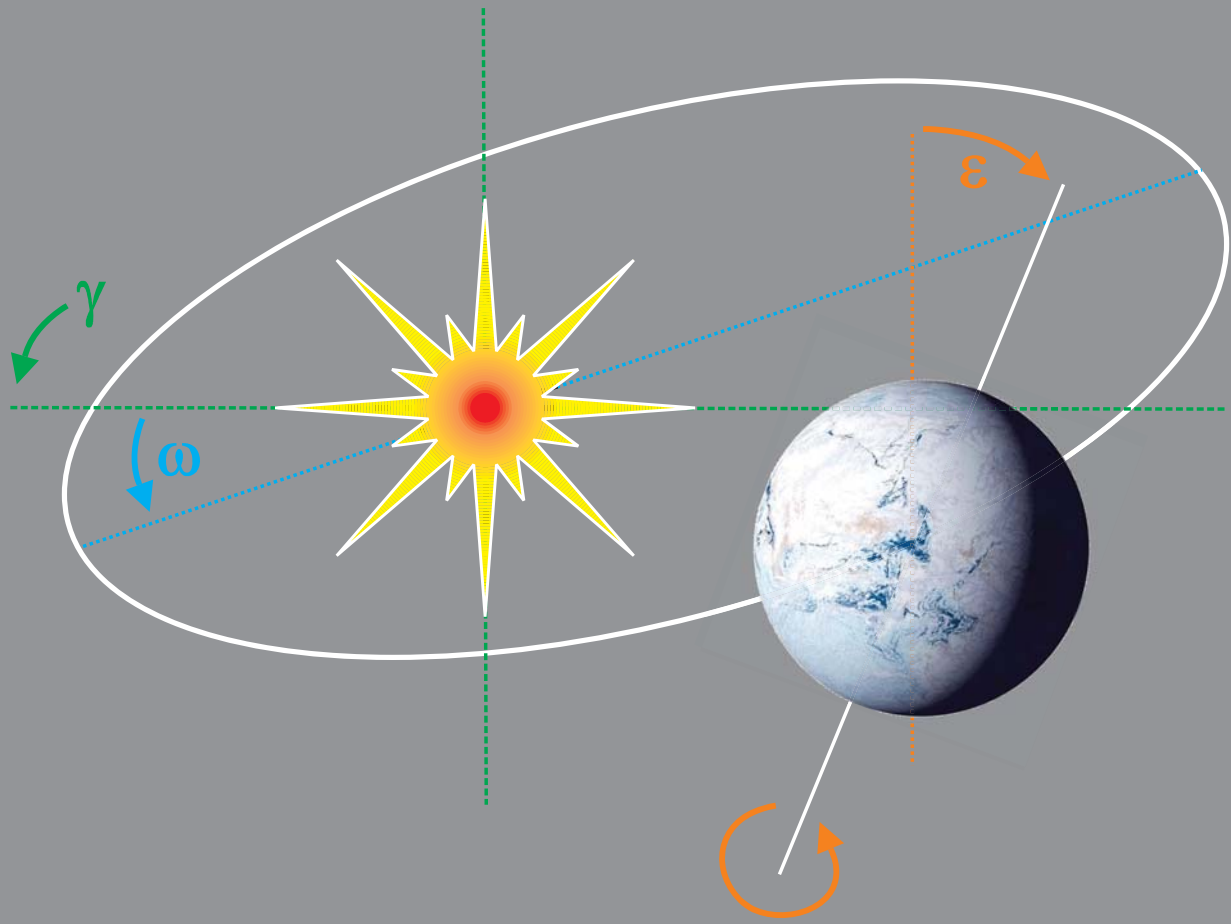


Snowball Earth

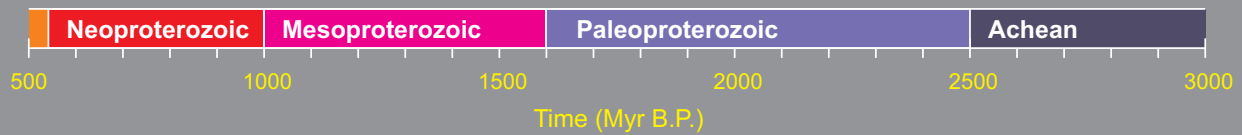


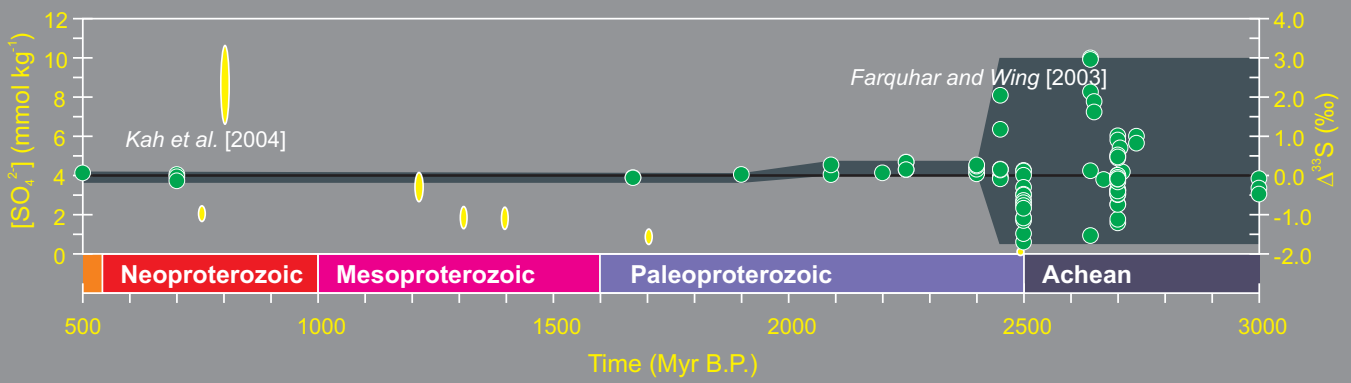
Fairchild and Kennedy [2007] (more recent and slightly more neutral/contrarian review)

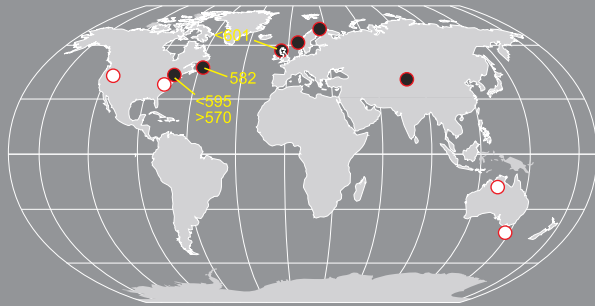
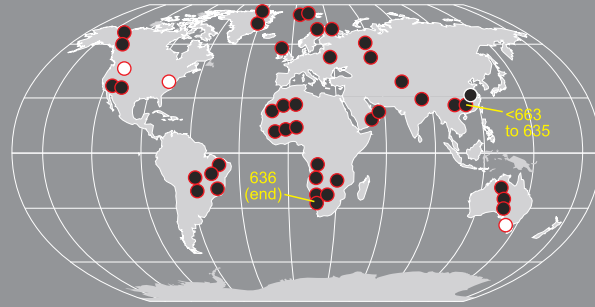
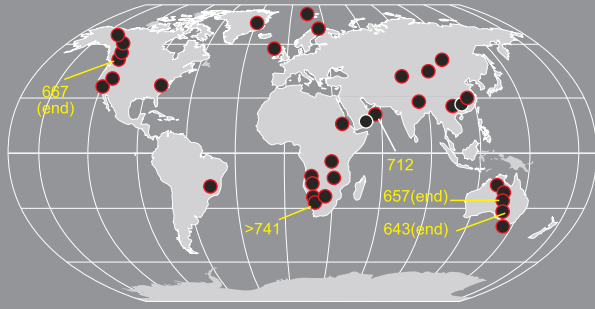
Hoffman and Schrag [2002] ('pro' snowball Earth hypothesis review)

Ridgwell and Kennedy [2004] (a different and more carbon cycle focussed view-point)

The Neoproterozoic: Gateway to a metazoan-dominated, oxygenated, 'modern-like' biosphere?



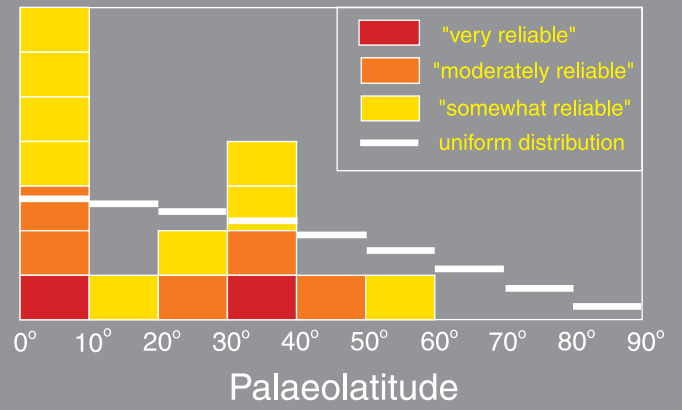
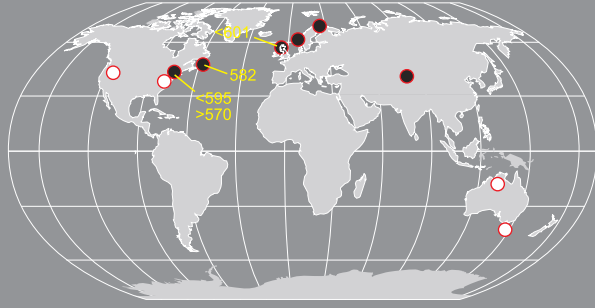
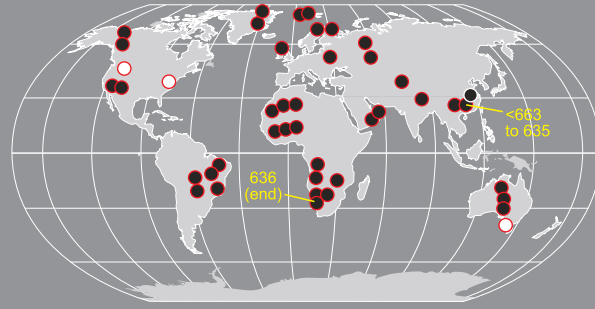
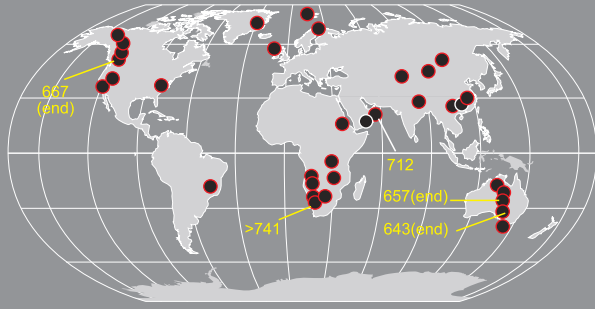




From: Fairchild and Kennedy [2007]

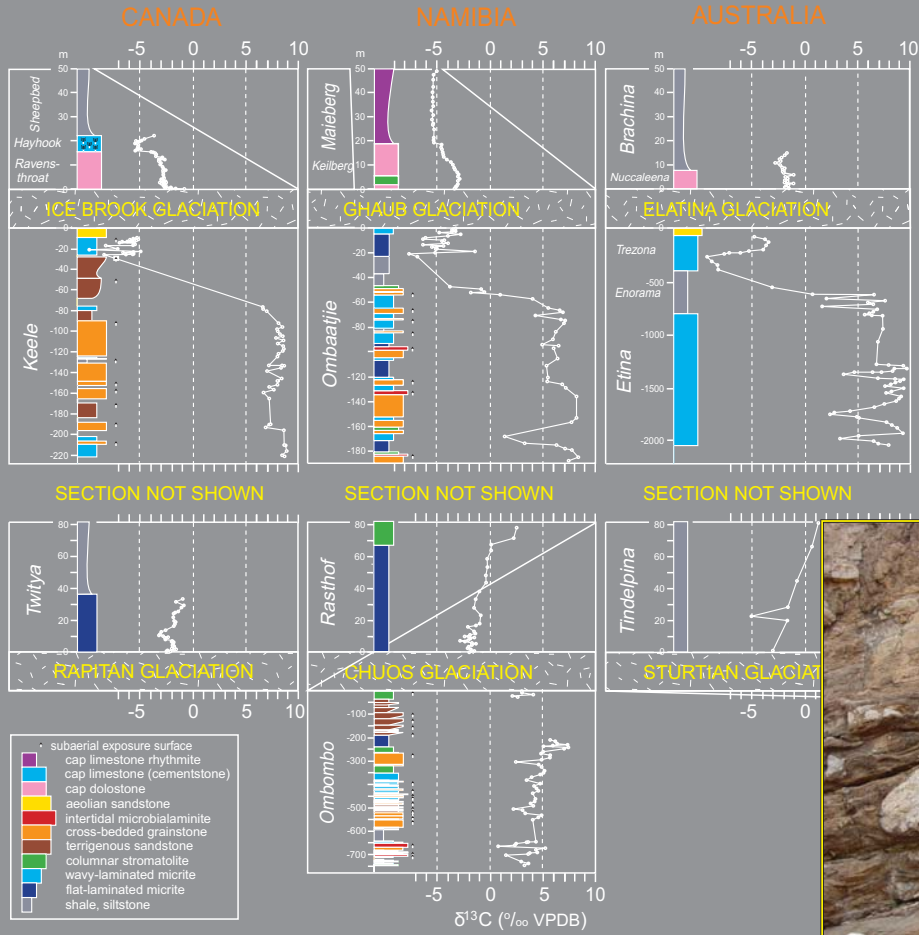


From: Hoffman and Schrag [2002]

