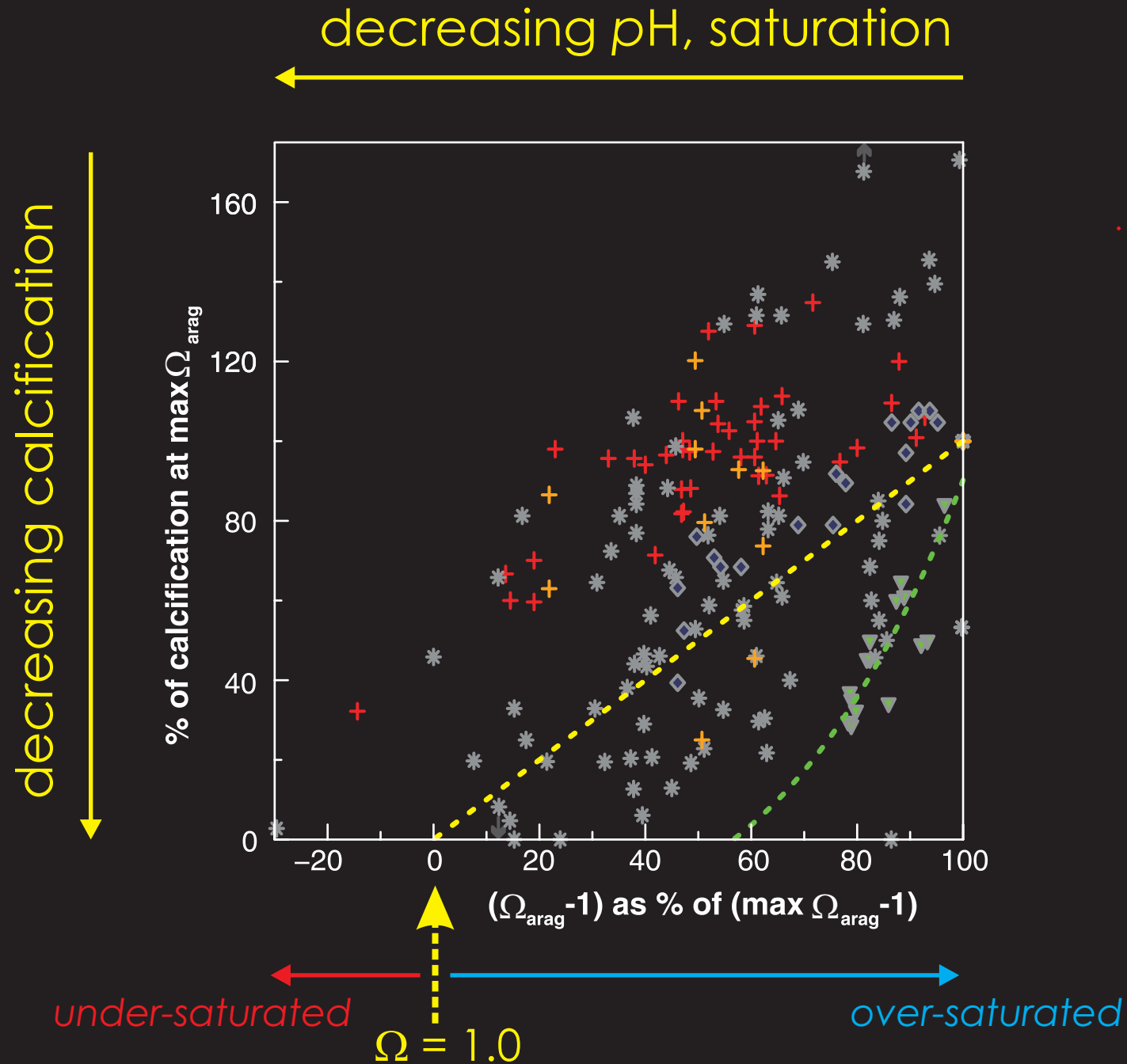


# Ocean biological consequences(?)

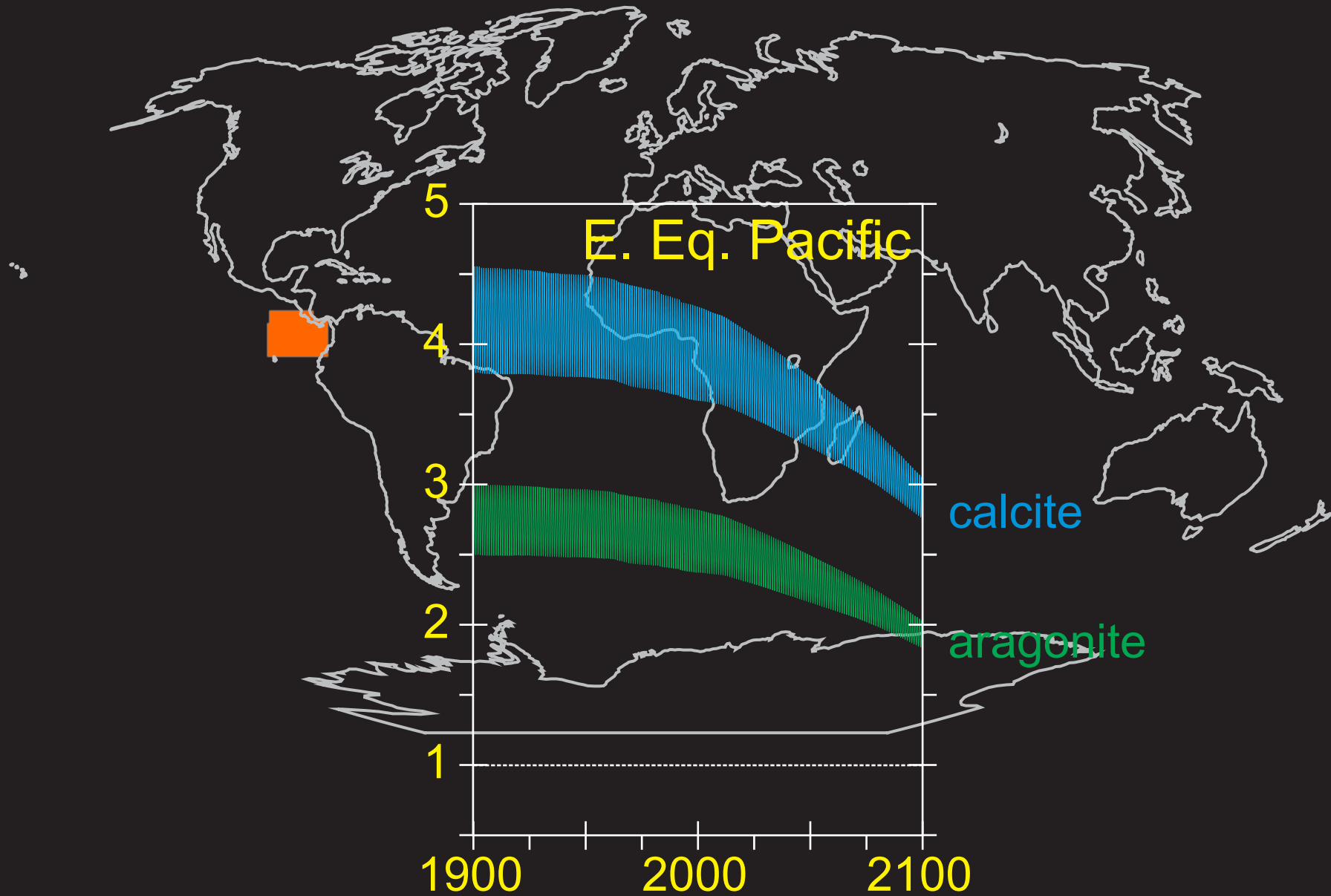




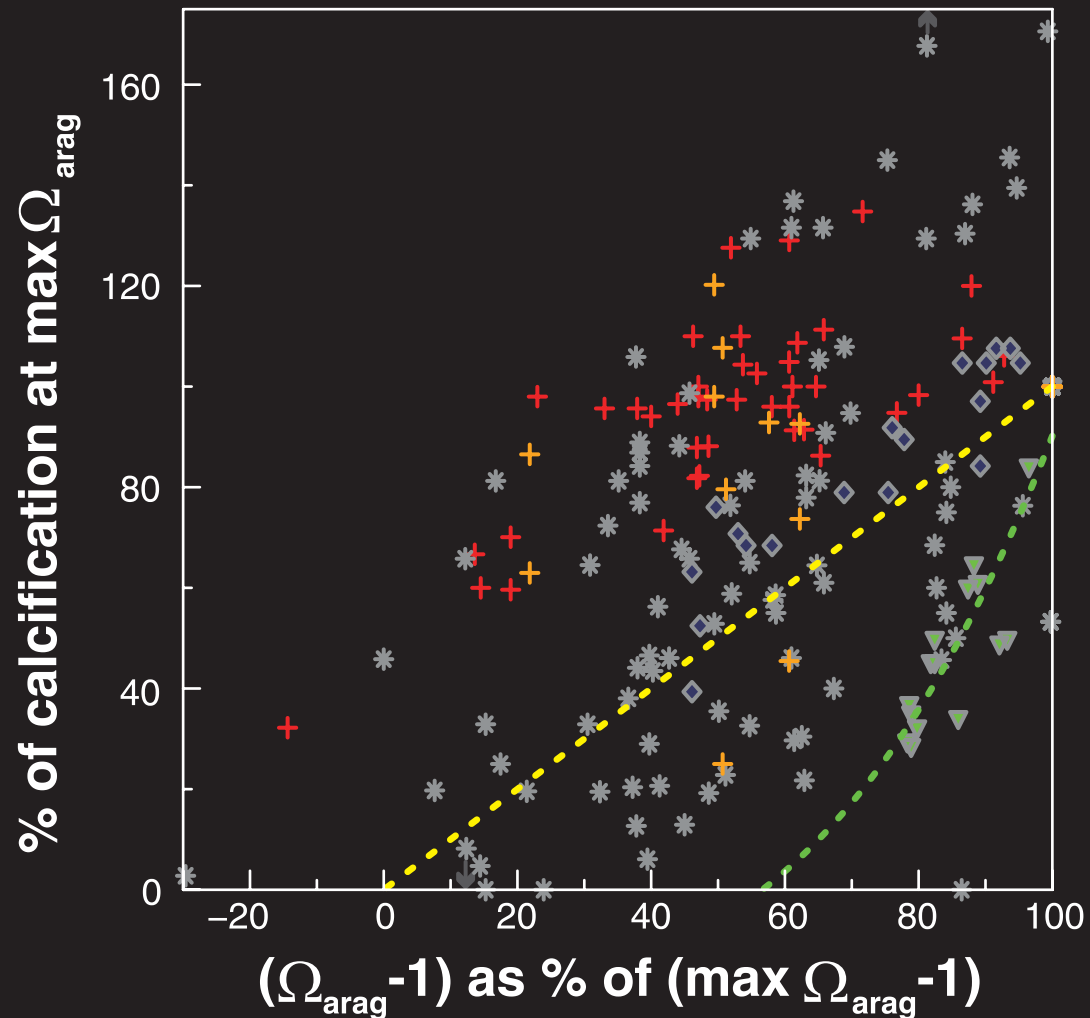
