Snowball Earth



Fairchild and Kennedy [2007] (more recent and slightly more neutral/contrarian review)

Hoffman and Schrag [2002] ('pro' snowball Earth hypothesis review)

Ridgwell and Kennedy [2004] (a different and more carbon cycle focussed viewpoint) The Neoproterozoic: Gateway to a metazoandominated, oxygenated, 'modern-like' biosphere?

	Neopro	terozoic	Mesoproterozoic				P	Paleoproterozoic							Achean				
			l I		l	<u> </u>							ļ	I					
00		1(000			1500				2000				250	0				300
						-	Time	(Myr	B.P.)										



Background



Background – evidence for glaciation

Snowball Earth





From: Hoffman and Schrag [2002]

Background – evidence for glaciation





From: Fairchild and Kennedy [2007]

From: Hoffman and Schrag [2002]

Background – evidence for glaciation



From: Hoffman and Schrag [2002]



Sanchez-Baracaldo et al. [2014]



Nodes of planktonic cyanobacteria first occurrence (dashed circle) and divergence (full circle)

Sanchez-Baracaldo et al. [2014]



Nodes of planktonic cyanobacteria first occurrence (dashed circle) and divergence (full circle)

Sanchez-Baracaldo et al. [2014]

Snowball Earth

Hoffman et al. [1998] (Science 281)











Snowball Earth



positive "ice-albedo" feedback

Snowball Earth







Temperature

Snow cover

 $= -1/2^{\circ}C$

Snowball Earth



Temperature

Snow cover

TOTAL CHANGE = -1/2°C

Snowball Earth





Temperature



Snow cover

TOTAL CHANGE = $-1/2^{\circ}C - 1/4^{\circ}C$

Snowball Earth



Temperature

Snow cover

TOTAL CHANGE = $-1/2^{\circ}C - 1/4^{\circ}C$

Snowball Earth





Temperature



Snow cover

TOTAL CHANGE = $-1/2^{\circ}C - 1/4^{\circ}C - 1/8^{\circ}C$

Snowball Earth



TOTAL CHANGE = $-1/2^{\circ}C - 1/4^{\circ}C - 1/8^{\circ}C - 1/16^{\circ}$ -

Snowball Earth



TOTAL CHANGE = $-1^{\circ}C - 2^{\circ}C - 4^{\circ}C - 8^{\circ}C$

• • • • • • •



Global mean surface

temperature (°C)

Snowball Earth

adapted from; *Hoffman and Schrag* [2002]

Global carbon cycling (modern)



Global carbon cycling (long-term controls)

Terrestrial weathering can be (approximately equally) divided into carbonate (CaCO₃) and calcium-silicate ('CaSiO₃') weathering:

- (1) $2CO_{2(aq)} + H_2O + CaSiO_3 \rightarrow Ca^{2+} + 2HCO_3^{-} + SiO_2$
- (2) $CO_{2(aq)} + H_2O + CaCO_3 \rightarrow Ca^{2+} + 2HCO_3^{--}$

Ultimately, the (alkalinity: Ca²⁺) weathering products must be removed through carbonate precipitation and burial in marine sediments:

(3) $Ca^{2+} + 2HCO_3^{-} \rightarrow CO_{2(aq)} + H_2O + CaCO_3$

It can be seen that in (2) + (3), that the CO₂ removed (from the atmosphere) during weathering, is returned upon carbonate precipitation (and burial). In (1) + (3) (silicate weathering) CO₂ is permanently removed to the geological reservoir. This CO₂ must be balanced by mantle (/volcanic) out-gassing on the very long term.

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Snowball Earth



adapted from; *Hoffman and* Schrag [2002]

Snowball Earth



adapted from; *Hoffman and Schrag* [2002]

All very well, but ...





Q 'snowball' or 'slushball'?

i.e., was the equatorial ocean ice-free, or frozen from pole-to-pole during the glacial? This has profound geochemical and climatic implications, and will constrain the mechanims responsible for going into and coming out of the glaciation.

From: Fairchild and Kennedy [2007]

All (numerical climate) models are wrong. Some may be useful.

ATM	SEA- ICE	OCN	ICE- SHT	<i>p</i> CO₂ threshold	(conclusions)
				1700 ppm	(snowball)
				<40 ppm	snowball unlikely
				130 ppm	slushball probable
				<340 ppm	slushball probable
				n/a	no snowball
				n/a	snowball less likely
				130 ppm	(snowball)
				130 ppm	snowball more likely
				500 - 990	slushball unlikely
				1800 ppm	(snowball)
				<149, 250	(snowball)
				200 ppm	(snowball)
		ATM SEA Image: Constraint of the second state	ATM CE OCN CC OCN COCN COCN COCN COCN COCN COC	ATM SEA- ICE OCN ICE- SHT I	ATM SEA ICE OCN ICE SHT PCO2 threshold I I 1700 ppm <40 ppm

KEY:

'ADVANCED' e.g. 3D GCM, thermodynamic sea-ice **'INTERMEDIATE'** e.g. 2D EBM, seasonal mixed layer ocean '**BASIC'** e.g. 1D EBM, slab ocean



<image>



CO₂ chemistry in seawater

From: Barker and Ridgwell [2012]

ocean

http://www.nature.com/scitable/knowledge/library/ ocean-acidification-25822734



CO₂ chemistry in seawater



CO₂ chemistry in seawater

ocean



carbonate ion

CO₂ chemistry in seawater

ocean



CO₂ chemistry & mineral phases



Aragonite: less stable orthorhombic polymorph (e.g., many corals, pteropods)



Calcite: more stable

(and more abundant) trigonal polymorph (e.g., coccolithophorides, foraminifera)



CO₂ chemistry & mineral phases

The addition of CO_2 to seawater results in a decrease in carbonate ion (CO_3^{2-}) concentration and 'ocean acidification'. A decrease in CO_3^{2-} , in turn, suppresses the stability of $CaCO_3$, defined by its saturation state:

 $\Omega = [Ca^{2+}] \times [CO_3^{2-}]/k$

⇒ The thermodynamic efficiency of precipitating $CaCO_3$ is a function of $[CO_3^{2-}]$ (and carbonate 'saturation').



See: GEOG 'World in crisis' lecture on 'Ocean Acidification' (PDF available from www.seao2.org/teaching.html)



Cao and Caldeira [2008]















NOTE: ocean composition format; [mean alkalinity, mean DIC] (µmol kg⁻¹)



Snowball

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Snowball Earth





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From: Hoffman and Schrag [2002]