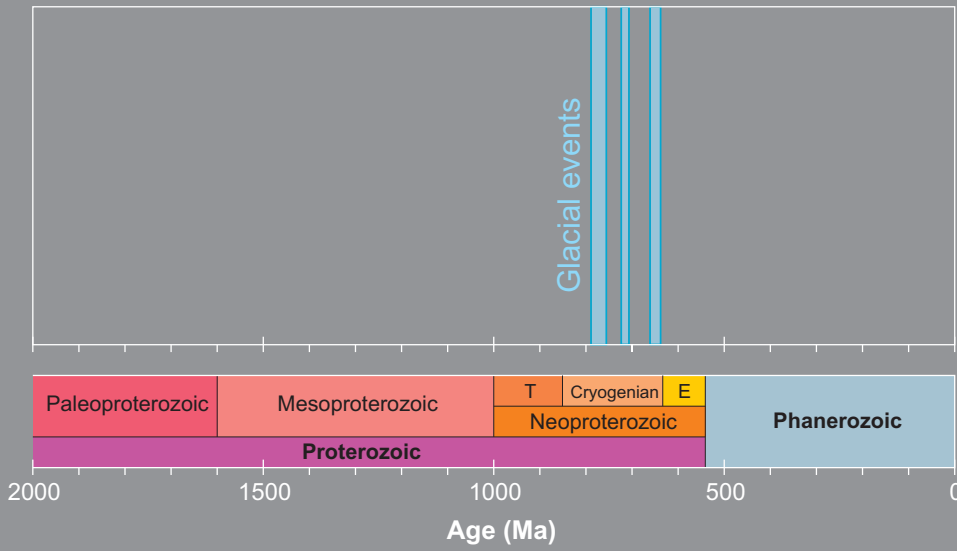


From: Fairchild and Kennedy [2007]

From: Hoffman and Schrag [2002]



From: Hoffman and Schrag [2002]

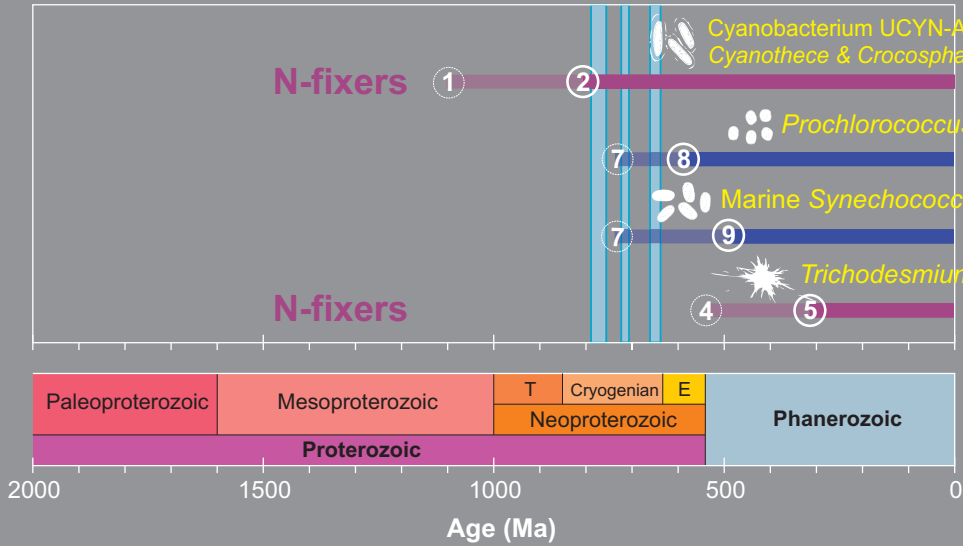


Sanchez-Baracaldo et al. [2014]

Low fixed N supply to the open ocean
Low open ocean primary production

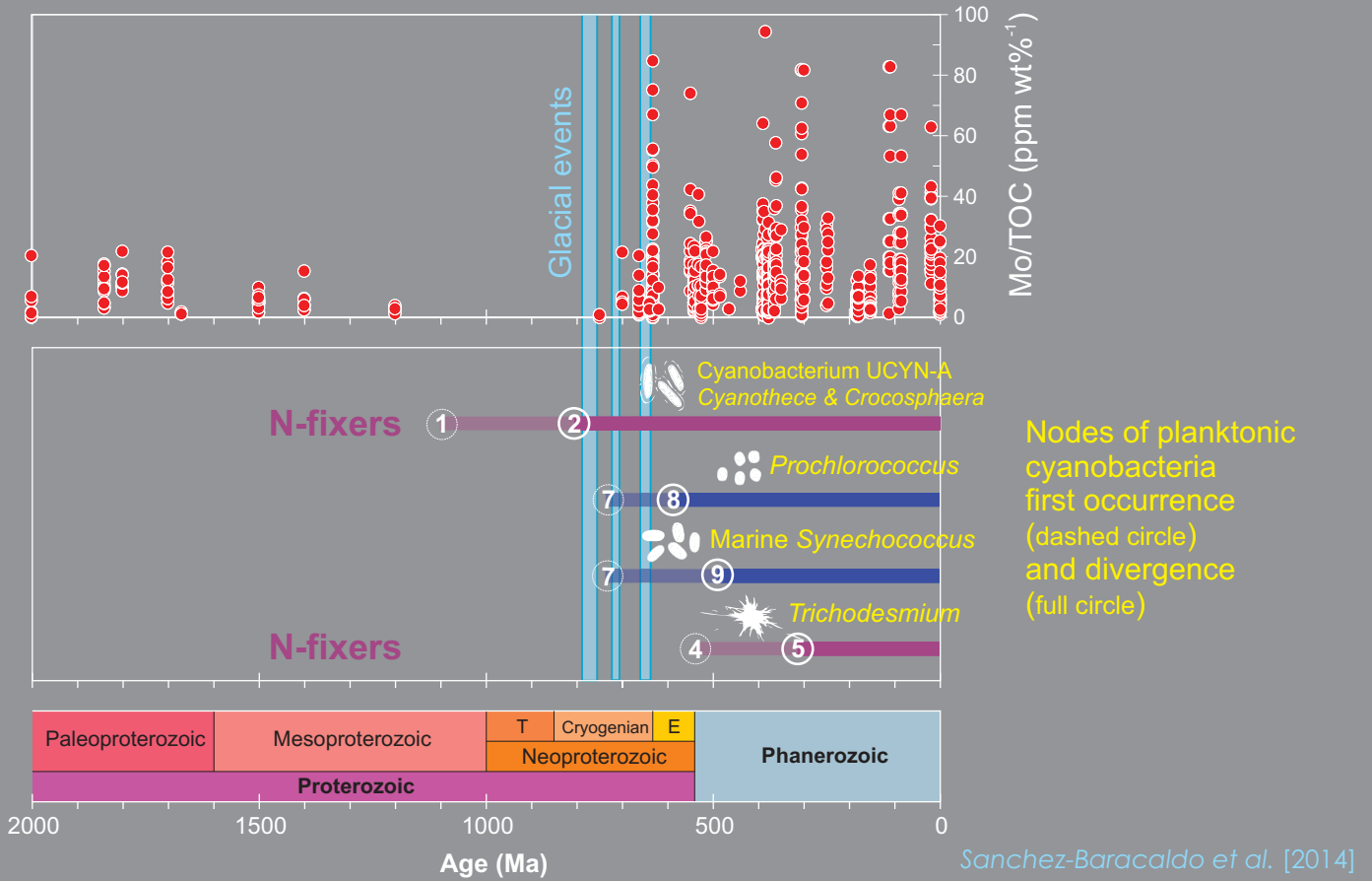
Transitional interval

High diversity of N fixers
High primary production

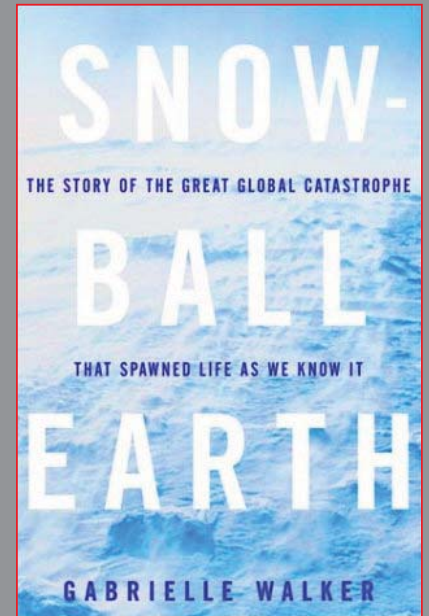


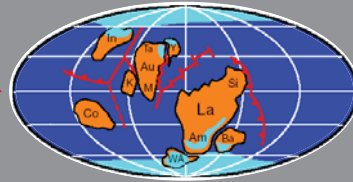
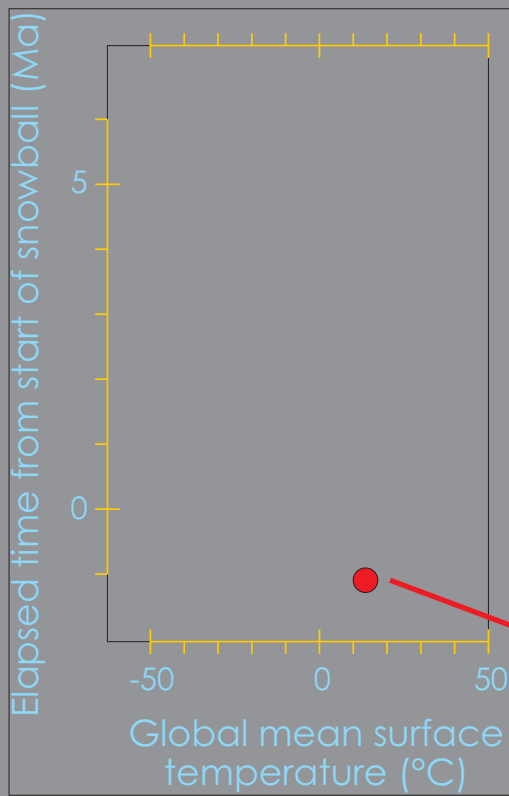
Nodes of planktonic cyanobacteria first occurrence (dashed circle) and divergence (full circle)

Sanchez-Baracaldo et al. [2014]

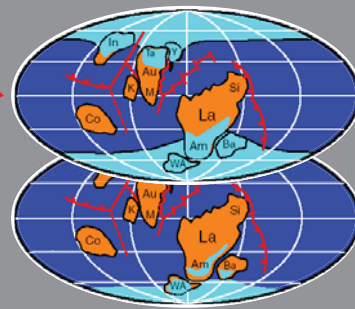
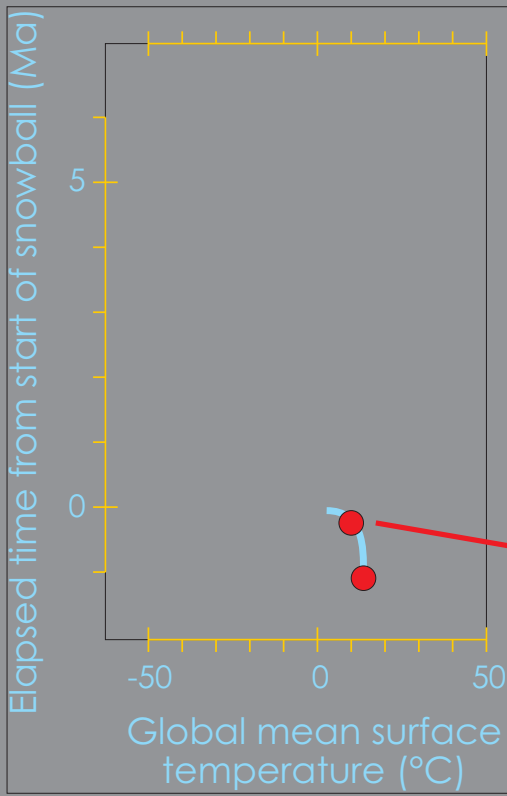


Hoffman et al. [1998] (*Science* 281)

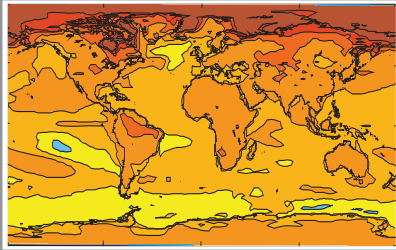




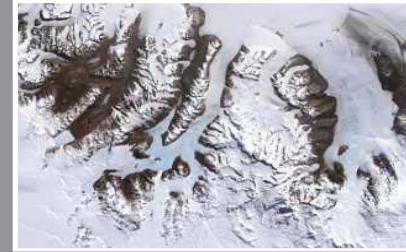
adapted from; Hoffman and Schrag [2002]



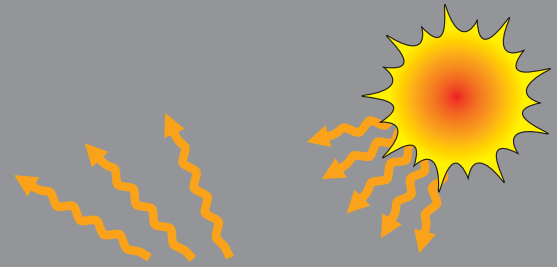
adapted from; Hoffman and Schrag [2002]



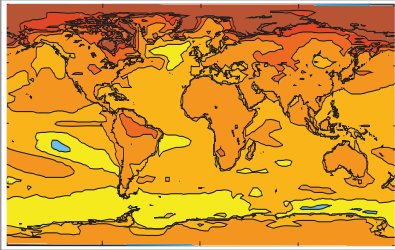
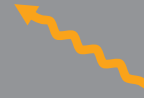
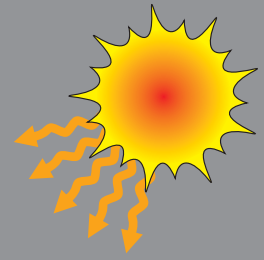
Temperature