# Ocean biological consequences(?)



# Ocean biological consequences(?)



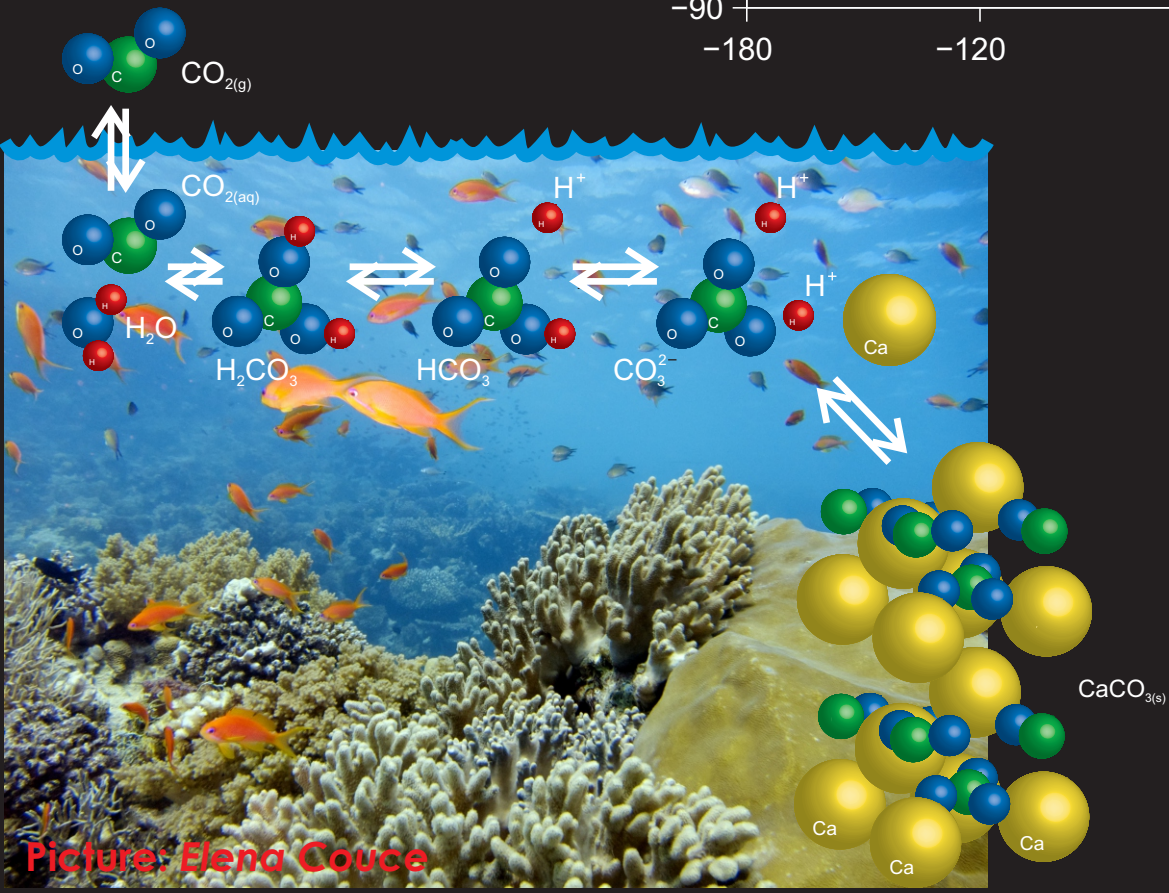
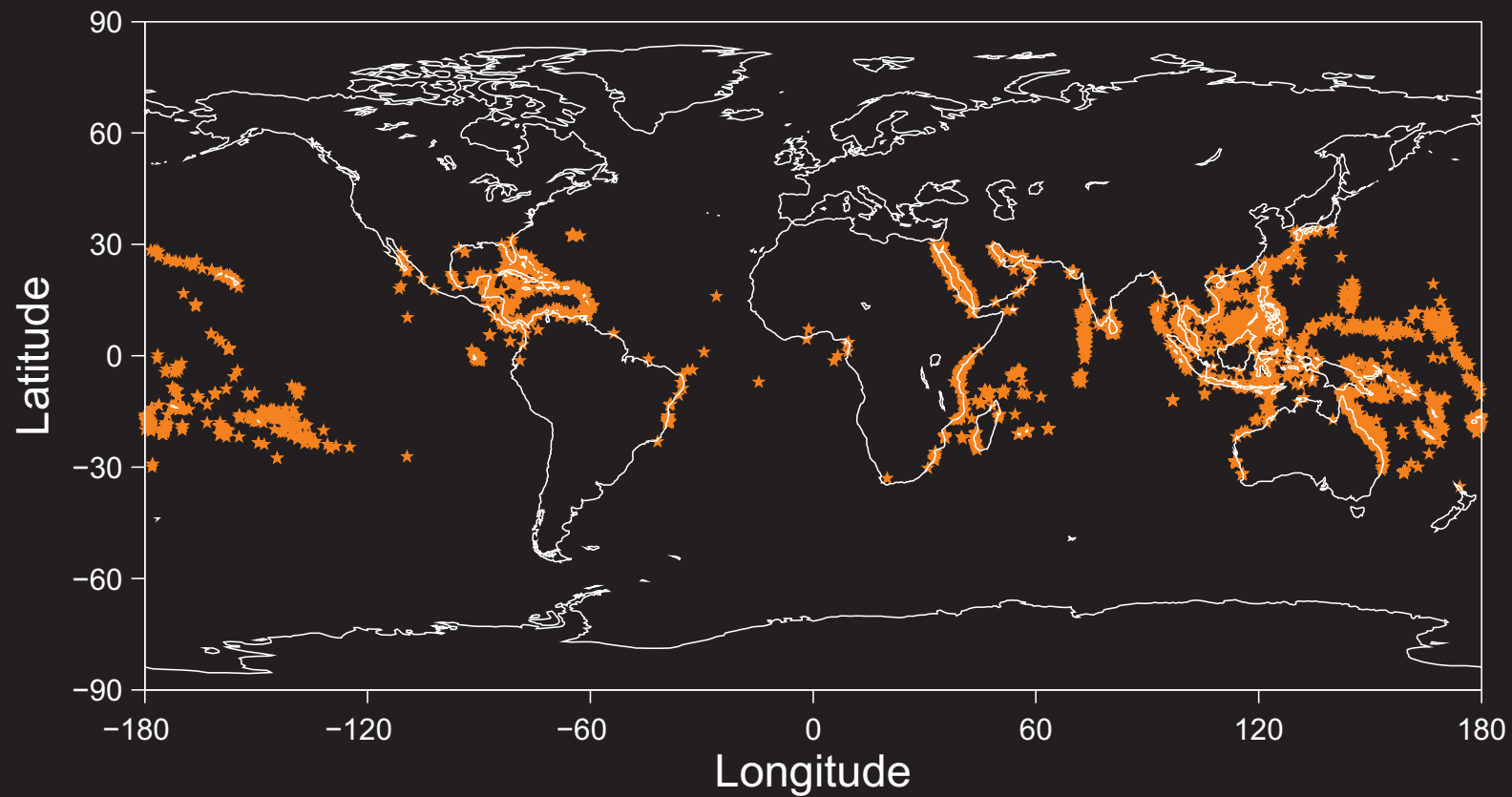
# Ocean biological consequences(?)