Snow cover

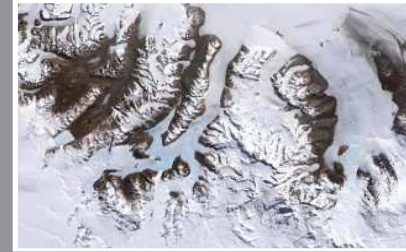


positive "ice-albedo" feedback



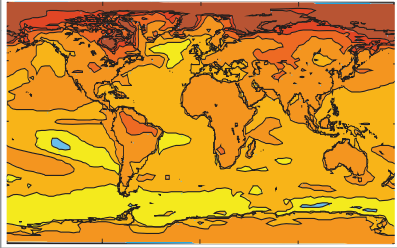
Temperature

= $-1/2^{\circ}\text{C}$

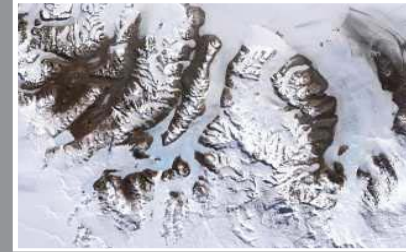


Snow cover



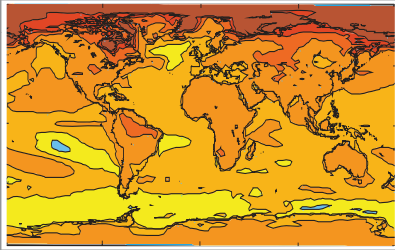
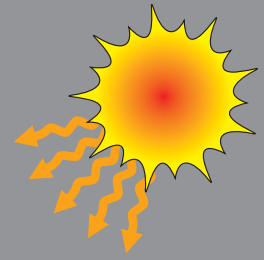


Temperature

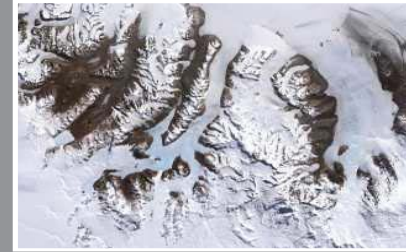


Snow cover

TOTAL CHANGE = $-1/2^{\circ}\text{C}$



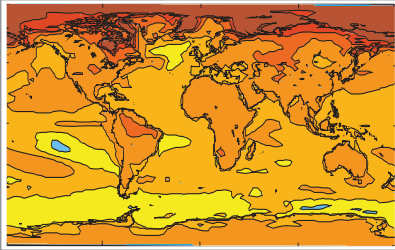
Temperature



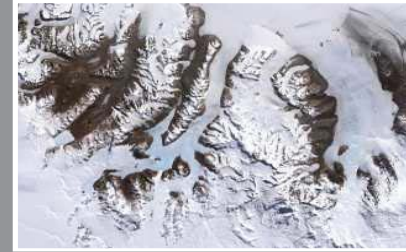
Snow cover



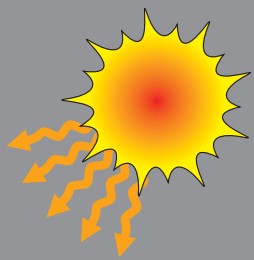
TOTAL CHANGE = $-1/2^{\circ}\text{C}$ - $1/4^{\circ}\text{C}$



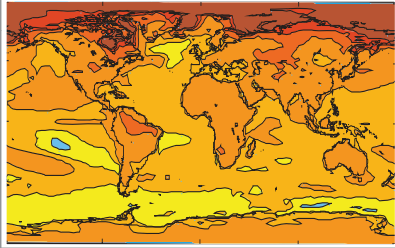
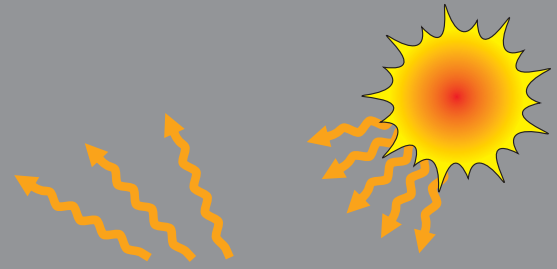
Temperature



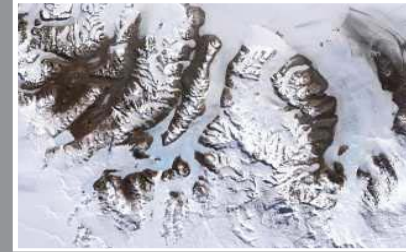
Snow cover



TOTAL CHANGE = $-1/2^{\circ}\text{C}$ - $1/4^{\circ}\text{C}$



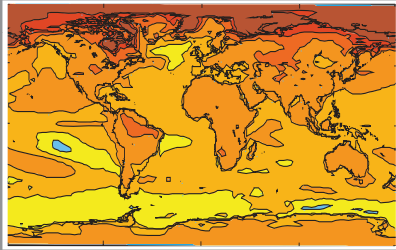
Temperature



Snow cover



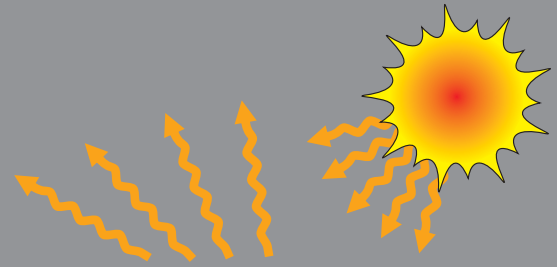
TOTAL CHANGE = $-1/2^{\circ}\text{C}$ - $1/4^{\circ}\text{C}$ - $1/8^{\circ}\text{C}$



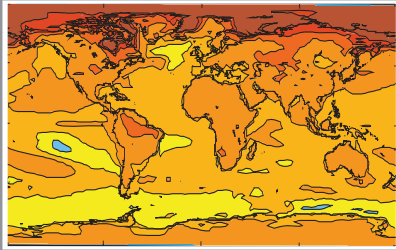
Temperature



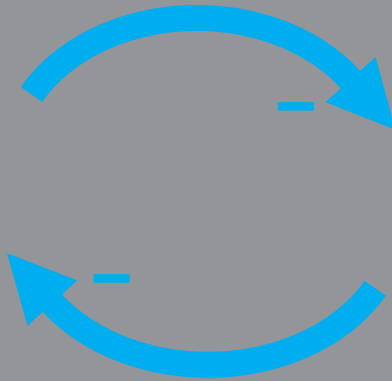
Snow cover



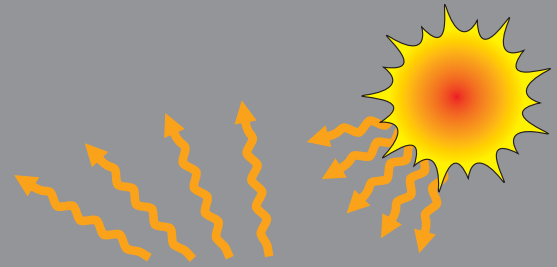
$$\text{TOTAL CHANGE} = -1/2^{\circ}\text{C} - 1/4^{\circ}\text{C} - 1/8^{\circ}\text{C} - 1/16^{\circ}\text{C} - \dots$$



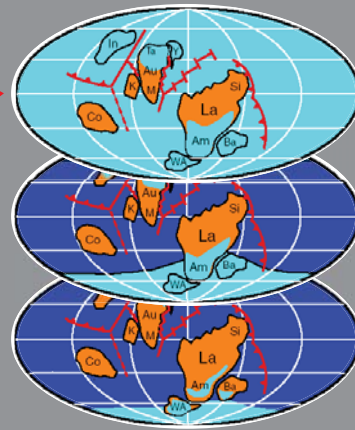
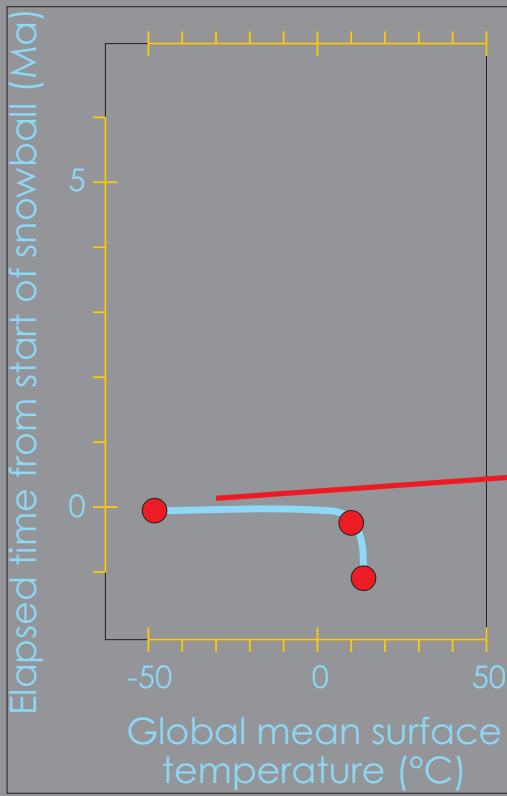
Temperature



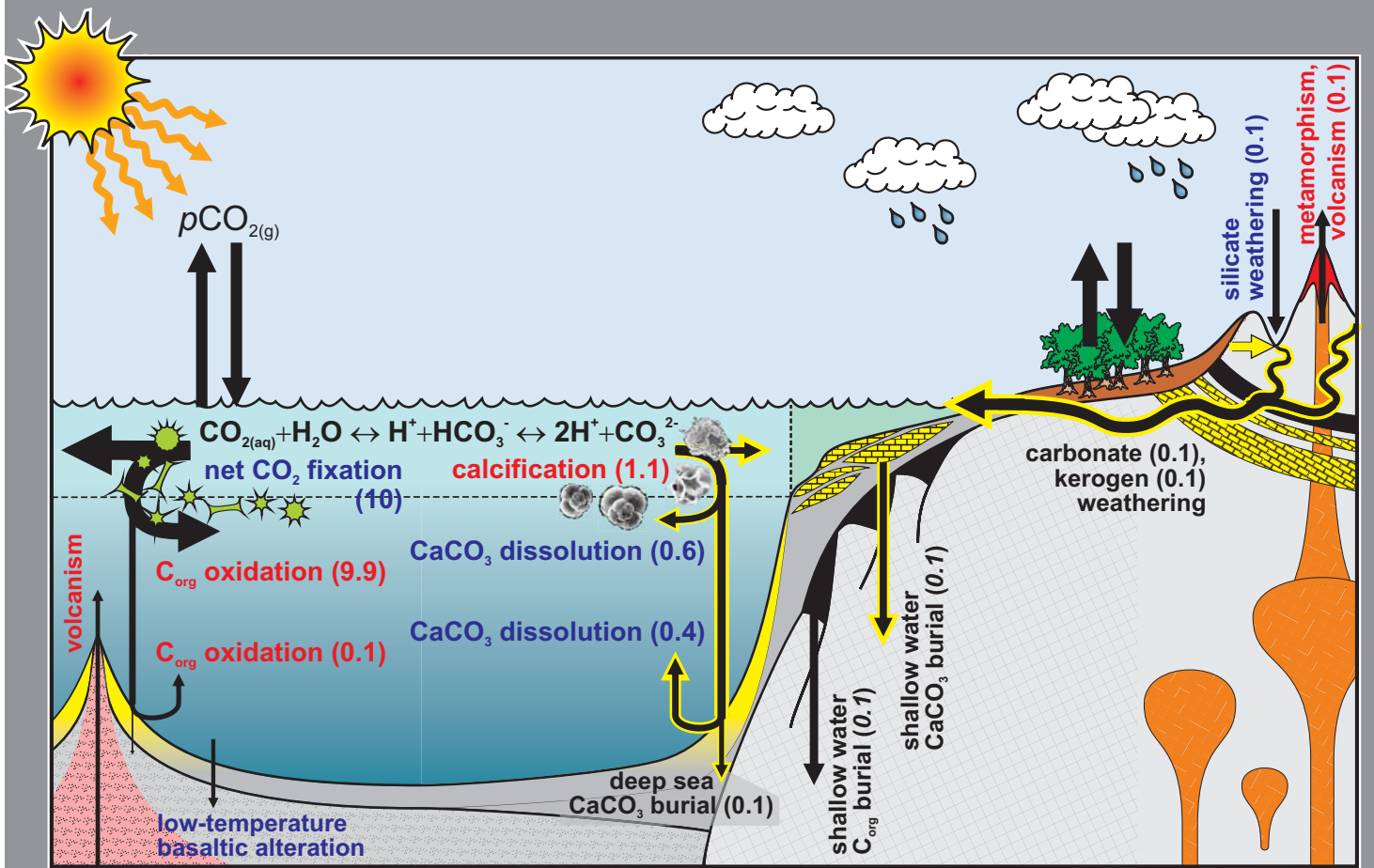
Snow cover



TOTAL CHANGE = $-1^{\circ}\text{C} - 2^{\circ}\text{C} - 4^{\circ}\text{C} - 8^{\circ}\text{C}$
-



adapted from; Hoffman and Schrag [2002]



Terrestrial weathering can be (approximately equally) divided into carbonate (CaCO_3) and calcium-silicate (' CaSiO_3 ') weathering:



Ultimately, the (alkalinity: Ca^{2+}) weathering products must be removed through carbonate precipitation and burial in marine sediments:



It can be seen that in (2) + (3), that the CO_2 removed (from the atmosphere) during weathering, is returned upon carbonate precipitation (and burial). In (1) + (3) (silicate weathering) CO_2 is permanently removed to the geological reservoir. This CO_2 must be balanced by mantle (/volcanic) out-gassing on the very long term.

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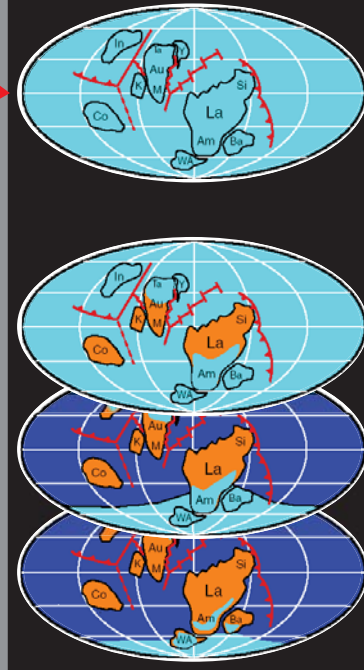
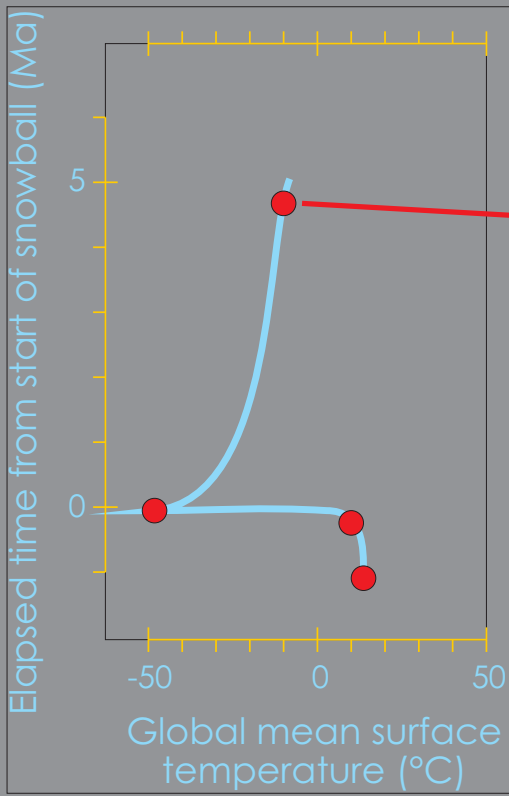
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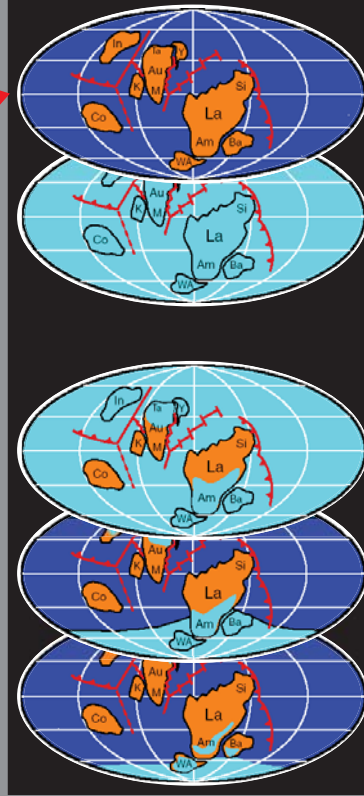
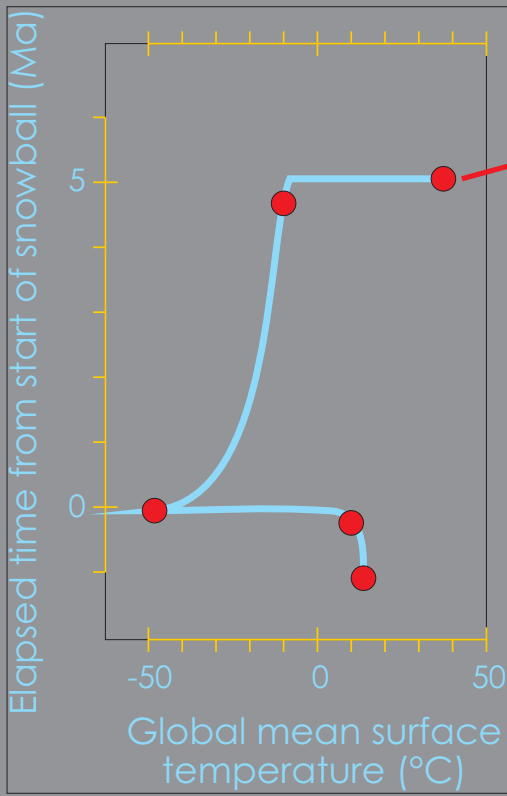
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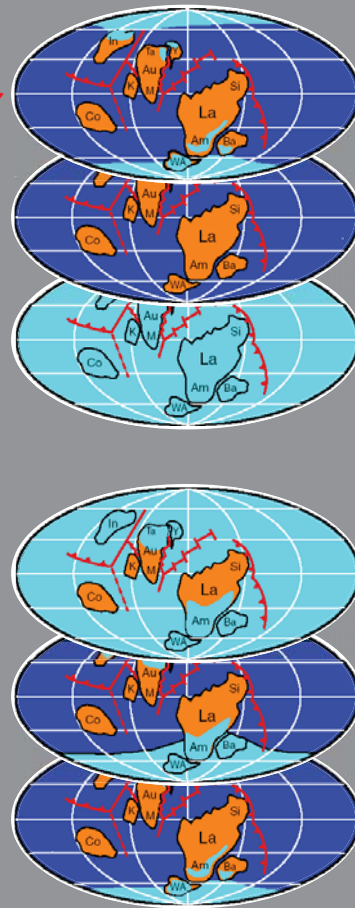
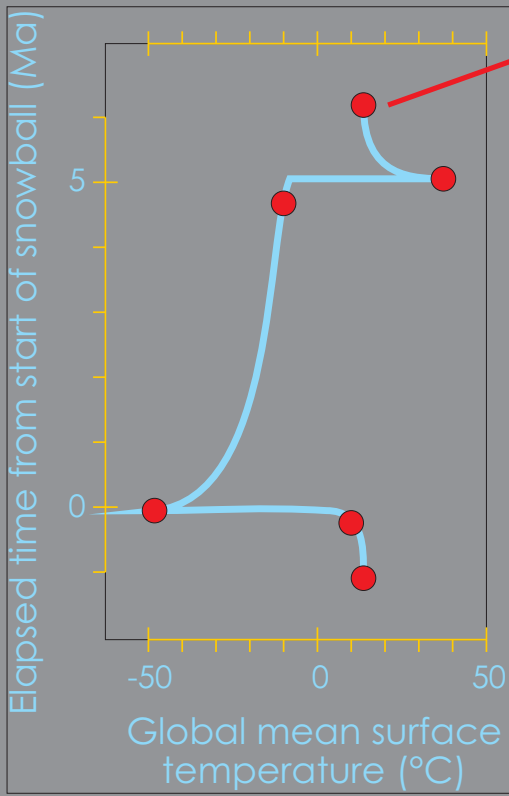
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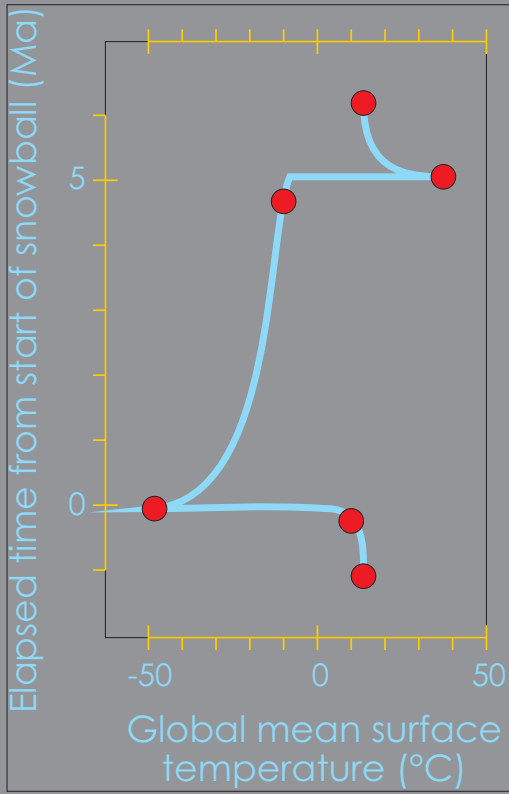
adapted from; Hoffman and Schrag [2002]



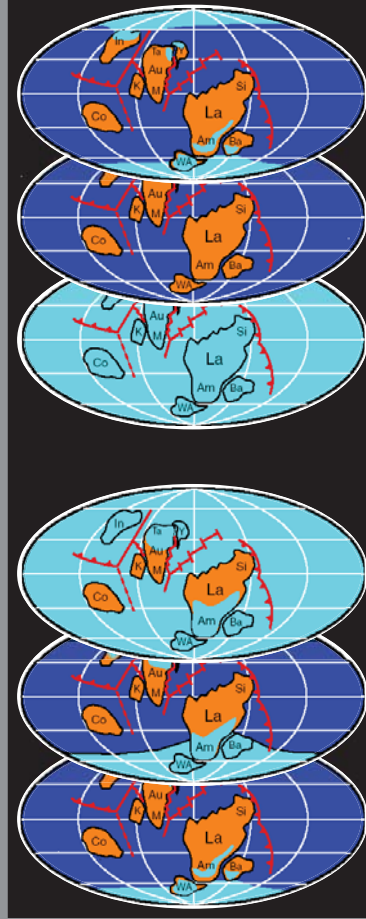
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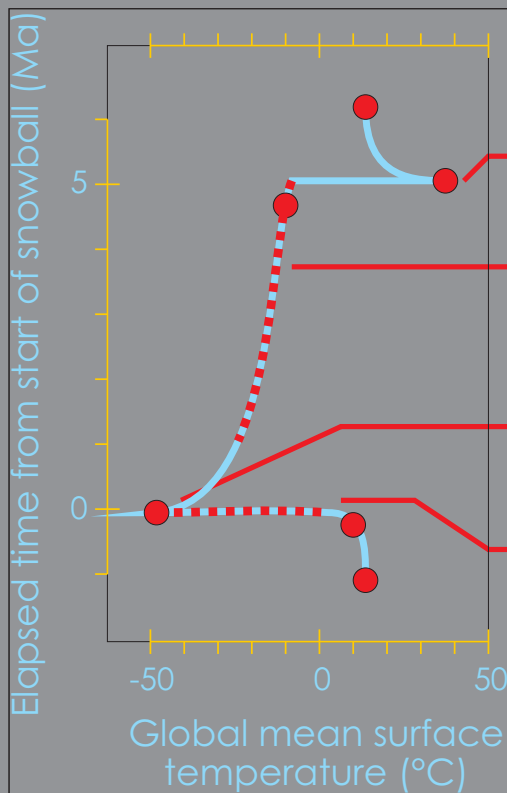


Progression of exotic states of the carbon-climate system



adapted from; Hoffman and Schrag [2002]

All very well, but ...



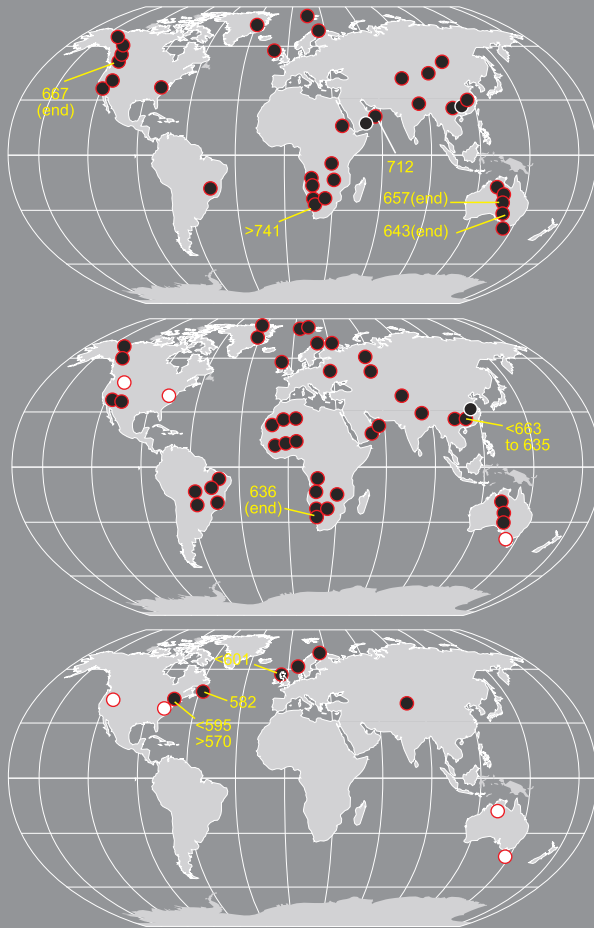
No proxy evidence for: (1) >0.1 bar CO_2 in the atmosphere, (2) the inferred mean global surface temperatures of ca. 50°C , or (3) intense weathering rates in the immediate aftermath of deglaciation.

No proxy evidence of complete cessation of weathering on land (required to build up CO_2 in the atmosphere).

No direct proxy evidence for a completely ice-covered ocean.

How does complex life persist?

Some coupled ocean-atmosphere models do not find an ice-albedo instability.



From: Fairchild and Kennedy [2007]





Q 'snowball' or 'slushball'?

i.e., was the equatorial ocean ice-free, or frozen from pole-to-pole during the glacial? This has profound geochemical and climatic implications, and will constrain the mechanisms responsible for going into and coming out of the glaciation.

All (numerical climate) models are wrong.
Some may be useful.