“A 10-20% decrease in  $\text{CaCO}_3$  production will pose a significant deficit for many coral reefs ... These might include high latitude reefs (for example, Bermuda), reefs in up-welling regions (for example, Galapagos), and many reefs experiencing anthropogenic stresses.”

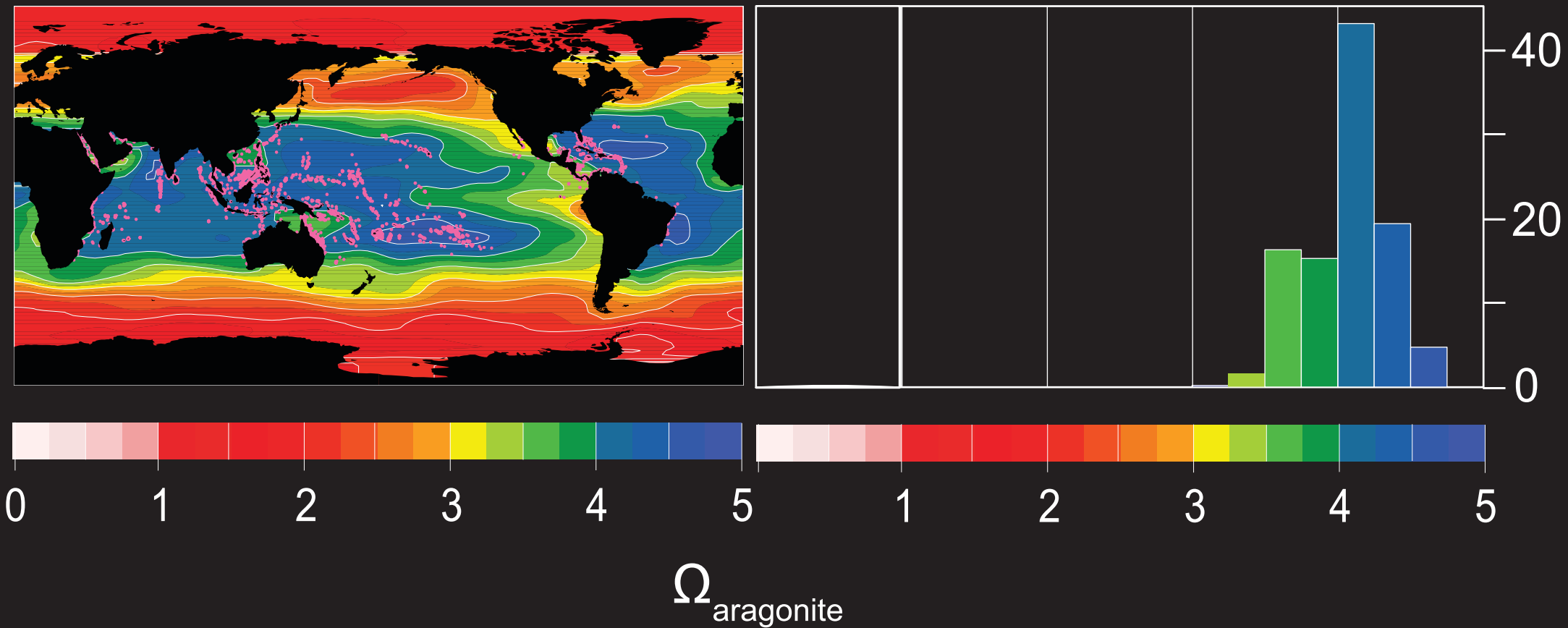
Kleypas et al. [1999] (Science 284)