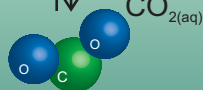
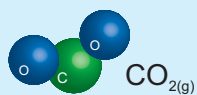
Study	ATM	SEA-ICE	OCN	ICE-SHT	pCO ₂ threshold	(conclusions)
Jenkins and Smith [1999]	Red	Red	Green	Black	1700 ppm	(snowball)
Chandler and Sohl [2000]	Red	Red	Yellow	Black	<40 ppm	snowball unlikely
Hyde et al. [2000]; Crowley et al. [2001]	Yellow	Red	Green	Red	130 ppm	slushball probable
Baum and Crowley [2001,2003]	Red	Red	Green	Black	<340 ppm	slushball probable
Poulsen et al. [2001,2]; Poulsen [2003]	Red	Yellow	Red	Black	n/a	no snowball
Bendtsen [2002]	Green	Green	Green	Black	n/a	snowball less likely
Godderis et al. [2003]	Green	Black	Black	Black	130 ppm	(snowball)
Goodman and Pierrehumbert [2003]	Diagonal (Red/Black)	Red	Diagonal (Red/Black)	Black	130 ppm	snowball more likely
Donnadieu et al. [2003]	Red	Green	Green	Diagonal (Red/Black)	500 - 990	slushball unlikely
Lewis et al. [2003,2004]	Yellow	Red	Red	Black	1800 ppm	(snowball)
Donnadieu et al. [2004a,b]	Red	Red	Yellow	Black	<149, 250	(snowball)
Edwards and Ridgwell [unpublished]	Yellow	Red	Red	Black	200 ppm	(snowball)

KEY:

-  **'ADVANCED'**
e.g. 3D GCM,
thermodynamic
sea-ice
-  **'INTERMEDIATE'**
e.g. 2D EBM,
seasonal mixed
layer ocean
-  **'BASIC'**
e.g. 1D EBM,
slab ocean
-  **DECOUPLED**



atmosphere



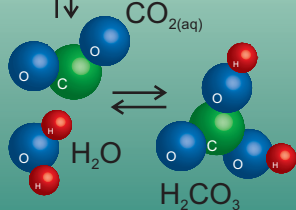
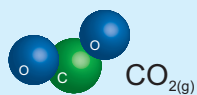
ocean

CO_2 chemistry in seawater

From: Barker and Ridgwell (2012)

<http://www.nature.com/scitable/knowledge/library/ocean-acidification-25822734>

atmosphere



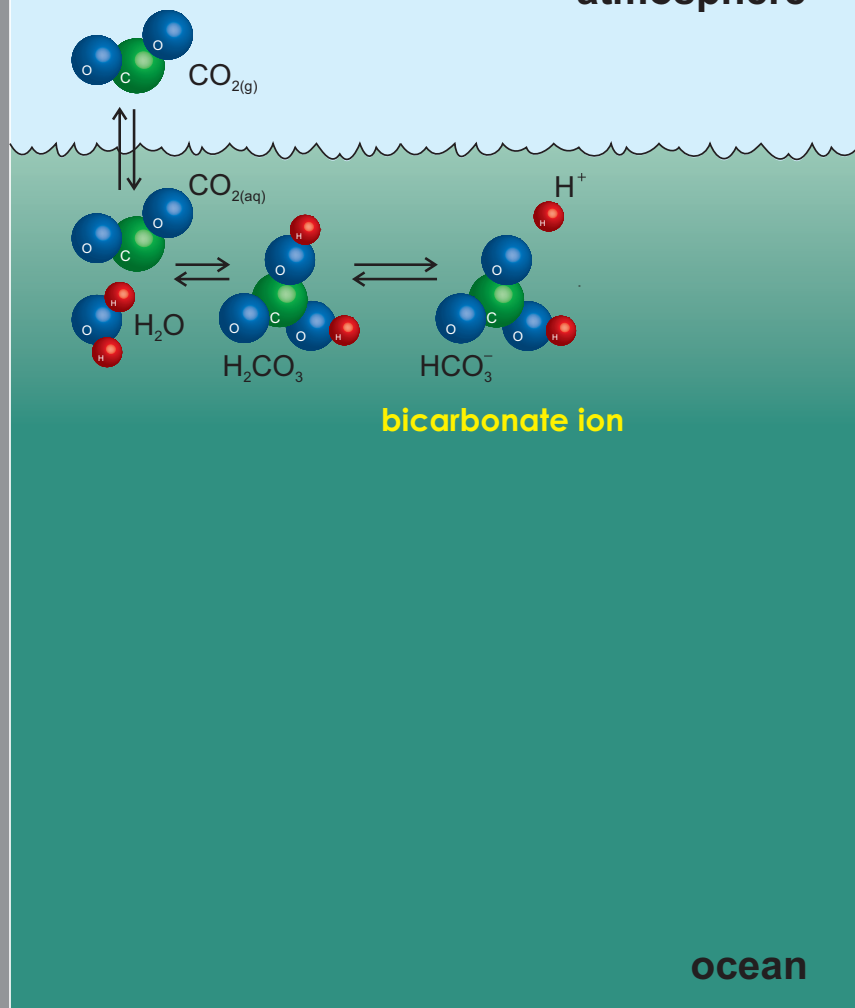
carbonic acid

ocean

CO_2 chemistry
in seawater

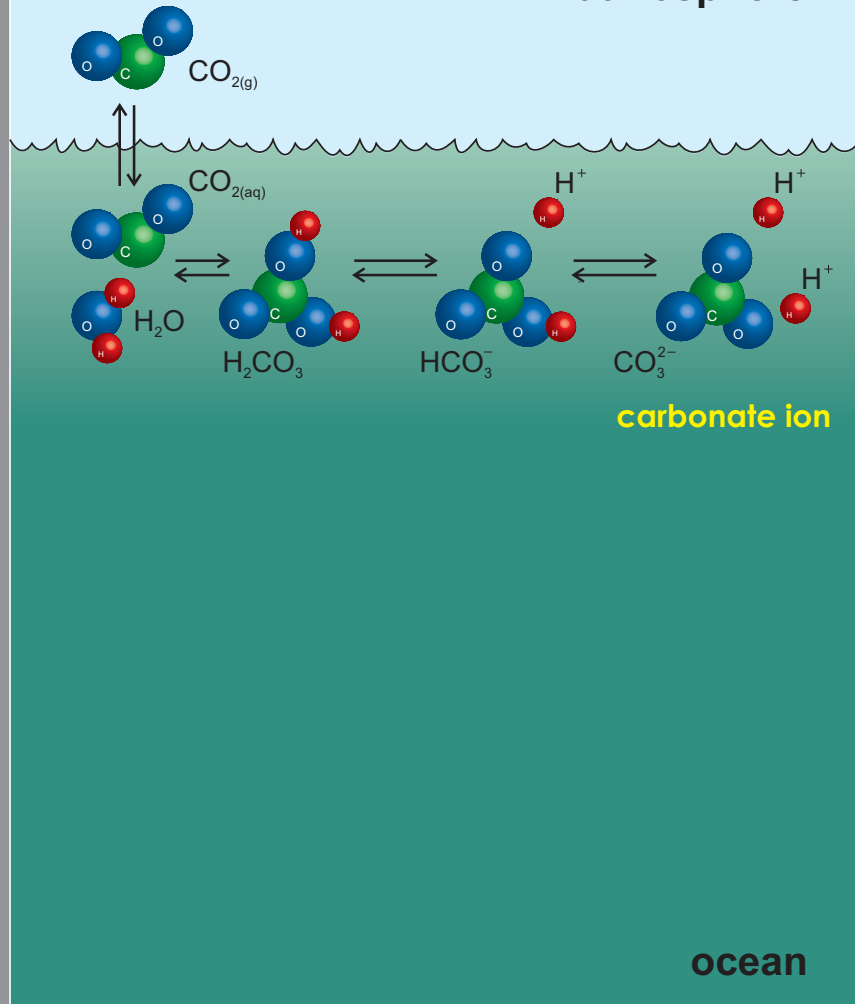
atmosphere

CO₂ chemistry
in seawater

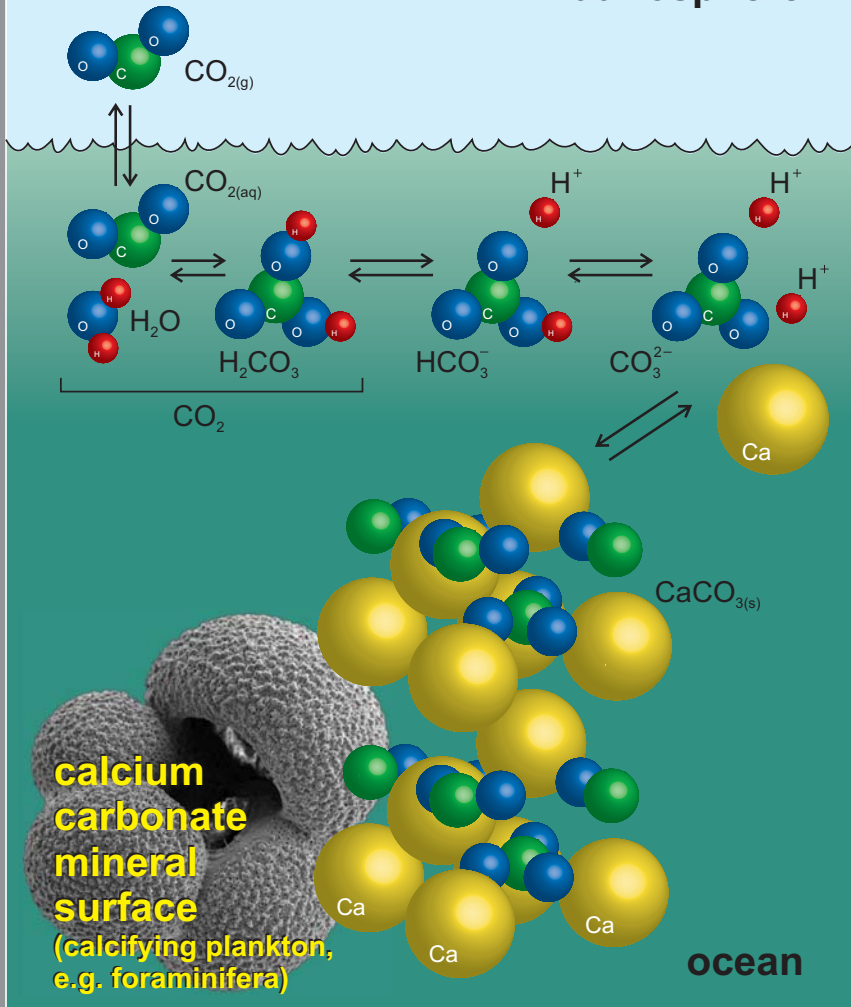


atmosphere

CO₂ chemistry
in seawater



atmosphere



calcium carbonate mineral surface
(calcifying plankton, e.g. foraminifera)

ocean

CO₂ chemistry & mineral phases

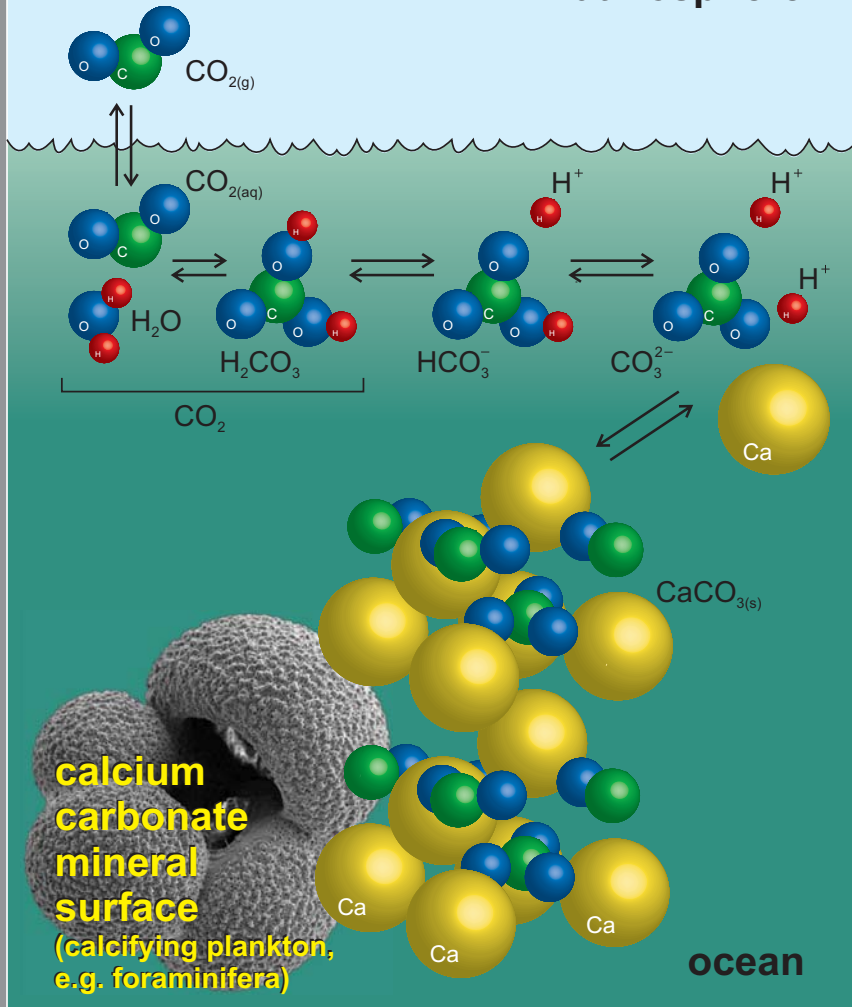


Aragonite: less stable
orthorhombic polymorph (e.g., many corals, pteropods)



Calcite: more stable
(and more abundant)
trigonal polymorph (e.g., coccolithophorides, foraminifera)

atmosphere



CO_2 chemistry & mineral phases

The addition of CO_2 to seawater results in a decrease in carbonate ion (CO_3^{2-}) concentration and 'ocean acidification'. A decrease in CO_3^{2-} , in turn, suppresses the stability of CaCO_3 , defined by its *saturation state*:

$$\Omega = [\text{Ca}^{2+}] \times [\text{CO}_3^{2-}] / k$$

⇒ The thermodynamic efficiency of precipitating CaCO_3 is a function of $[\text{CO}_3^{2-}]$ (and carbonate 'saturation').

**calcium
carbonate
mineral
surface**

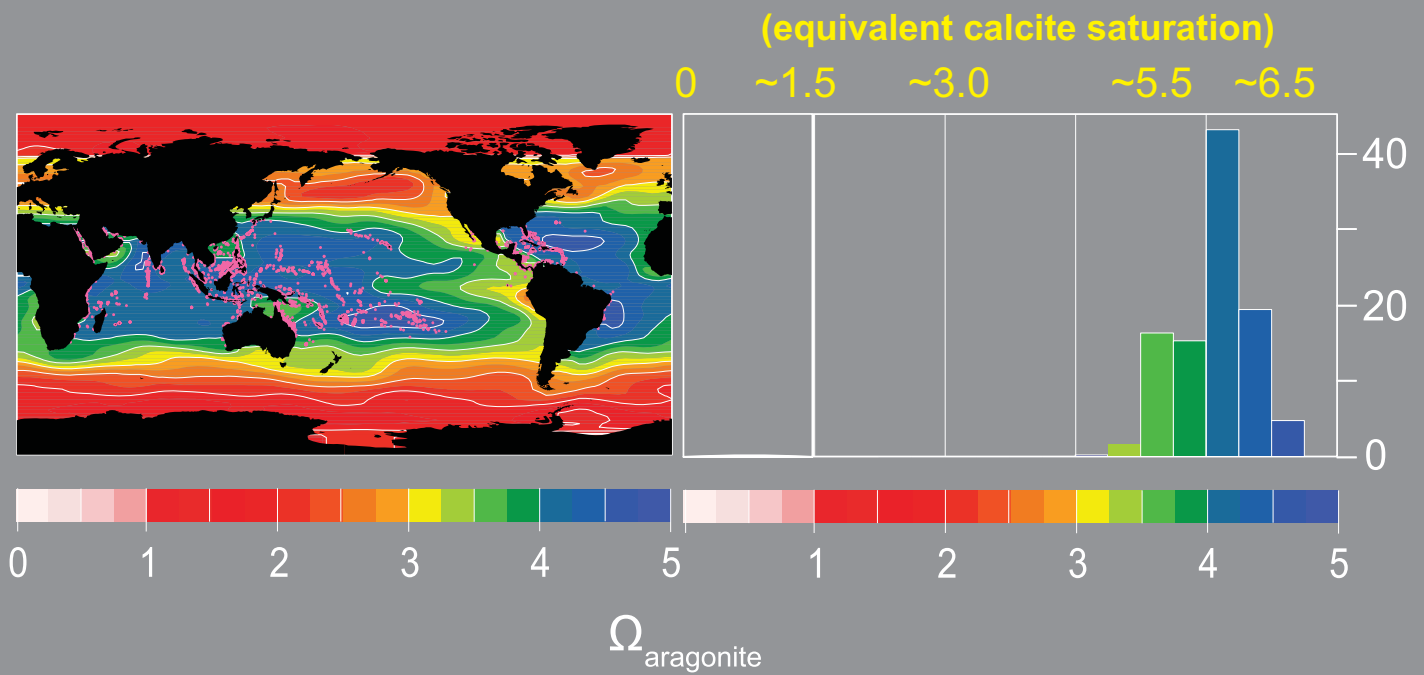
(calcifying plankton,
e.g. foraminifera)

ocean

Aqueous carbonate equilibrium; $\text{H}_2\text{O} + \text{CO}_{2(\text{aq})} + \text{CO}_3^{2-} \leftrightarrow 2\text{HCO}_3^-$
Stability of CaCO_3 defined relative to saturation state; $\Omega = [\text{Ca}^{2+}] \times [\text{CO}_3^{2-}] / k$

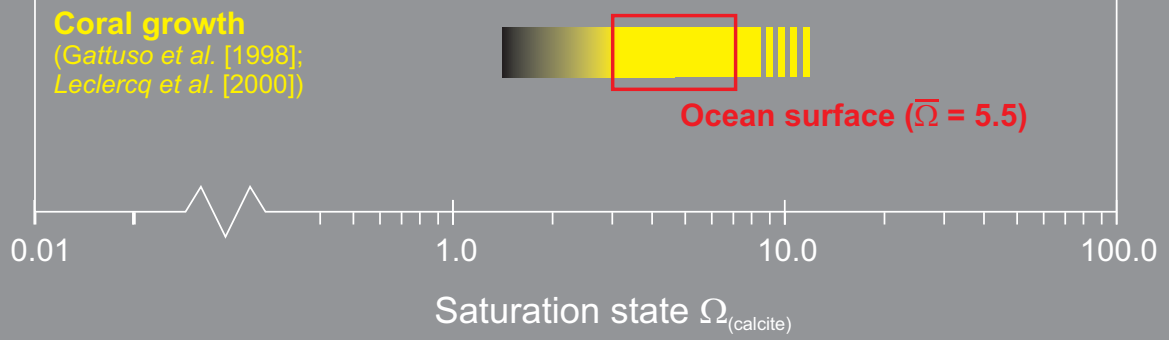


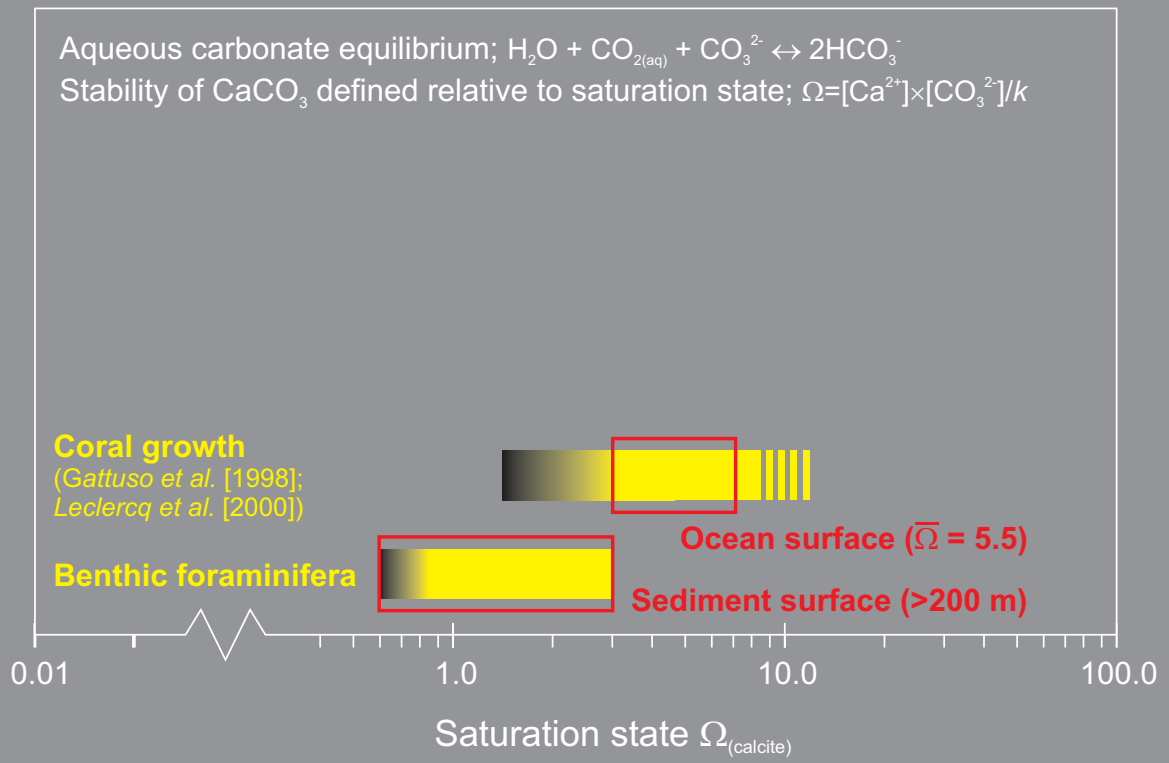
See: GEOG 'World in crisis' lecture on 'Ocean Acidification'
(PDF available from www.seao2.org/teaching.html)

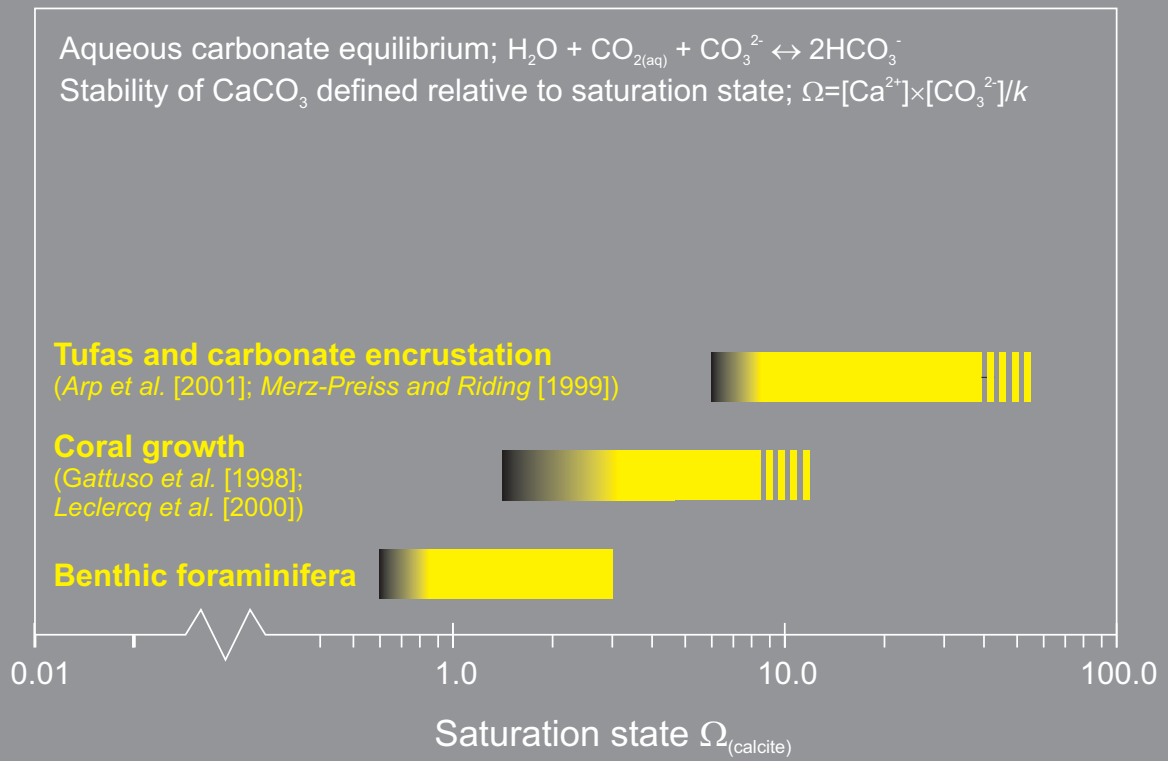


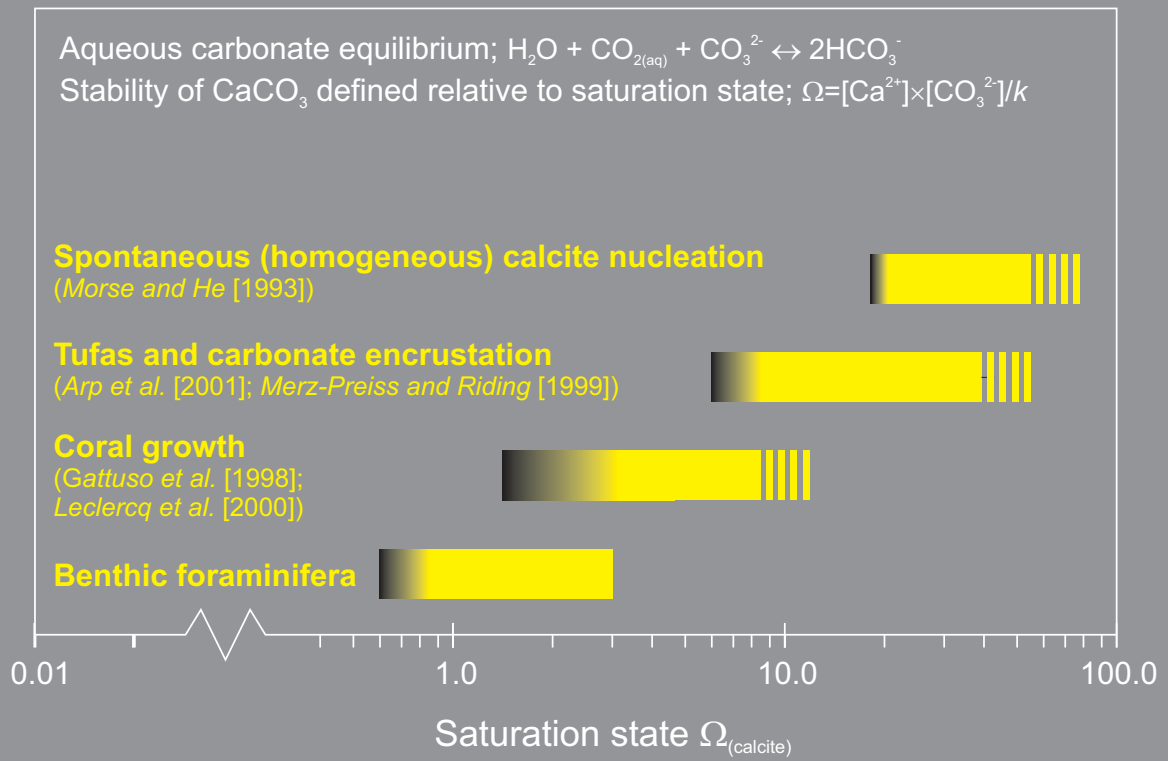
Cao and Caldeira [2008]