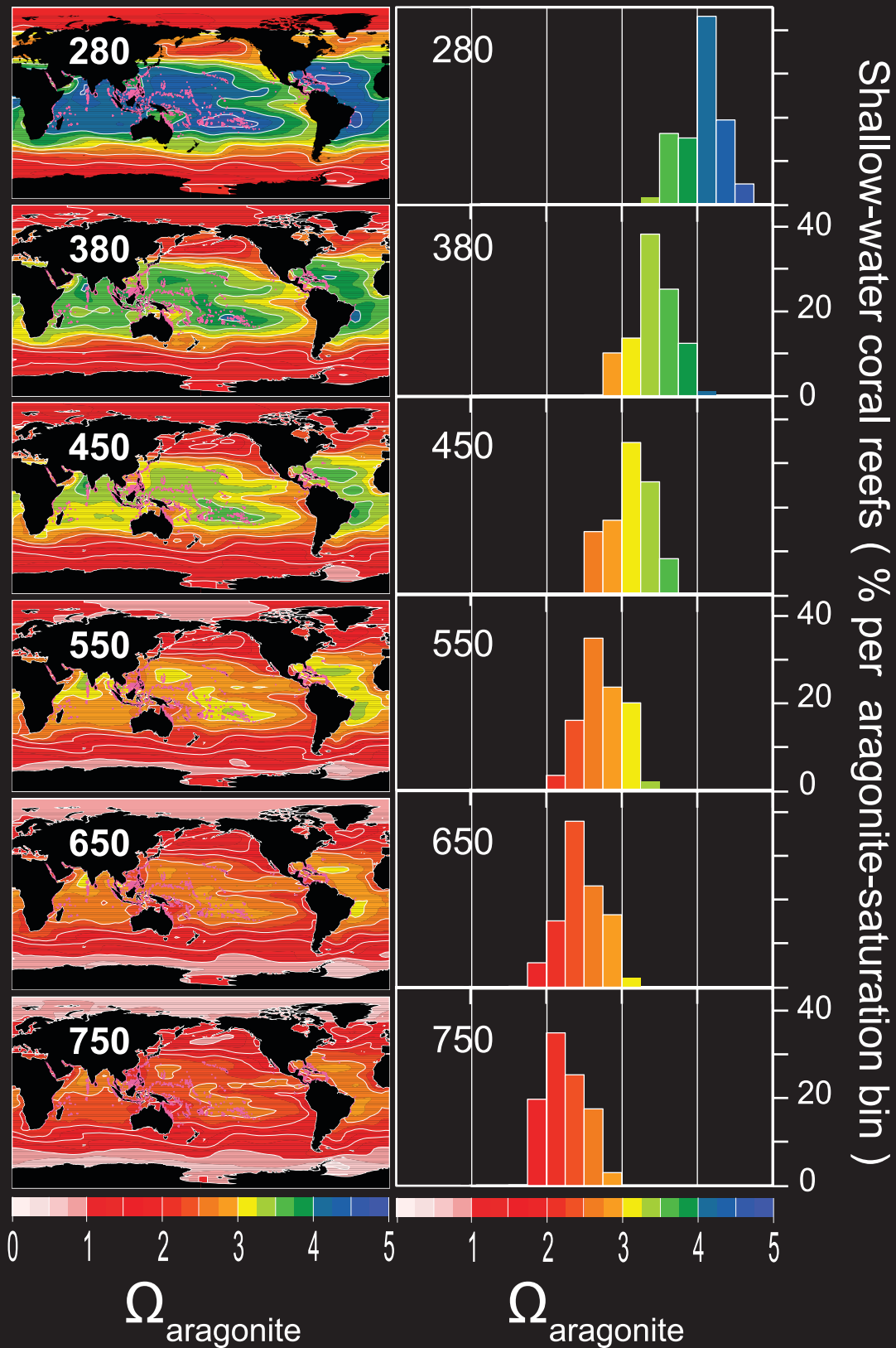




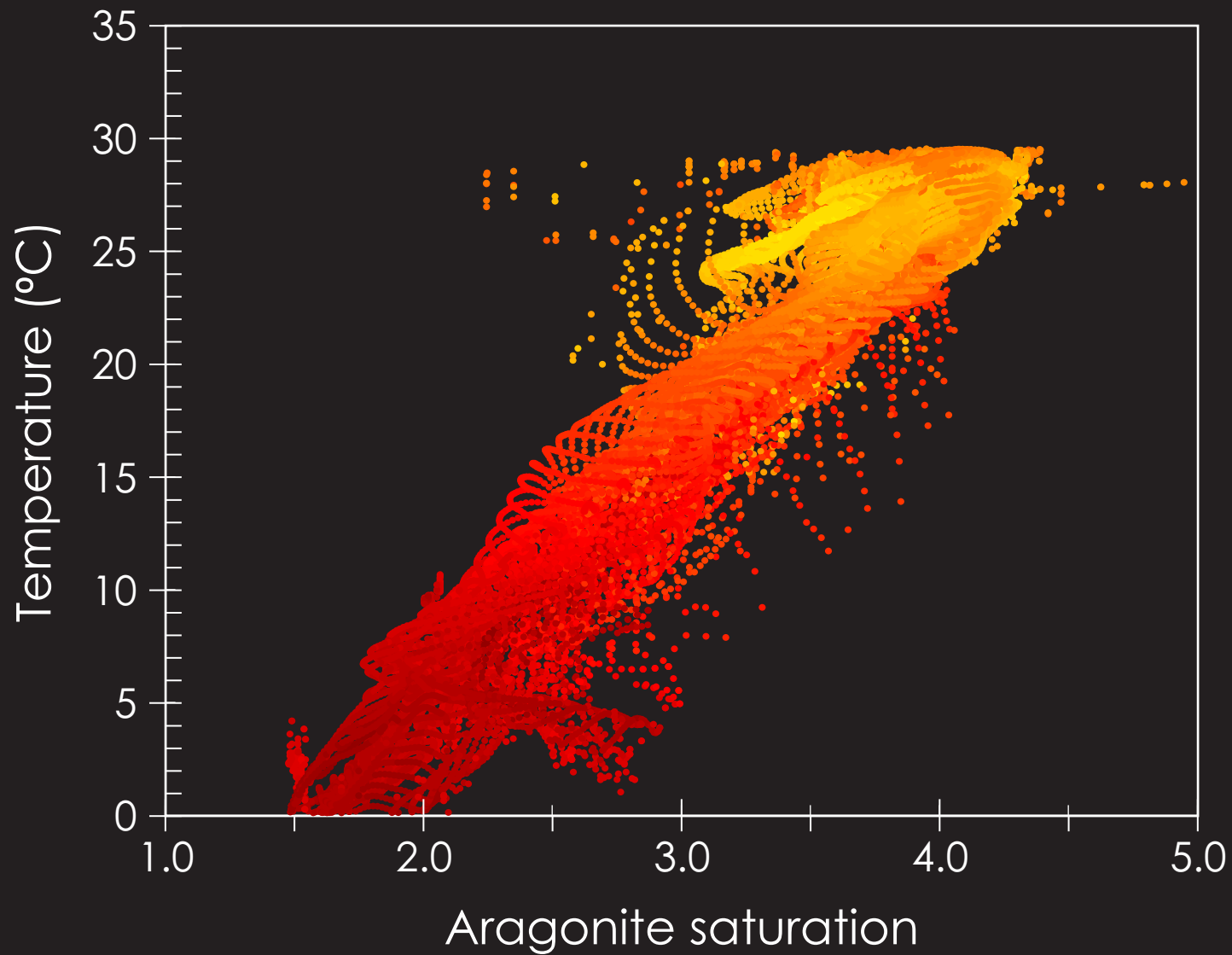
Picture: Elena Couce

# Ocean biological consequences(?)

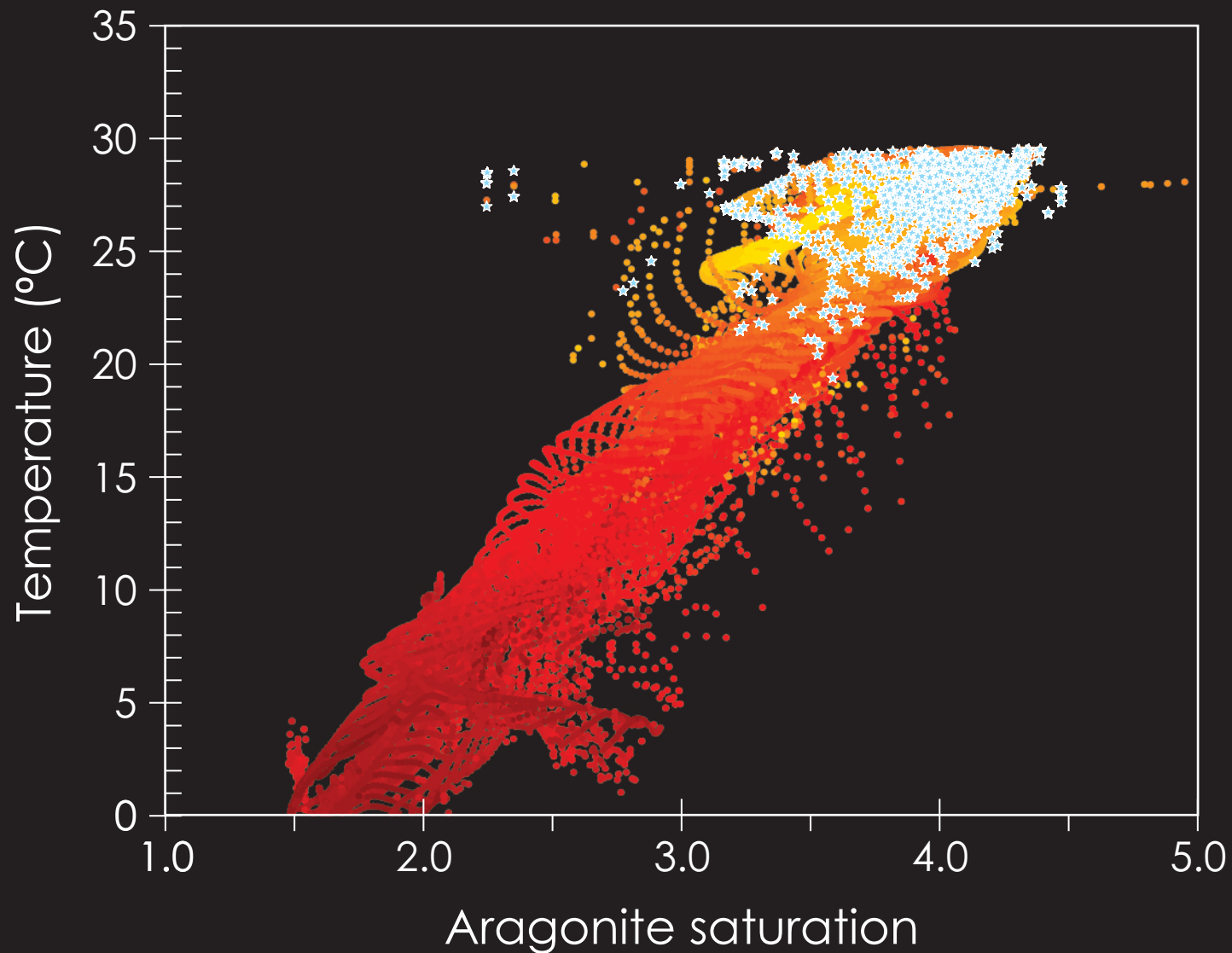




# *Modern ocean surface environmental relationships*

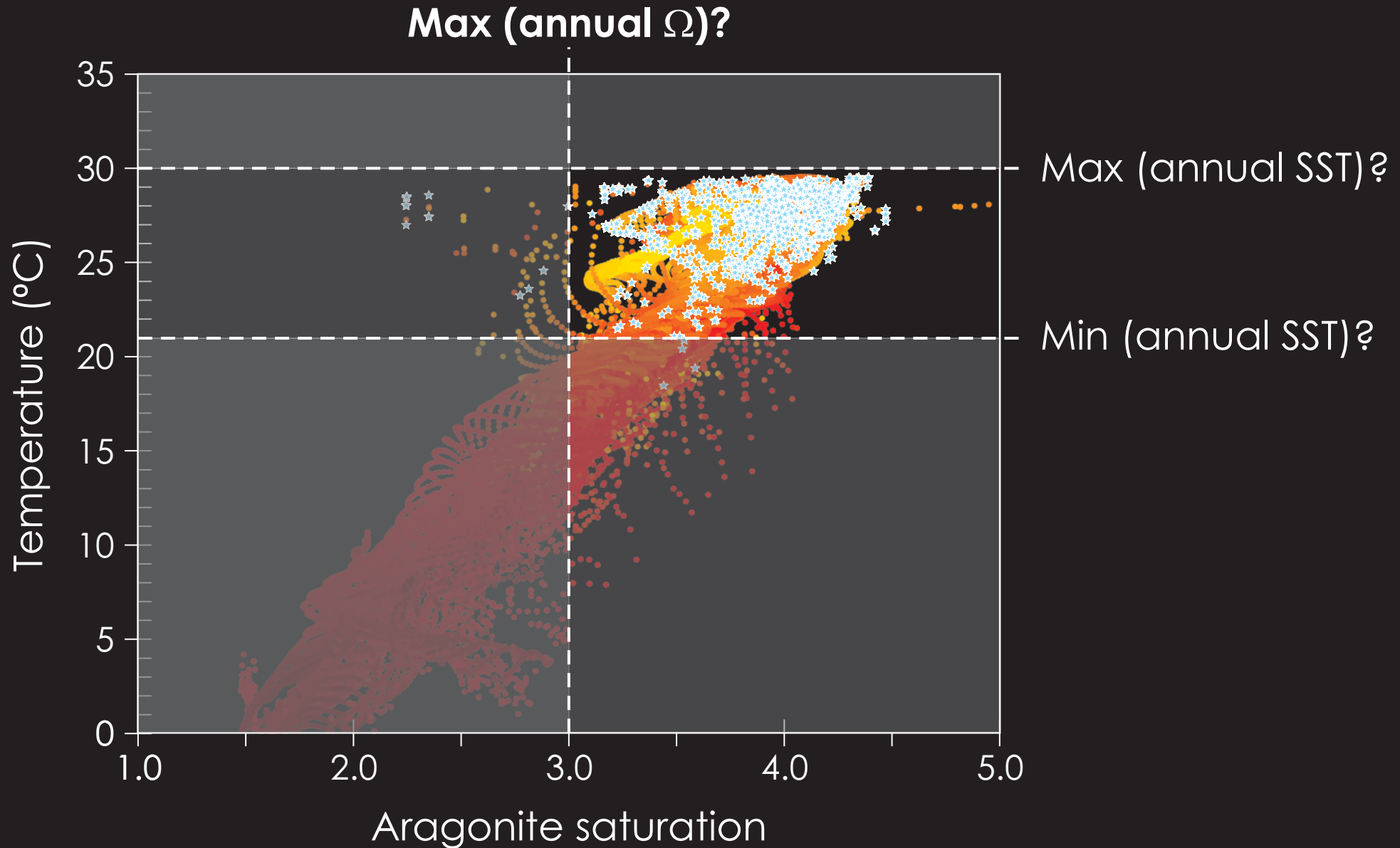