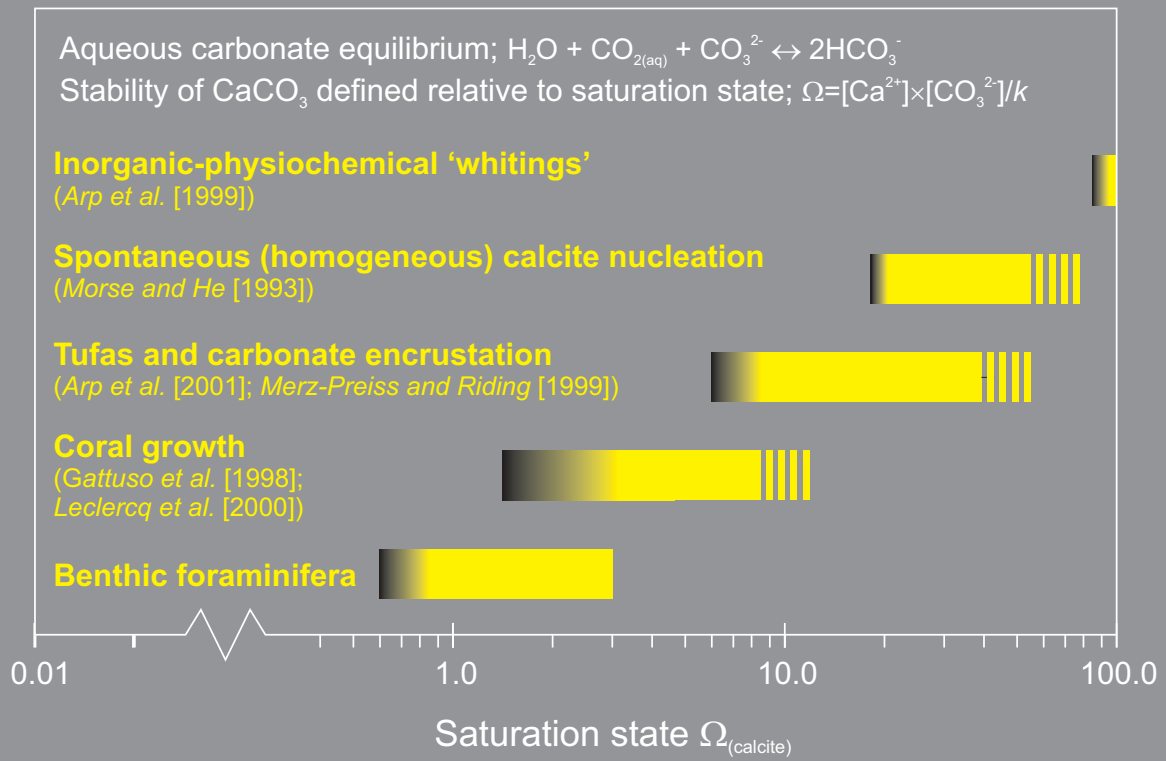
Aqueous carbonate equilibrium; $\text{H}_2\text{O} + \text{CO}_{2(\text{aq})} + \text{CO}_3^{2-} \leftrightarrow 2\text{HCO}_3^-$
Stability of CaCO_3 defined relative to saturation state; $\Omega = [\text{Ca}^{2+}] \times [\text{CO}_3^{2-}] / k$

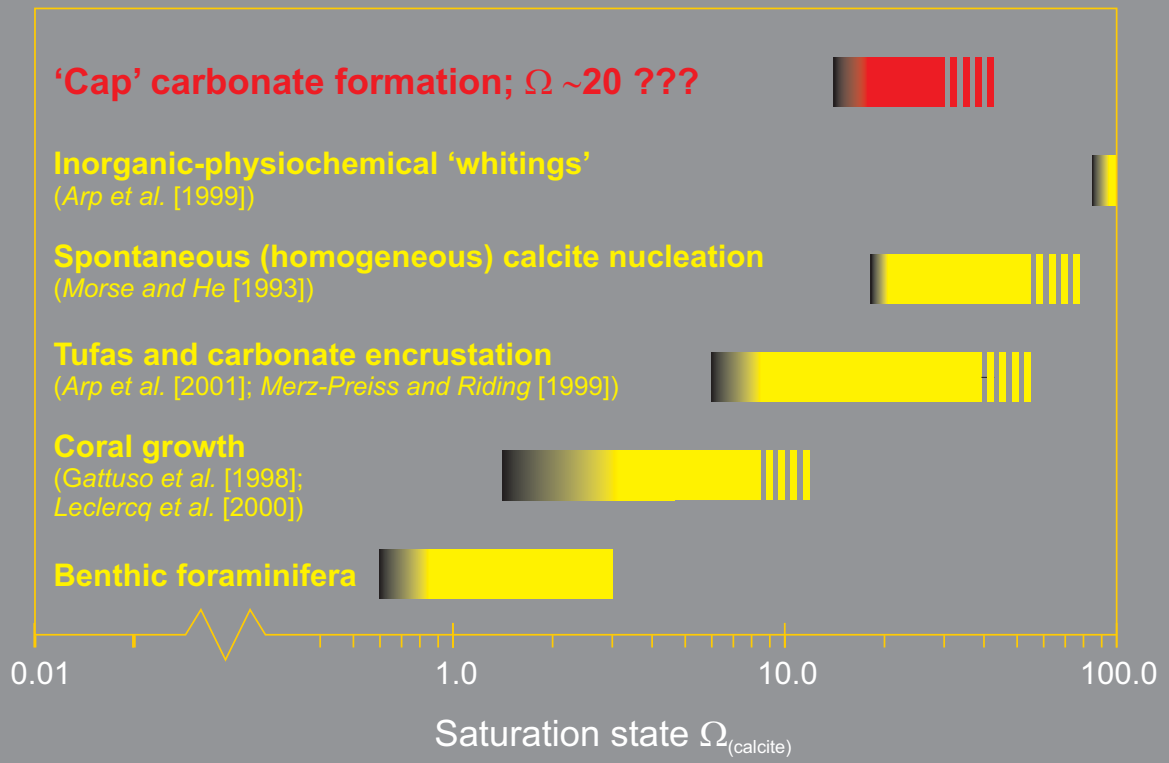


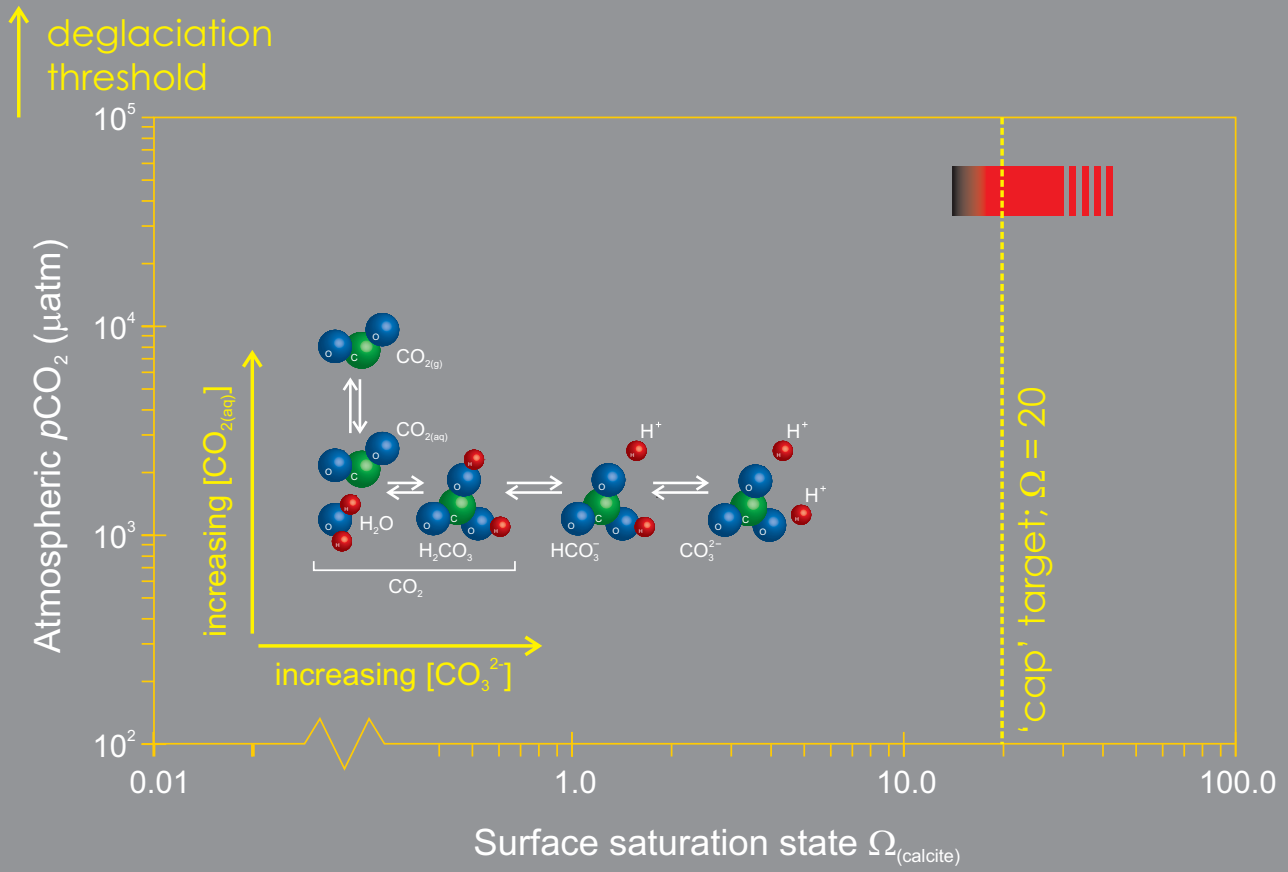




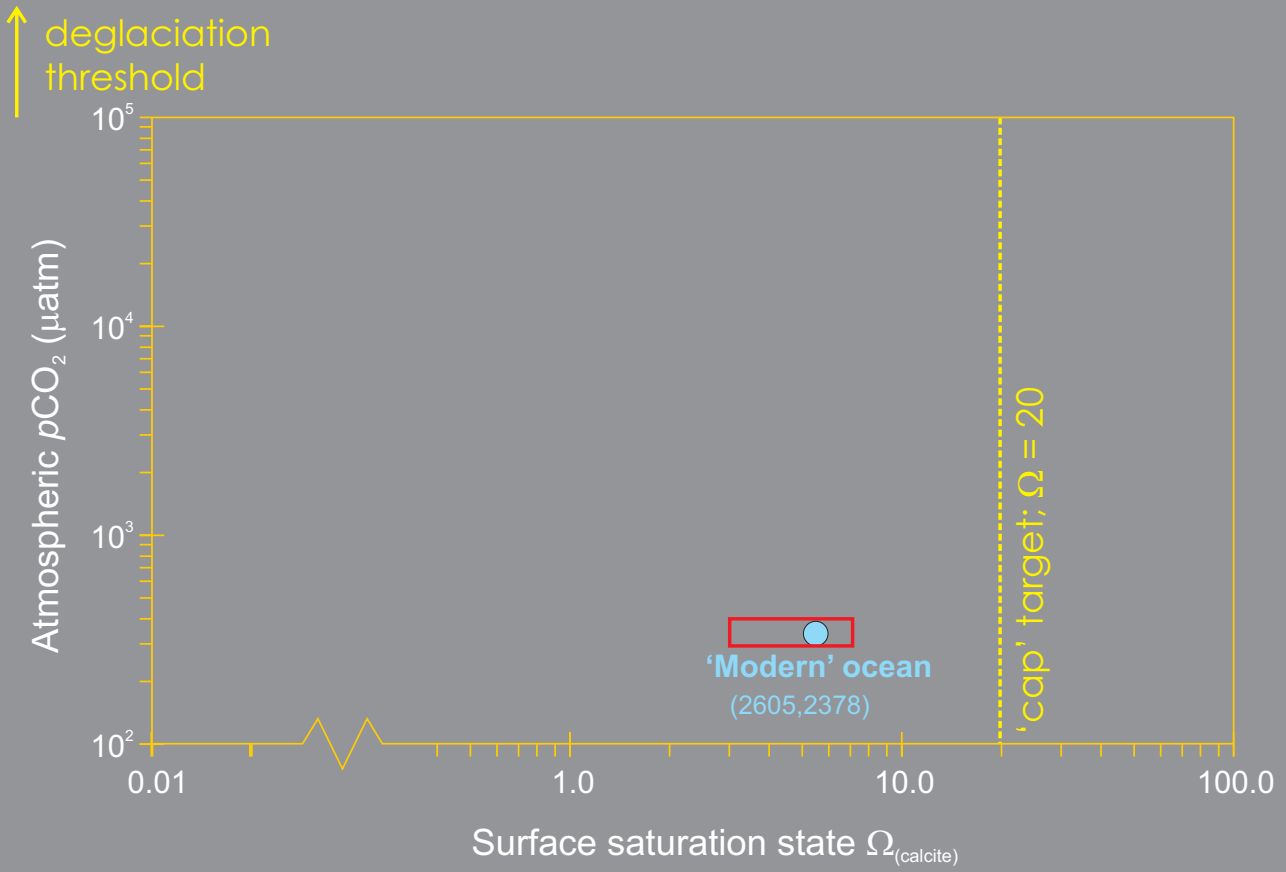




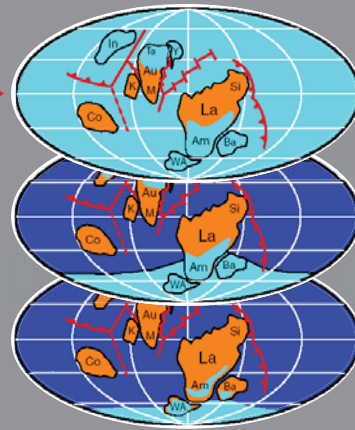
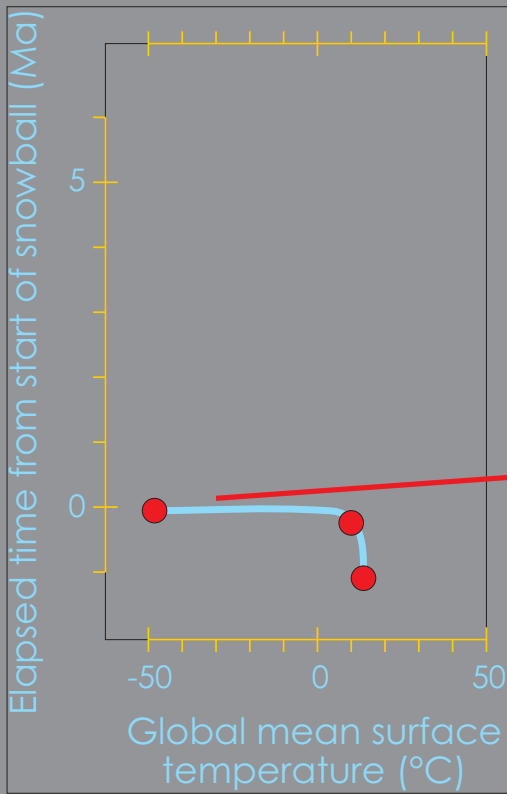