*Modern ocean surface environmental relationships*  
Tropical coral reef locations (ReefBase)



# Modern ocean surface environmental relationships

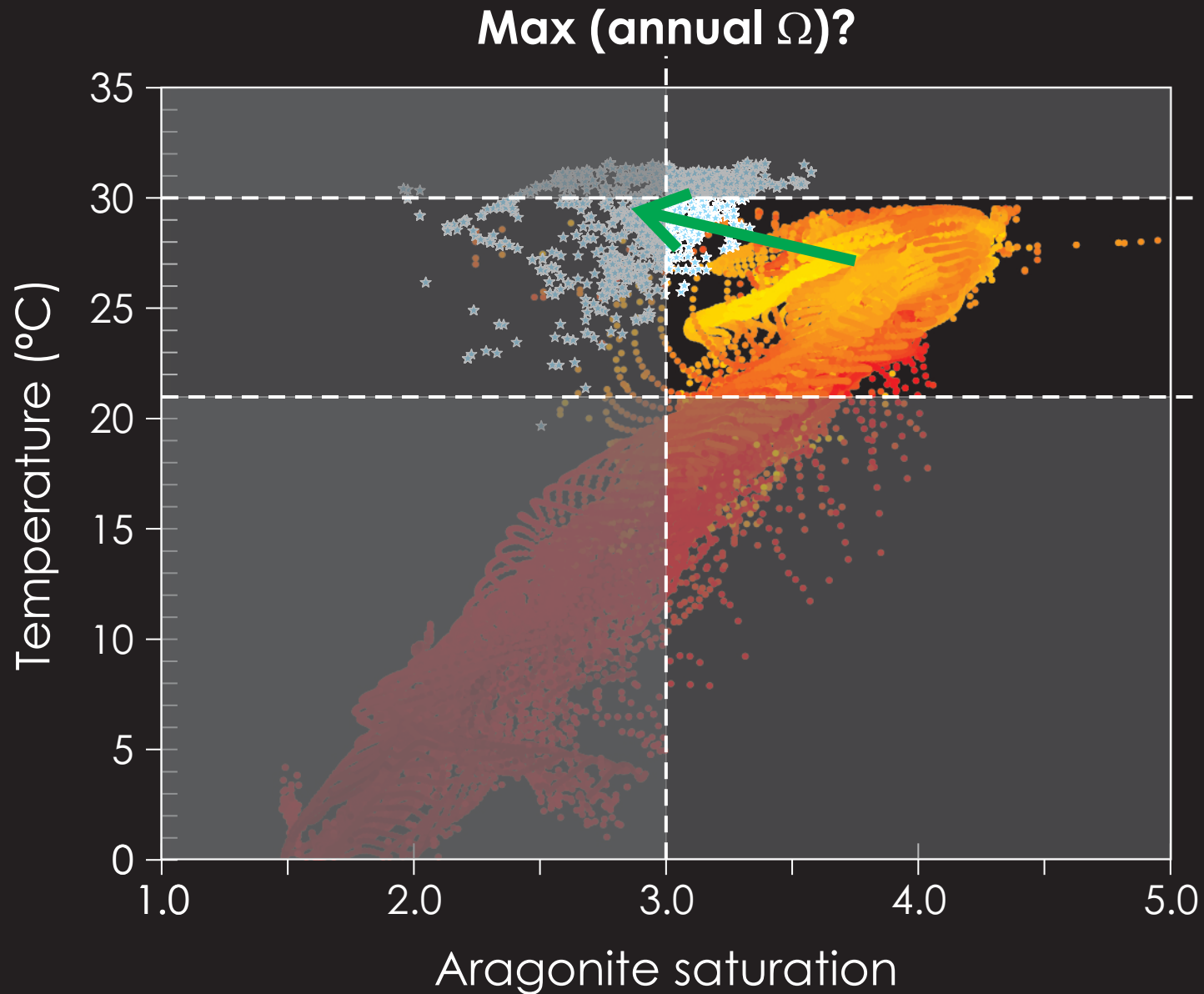
Tropical coral reef locations (ReefBase)





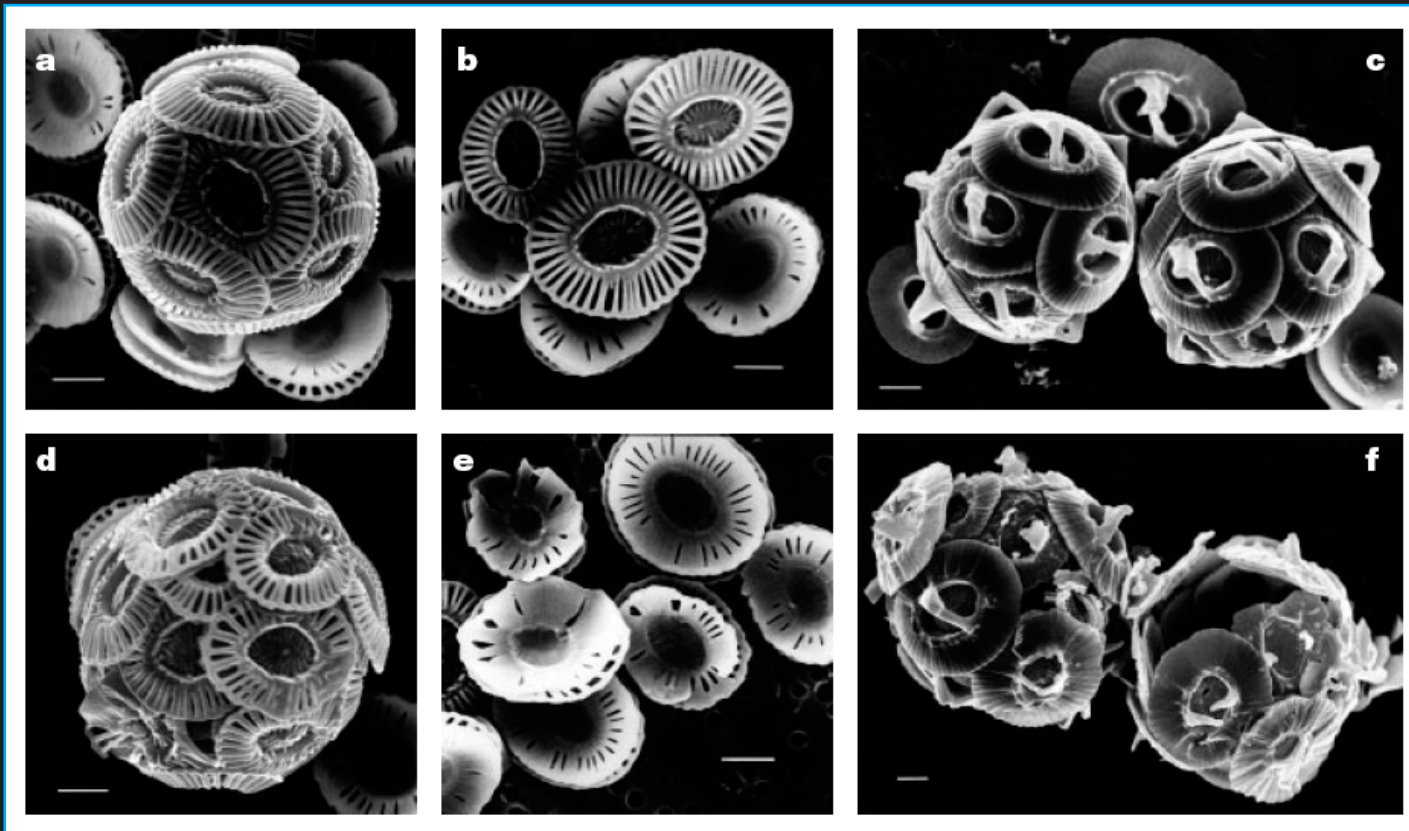
# Modern ocean surface environmental relationships

Reef environmental conditions @2070 under RCP 8.5





# Ocean biological consequences(?)



low CO<sub>2</sub> (high pH)

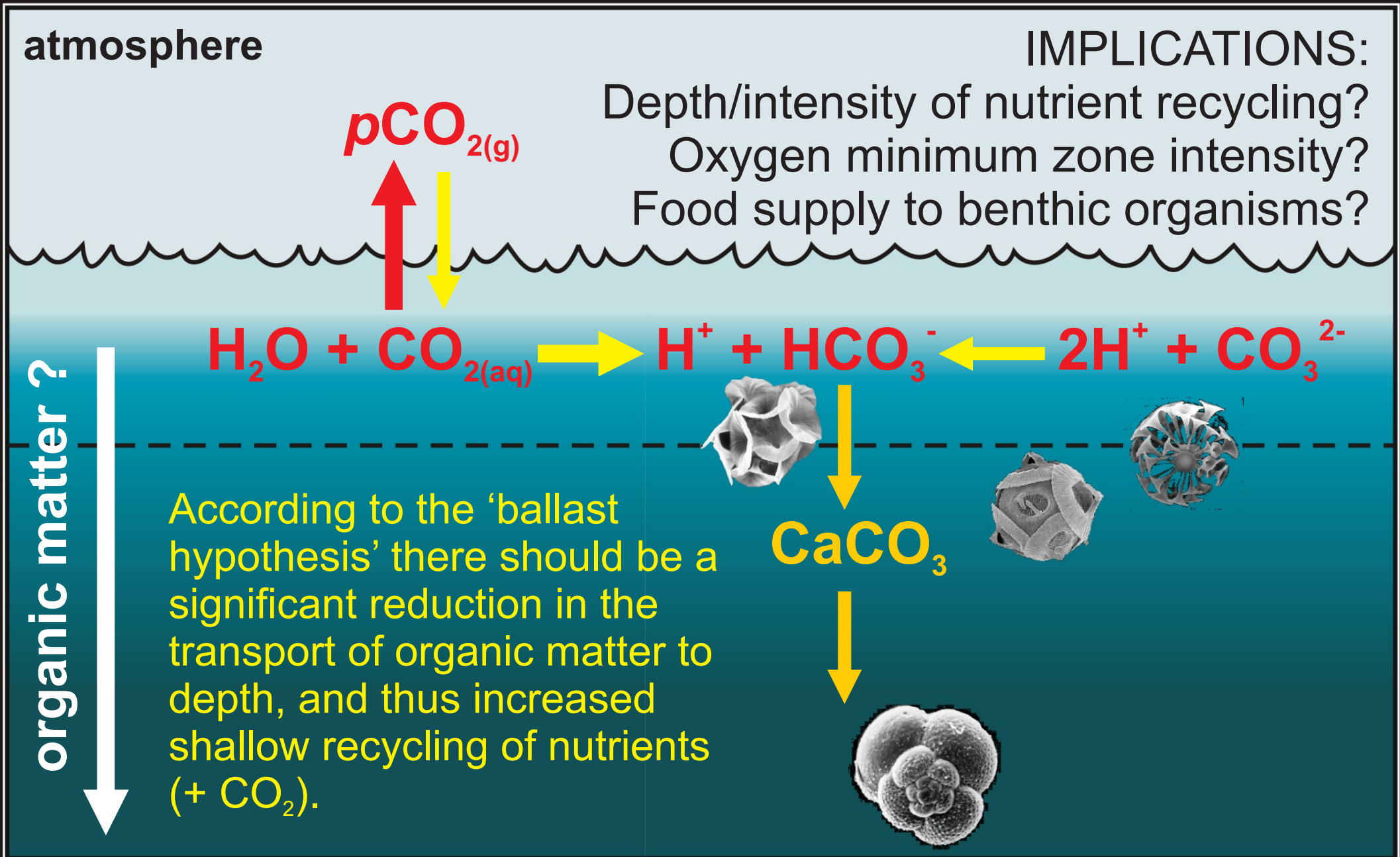
high CO<sub>2</sub> (low pH)

SEM micrographs of coccolithophorids under different CO<sub>2</sub> conditions  
*Riebesell et al. [2000] (Nature 407)*