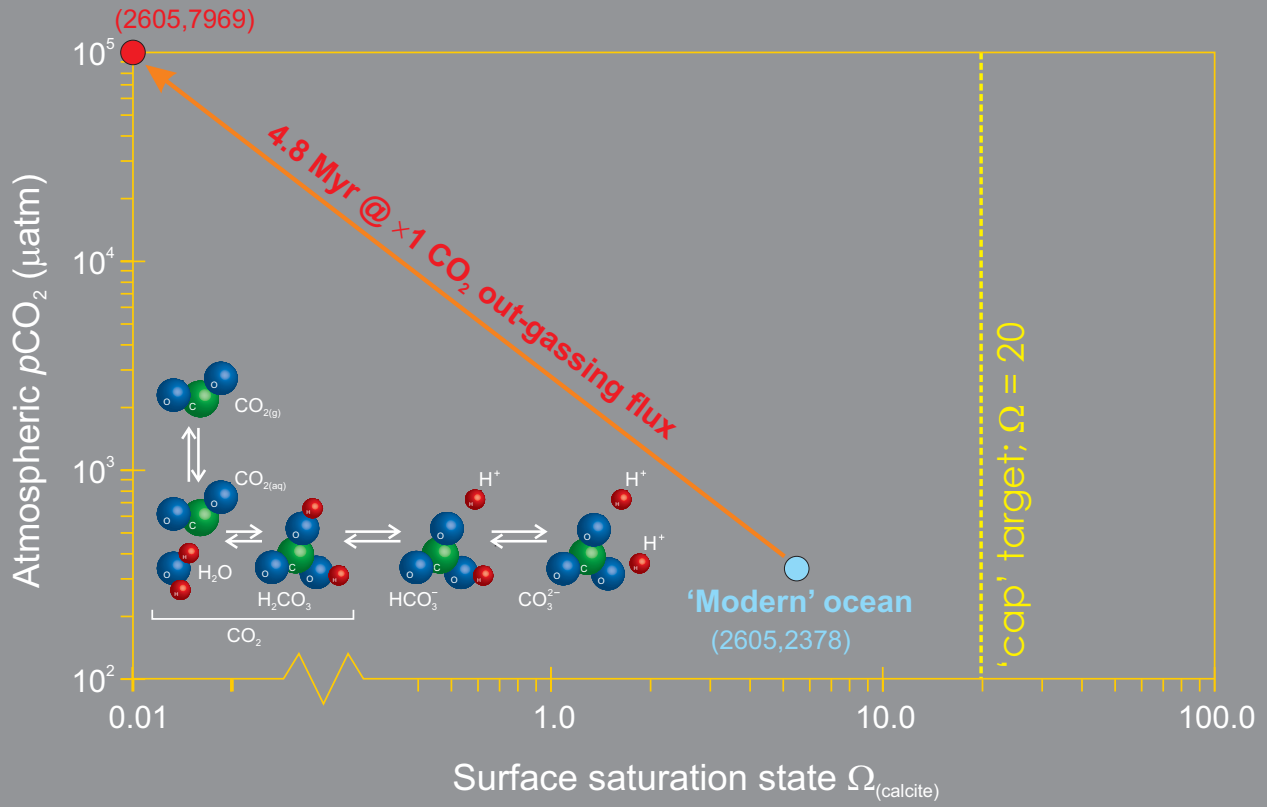
NOTE: ocean composition format;
[mean alkalinity, mean DIC] ($\mu\text{mol kg}^{-1}$)



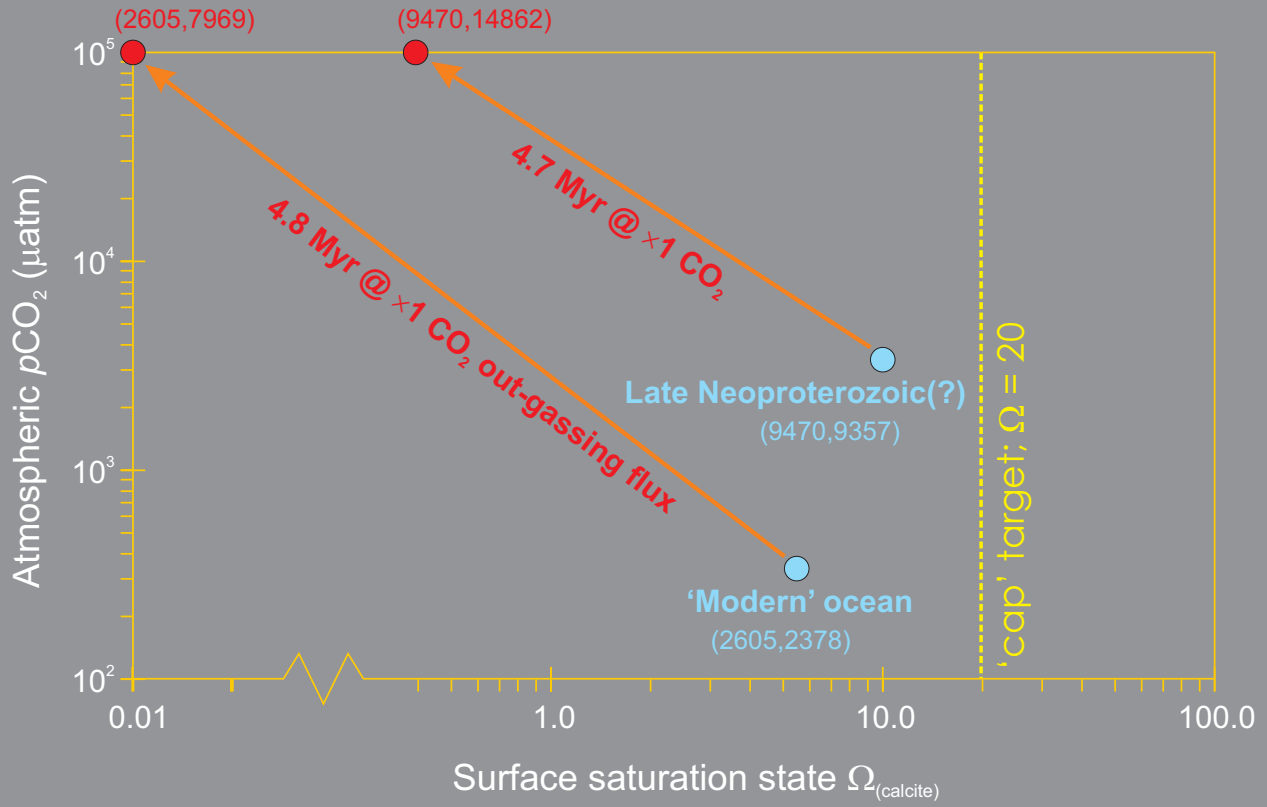
NOTE: ocean composition format;
[mean alkalinity, mean DIC] ($\mu\text{mol kg}^{-1}$)



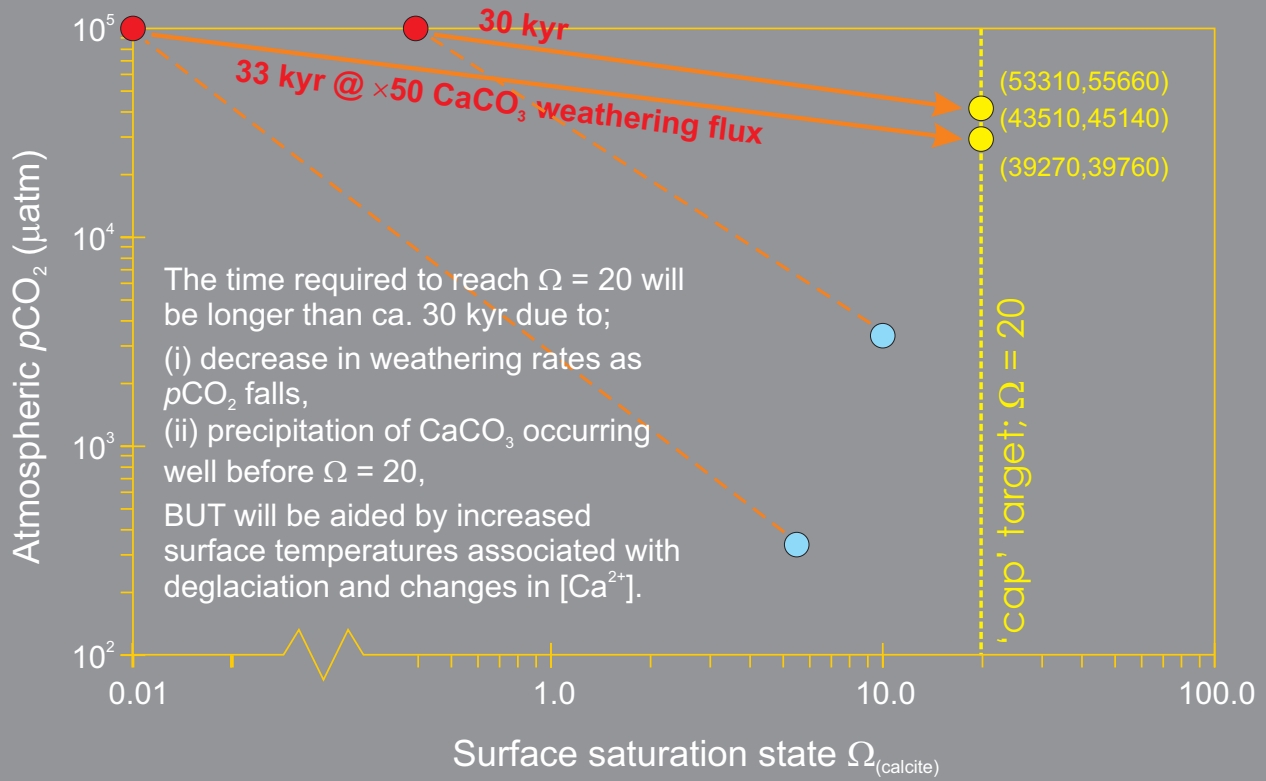
adapted from; Hoffman and Schrag [2002]



NOTE: ocean composition format;
[mean alkalinity, mean DIC] ($\mu\text{mol kg}^{-1}$)

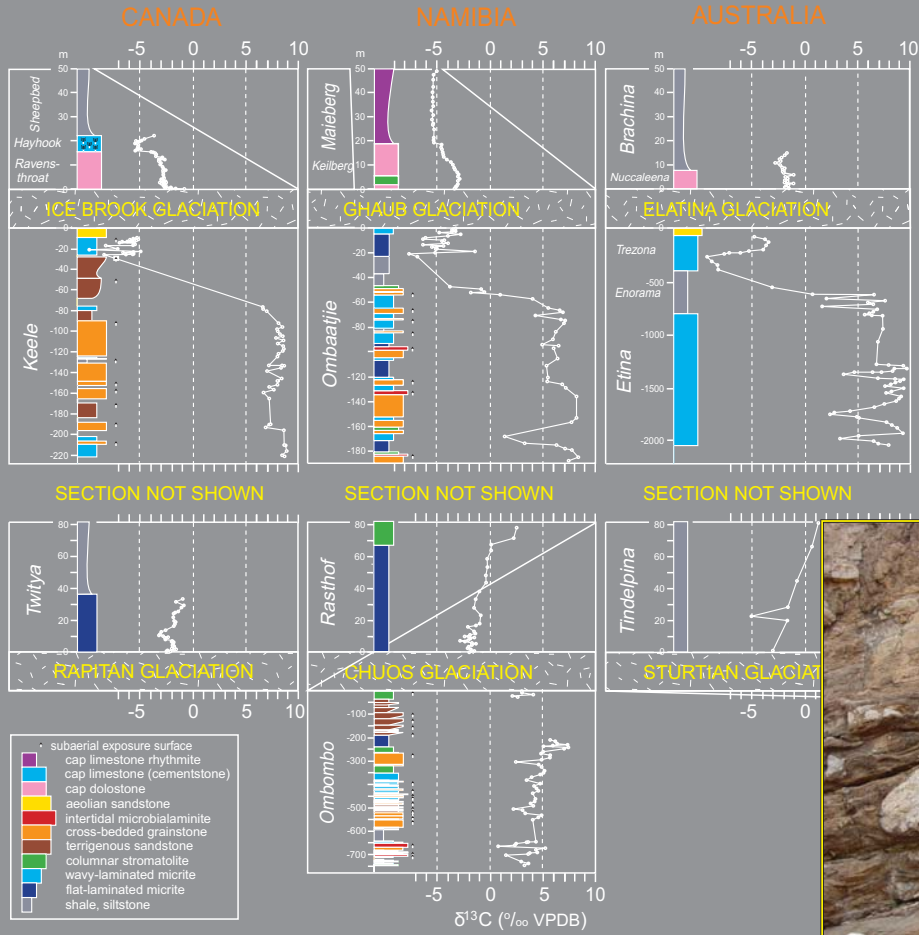


NOTE: ocean composition format;
[mean alkalinity, mean DIC] ($\mu mol kg^{-1}$)



NOTE: ocean composition format;
[mean alkalinity, mean DIC] ($\mu\text{mol kg}^{-1}$)

The enigma of the 'cap carbonates'



From: Hoffman and Schrag [2002]