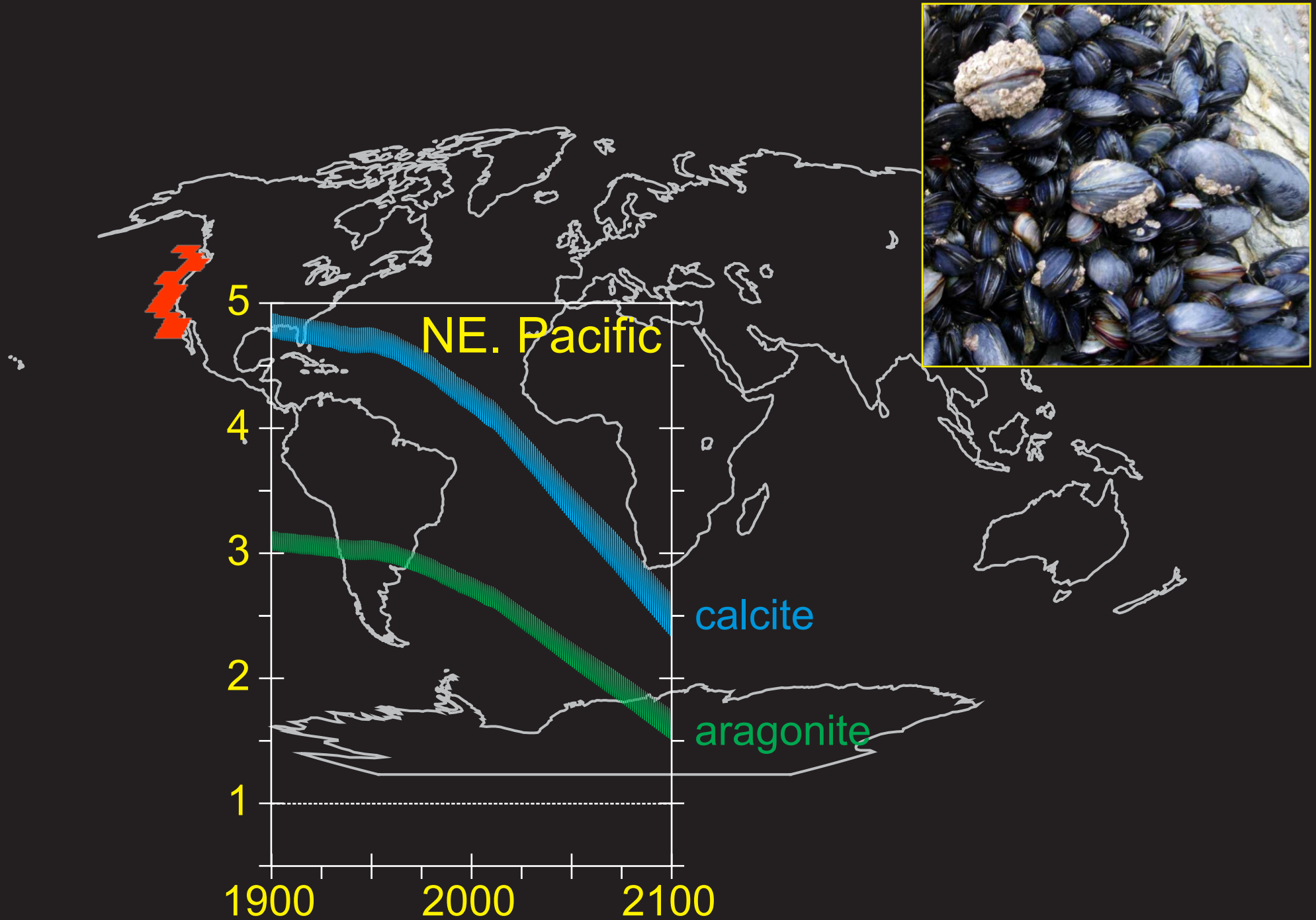
# Calcification responses ( $\text{CaCO}_3$ per cell per day) at elevated ( $\sim \times 2$ to $\times 3$ ) $\text{CO}_2$

Species	Strain	Year location	Exp. design	Manipulation		Reference
<i>Emiliana huxleyi</i>	PML B92/11A	1992 North Sea	laboratory culture	acid/base	↓	Riebesell et al. [2000] Zondervan et al. [2001]
<i>Emiliana huxleyi</i>	PML B92/11A	1992 North Sea	laboratory culture	acid/base	↓	Riebesell et al. [2000] Zondervan et al. [2001]
<i>Emiliana huxleyi</i>	CAWPO6	1992 South Pacific	laboratory culture	$\text{CO}_2$ bubbling	↑	Iglesias-Rodriguez et al. [2008]
<i>Emiliana huxleyi</i>	MBA 61/12/4	1991 N. Atlantic	laboratory culture	$\text{CO}_2$ bubbling	↑	Iglesias-Rodriguez et al. [2008] (pers com)
<i>Emiliana huxleyi</i>	CCMP 371	1987 Sargasso Sea	laboratory culture	$\text{CO}_2$ bubbling	↓	Feng et al. [2008]
<i>Emiliana huxleyi</i>	CCMP 371	1987 Sargasso Sea	laboratory culture	$\text{CO}_2$ bubbling	↓	Feng et al. [2008]
<i>Emiliana huxleyi</i>	TW1	2001 W. Mediterranean	laboratory culture	$\text{CO}_2$ bubbling	↓	Sciandra et al. [2003]
<i>Emiliana huxleyi</i>	Ch 24-90	1991 North Sea	laboratory culture	$\text{CO}_2$ bubbling	↔	Buitenhuis et al. [1999]
<i>Emiliana huxleyi</i>	CAWPO6	1992 South Pacific	laboratory culture	$\text{CO}_2$ bubbling	↑	Shi et al. [2009]
<i>Gephyrocapsa oceanica</i>	PC7/1	1998 Portuguese shelf	laboratory culture	acid/base	↓	Riebesell et al. [2000] Zondervan et al. [2001]
<i>Calcidiscus leptoporus</i>	AC365	2000 S. Atlantic	laboratory culture	acid/base	↓↑	Langer et al. [2006]
<i>Coccolithus pelagicus</i>	AC400	2000 S. Atlantic	laboratory culture	acid/base	↔	Langer et al. [2006]

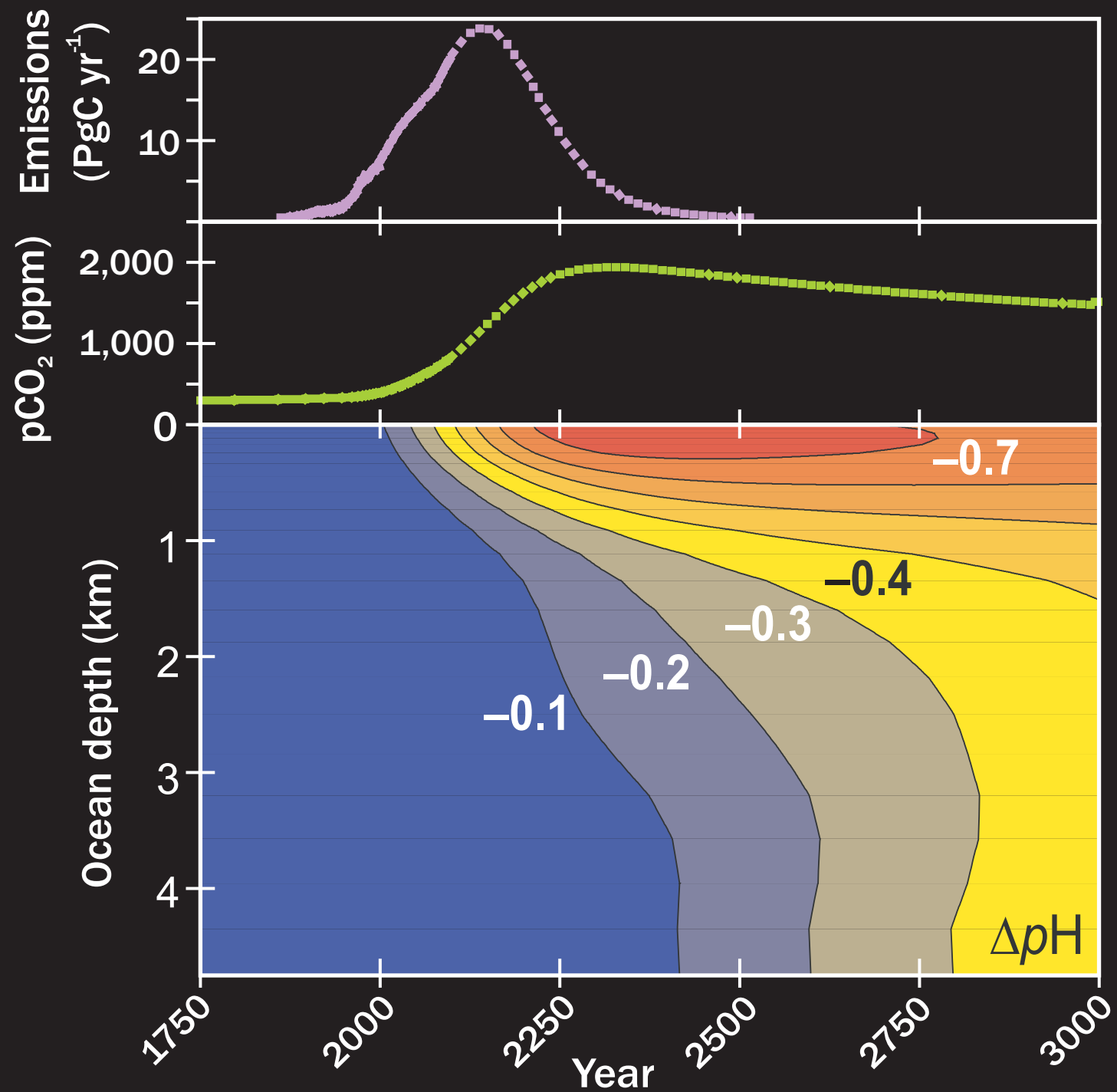
# Global biogeochemical impacts in a High CO<sub>2</sub> World



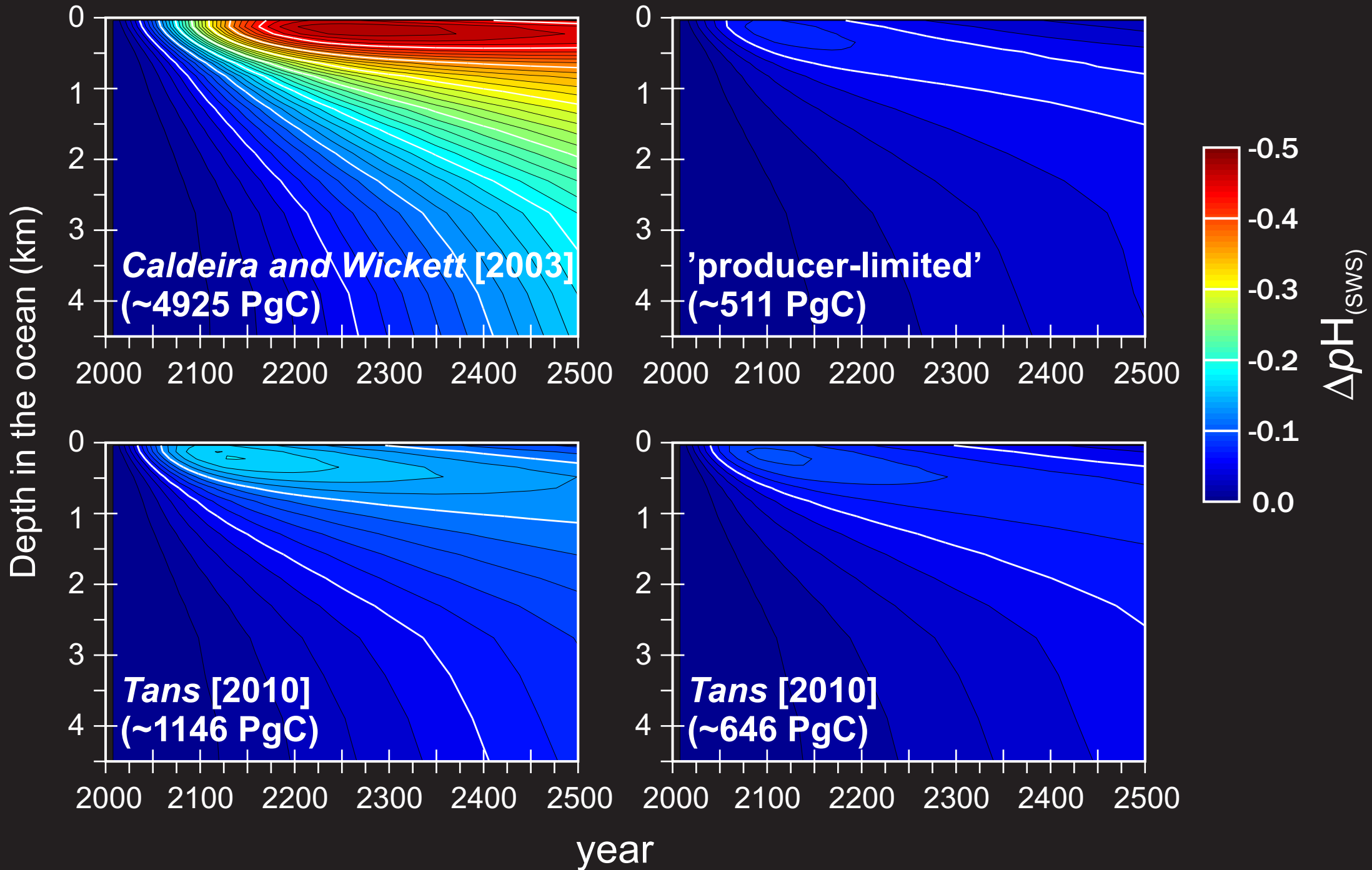
# Ocean biological consequences(?)



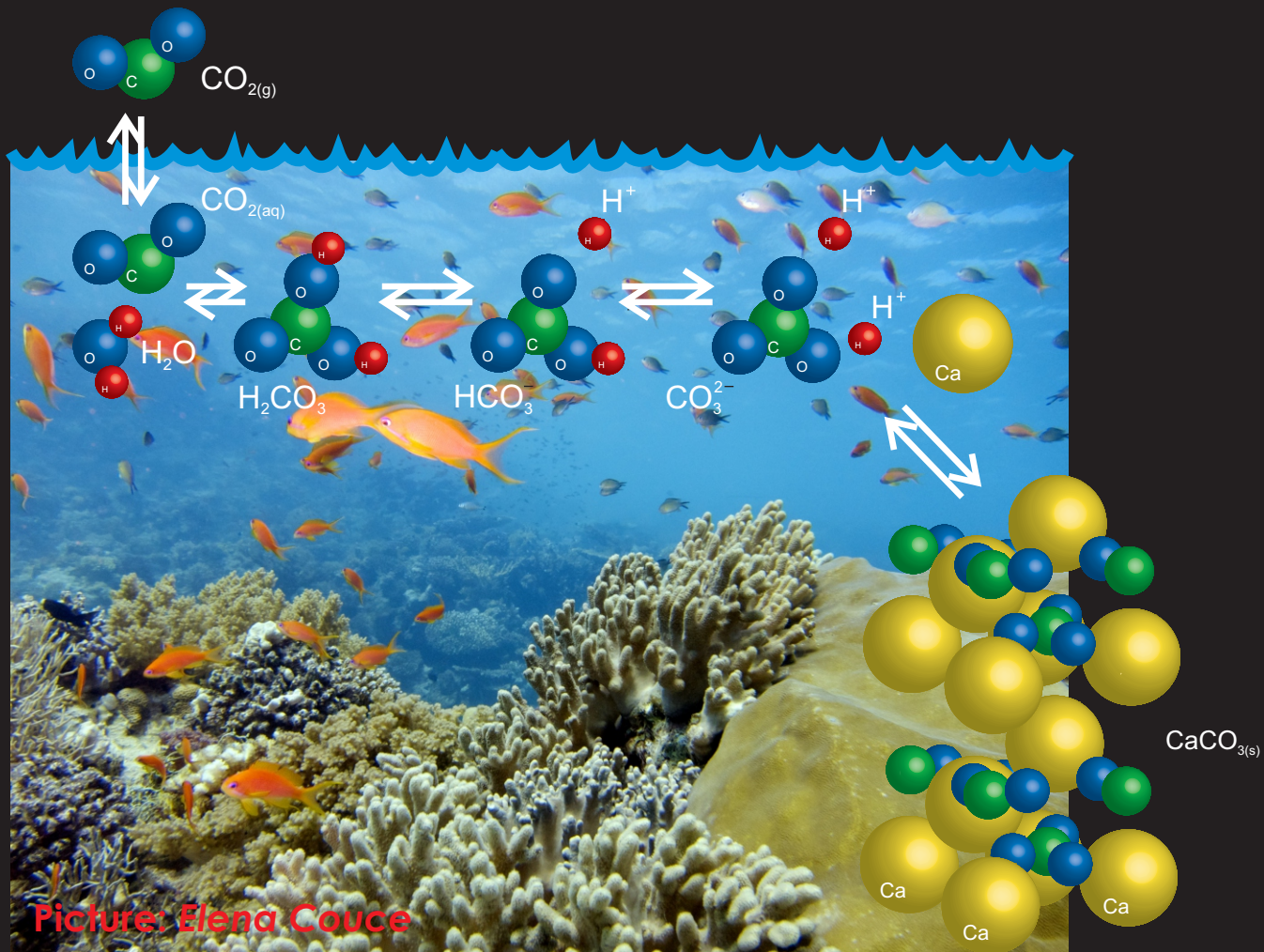
# Future trajectories of ocean pH



# Future ocean pH projections







Picture: Elena Couce